

→ POLINSAR 2015

# PRECIPITATING CLOUD EFFECTS ON THE RADAR POLARIMETRIC SIGNATURE AT KA BAND

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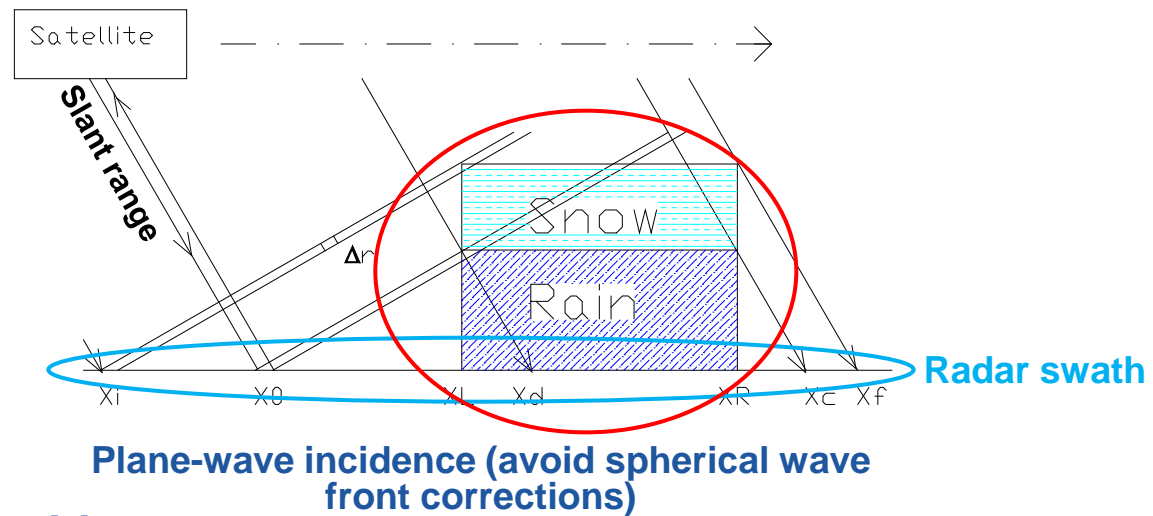
# WHY CLOUDS DEALING WITH SAR ?



- **POTENTIAL of X/Ku/Ka-band SAR for RAIN retrieval**
  - At higher frequency precipitating clouds may produce significant **attenuation/scattering/depolarization** effects
  - The high spatial resolution of SAR sensors might provide new insights into the **structure of precipitating clouds** from space.
  - **Large availability of a new generation of X-SAR satellites** near fully polarimetric
- **Atmospheric artifacts on high frequency radar imaging**
  - **Rainfall signatures** have been already revealed by previous X-SARs measurements (e.g. SAR-X SIR-C in 1994)
  - It can affect the interpretation of SAR data at X and higher frequencies, **modifying the polarimetric signature** of the ground
  - Assessing these effect can support to the **design and performance assessment** of future high frequency radar (e.g., Ka band interferometer, ESA funded project N. 4000109477/13/nl/lvh)

- **Introduction**
  - Context and examples
- **Modeling of SAR observations due to precipitation**
  - SAR response model
  - Polarimetric SAR observables
  - High-resolution polarimetric simulated scenarios
- **Polarimetric signature of precipitation**
  - Sensitivity analysis
  - Polarimetric frequency diversity
- **Parametric estimation of precipitation from X-SAR**
  - Case study
- **Conclusions**

# SAR cloud response and observables



## SAR polarimetric observables

$$\sigma^{\circ}_{SAR}(x) = \sigma^{\circ}_{SRF}(x) + \sigma^{\circ}_{VOL}(x)$$

$$Z_{SARco}(x) = \frac{\sigma^0_{SARhh}(x)}{\sigma^0_{SARvv}(x)}$$

$$\rho_{SARco}(x) = \frac{\langle S_{SARhh}(x) S_{SARvv}^*(x) \rangle}{\sqrt{\langle |S_{SARhh}(x)|^2 \rangle} \sqrt{\langle |S_{SARvv}(x)|^2 \rangle}}$$

$$= |\rho_{SARco}(x)| e^{j\Psi_{SARco}(x)}$$

- $\sigma^0_{SARpq}$ :  **$pq$ -polarized normalized radar cross section (NRCS)**

- $Z_{SARco}$ : **co-polar ratio**

- $\rho_{SARco}$ : **complex correlation coefficient**



# Modelling SAR response: NRCS



- For a given pixel  $(x,y)$  the SAR NRCS can be formally expressed as follows:

$$\sigma_{SARpq}^0(x,y) = \sigma_{SRFpq}^0(x,y) + \sigma_{VOLpq}^0(x,y)$$

- $\sigma_{SRFpq}^0(x,y)$ : surface backscatter, attenuated by the two-way path through the precipitating atmosphere
- $\sigma_{VOLpq}^0(x,y)$ : volume backscattering due to hydrometeor reflectivity, weighted by the two-way path through precipitating atmosphere

$$\sigma_{SRFpq}^0(x,y) = \sigma_{pq}^{ground}(x,y) \exp \left( - \int_{\Delta l(x,y)} k_{pp}(l) dl - \int_{\Delta l(x,y)} k_{qq}(l) dl \right)$$

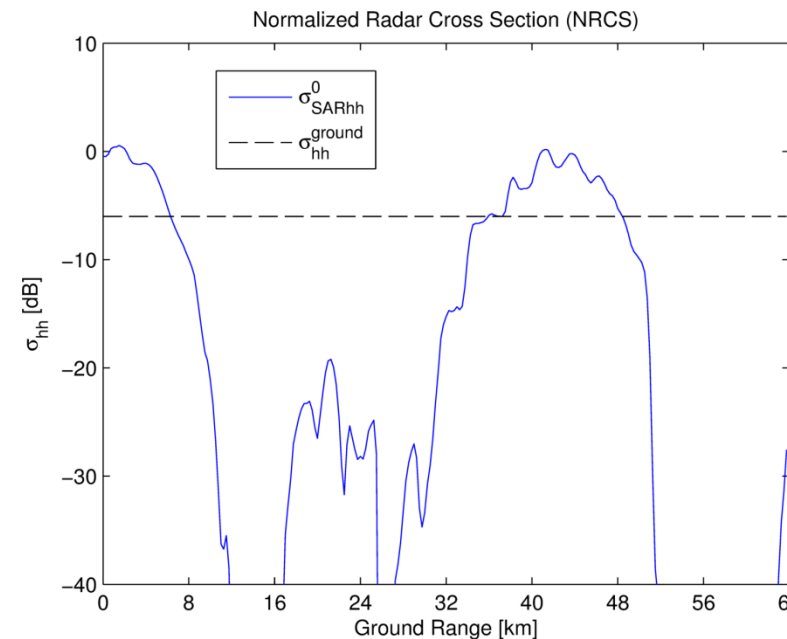
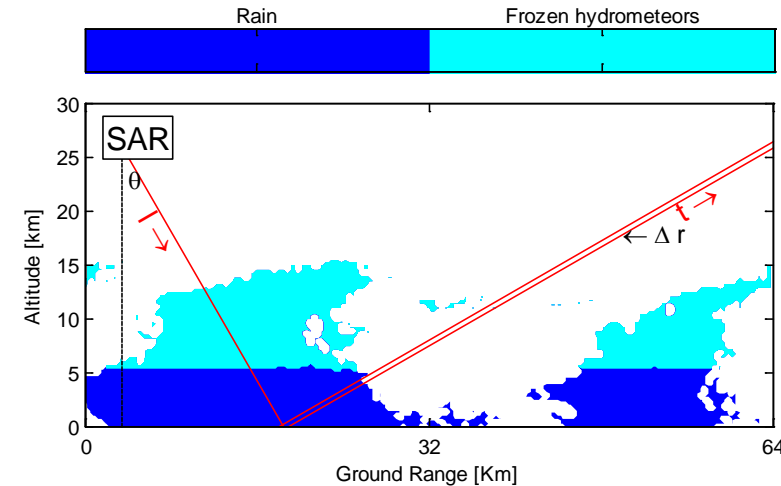
$$\sigma_{VOLpq}^0(x,y) = \int_{\Delta t(x,y)} \eta_{pq}(t) \exp \left( - \int_{\Delta l(x,y)} k_{pp}(l) dl - \int_{\Delta l(x,y)} k_{qq}(l) dl \right) dt$$

- $\sigma_{pq}^{ground}$ : surface target NRCS
- $\eta_{pq}, k_{pq}$ : hydrometeors reflectivity and specific attenuation

$$\eta_{pq} = 4\pi \langle |S_{pq}|^2 \rangle = \int_0^\infty \int_0^{\pi/2} 8\pi^2 |S_{pq}(D,\phi)|^2 N(D) p(\phi) dD \sin\phi d\phi$$

$$k_{pq} = -2\lambda \langle \text{Im}(F_{pq}) \rangle$$

- $S_{pq}, F_{pq}$ : element of the complex back or forward hydrometeor scattering matrix
- $N(D)$ : particle size distribution
- $p(\phi)$ : particle orientation probability density function
- $\lambda$ : wavelength



# Modelling SAR response: Correlation Coeff.



- For a ground point (x,y) the observable SAR complex correlation coefficient is given by:

$$\rho_{SARco} = \frac{\sqrt{\sigma_{hh}^0} \sqrt{\sigma_{vv}^0} \rho_{co}^{ground} e^{-\int_{\Delta l(x)} k_{hh}(l) dl} e^{-\int_{\Delta l(x)} k_{vv}(l) dl} e^{j2\Phi_{co}(x)} + \sin\theta \int_{\Delta l(x)} C_{VOL}(t) dt}{\sqrt{\sigma_{SARhh}^0} \sqrt{\sigma_{SARvv}^0}}$$

$$\Psi_{SARco}(x, y) = \arg\{\rho_{SARco}(x, y)\}$$

$$C_{VOL}(t) = \sqrt{\eta(t)_{vv}} \sqrt{\eta(t)_{hh}} \rho_{co}^{vol}(t) \exp\left(-\int_{\Delta l(t)} k_{vv}(l) dl\right) \exp\left(-\int_{\Delta l(t)} k_{hh}(l) dl\right) e^{j2\Phi_{co}(t)}$$

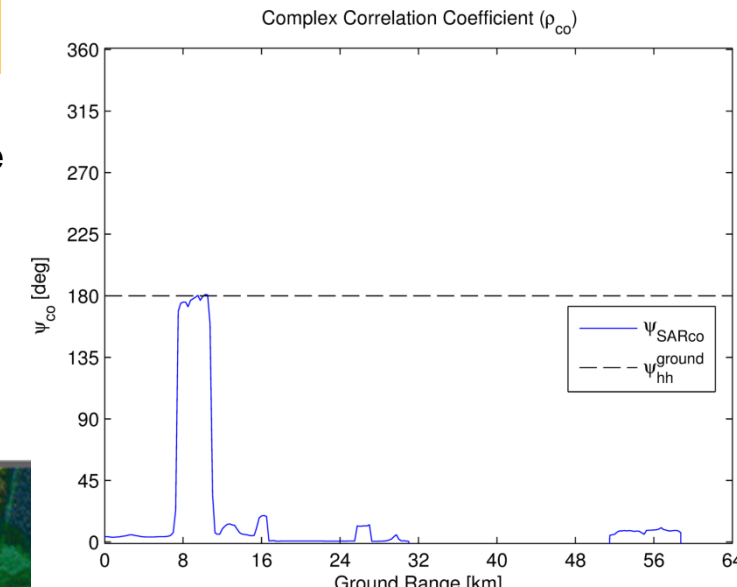
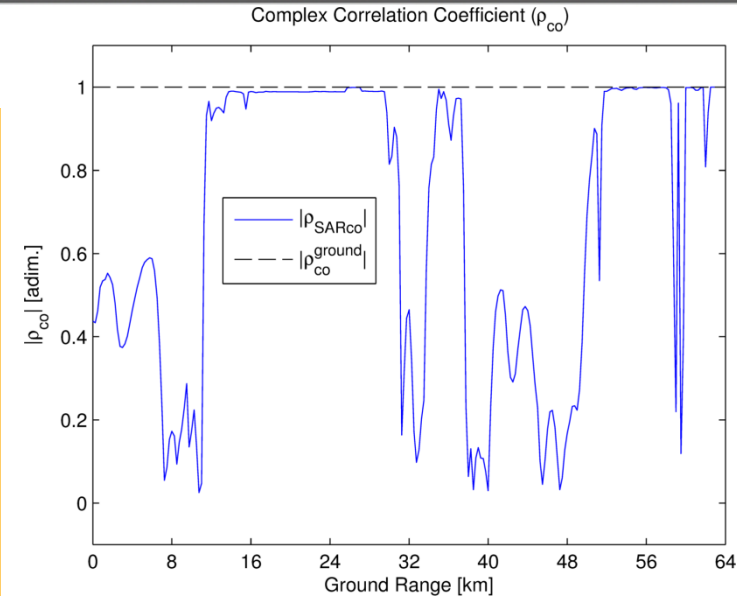
$$\rho_{co} = |\rho_{co}| e^{j\delta_{co}}$$

$$\delta_{co} = \delta_{vv} - \delta_{hh} = \arg(\rho_{co})$$

$$K_{co} = \lambda < \text{Re}(F_{hh} - F_{vv}) >$$

$$\Phi_{co} = \int_{\Delta l(t)} K_{co}(l) dl$$

- $\rho_{SARco}$ ,  $\rho_{co}^{ground}$ ,  $\rho_{co}^{vol}$ : complex correlation coefficients of observed resolution cell, surface target and volume bin
- $\delta_{co}$ : backscatter differential phase
- $S_{pq}$ ,  $F_{pq}$ : elements of the hydrometeoros complex back or forward scattering matrix
- $\eta_{pq}$ : hydrometeor reflectivity
- $K_{co}$ : hydrometeor copolar specific differential phase



# Hydrometeor e.m. parameterization



Hydrometeors polarimetric parameters can be modelled as function of water content by mean of power laws:

$$Z_{epq}(x, y, z) = \frac{\lambda^4}{\pi^5 |K|^2} \eta_{pq}(x, y, z) = a_{Zpq} W(x, y, z)^{b_{Zpq}}$$

$$k_{pq}(x, y, z) = a_{kpq} W(x, y, z)^{b_{kpq}}$$

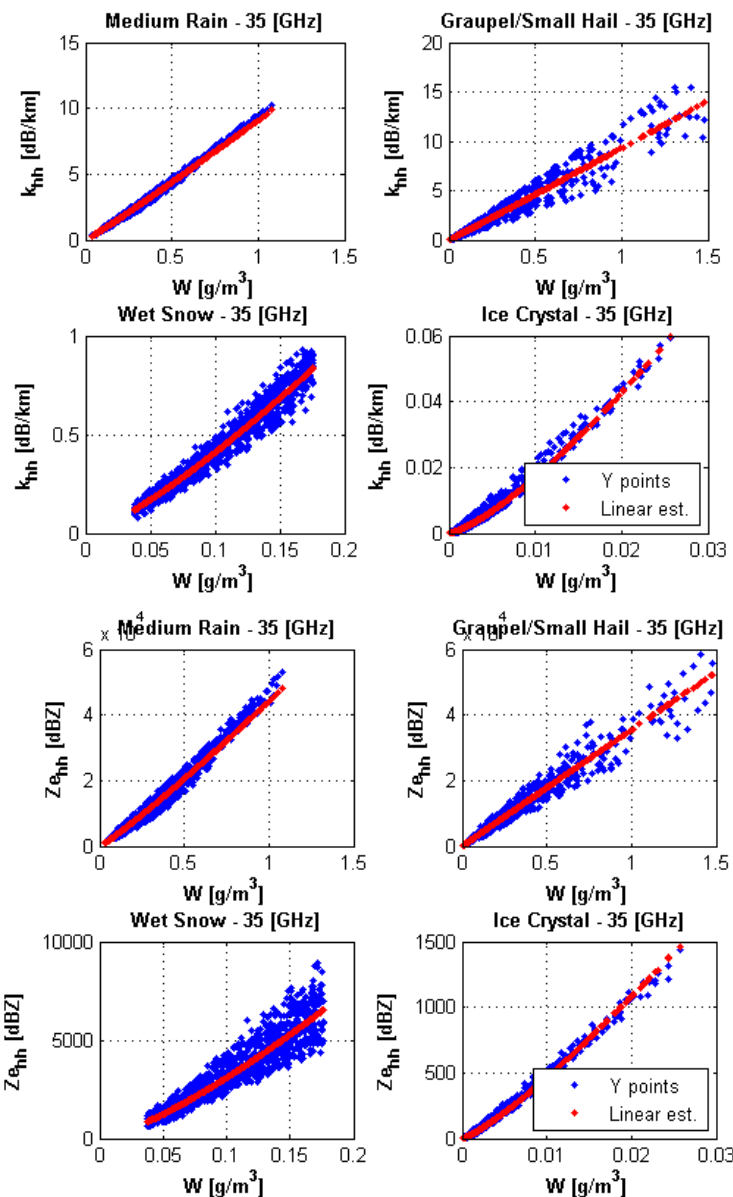
$$K_{co}(x, y, z) = a_{Kco} W(x, y, z)^{b_{Kco}}$$

$$|\rho_{co}|(x, y, z) = a_{\rho co} W(x, y, z)^{b_{\rho co}}$$

$$\delta_{co}(x, y, z) = a_{\delta co} W(x, y, z)^{b_{\delta co}}$$

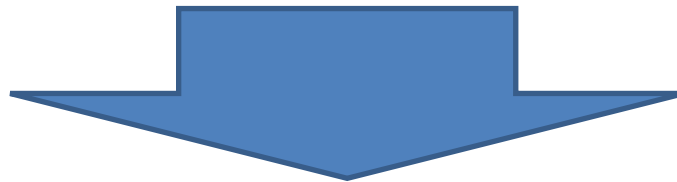
- $W$  = water content [g/m<sup>3</sup>]
- $K_{co}$  = differential phase. [°/km]
- $k_{pq}$  = specific attenuation [dB/km]
- $Z_{epq}$  = equivalent reflectivity [mm<sup>6</sup>/m<sup>3</sup>]
- $|\rho_{co}|$  = mod. of the copolar corr. Coefficient
- $\delta_{co}$  = arg. of the copolar corr. coefficient
- $\lambda$  = wavelength [cm]
- $|K|^2$  = 0.93 for water and 0.19 for ice

$a_{Xpq}$ ,  $b_{Xpq}$  coefficients have been obtained by using **HESS T-Matrix radar scattering model**, as described in [Marzano et al., 2007] and [Marzano et al., 2010].





- The proposed SAR response model requires as input a
  - **cloud structure** (2D geometry, hydrometeor types, hydrometeor water content distribution, hydrometeor e.m. response parameterization) and
  - the polarimetric characterization of the **ground target**



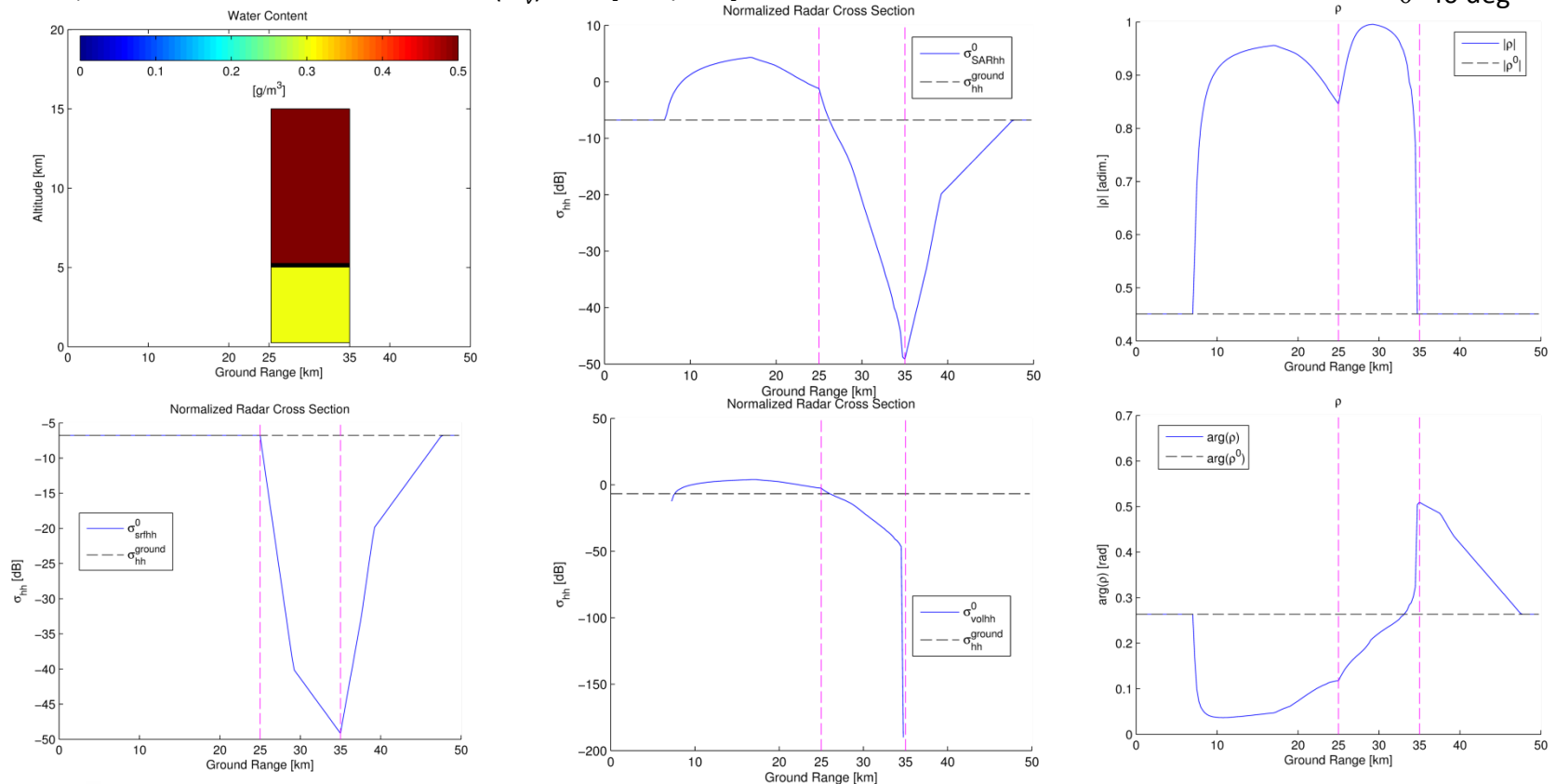
- The **ground target** polarimetric covariance matrix can be given by:
  - **Models** (e.g., bare soil polarimetric response by [Oh et al., 2002])
  - Polarimetric signature of **canonical targets** (i.e., spheres with  $\rho_{co}^{ground}=1$ , dihedrals with  $\rho_{co}^{ground}=-1$ , others)
- The **cloud structure** (i.e., hydrometeors distributions) can be derived by
  - ad-hoc **synthetic distributions**, with simplified shape (e.g. parallelepiped), to assess main effects in a simple environment.
  - Realistic fields simulated by a **3-D high-resolution mesoscale** cloud-resolving models (CRMs)



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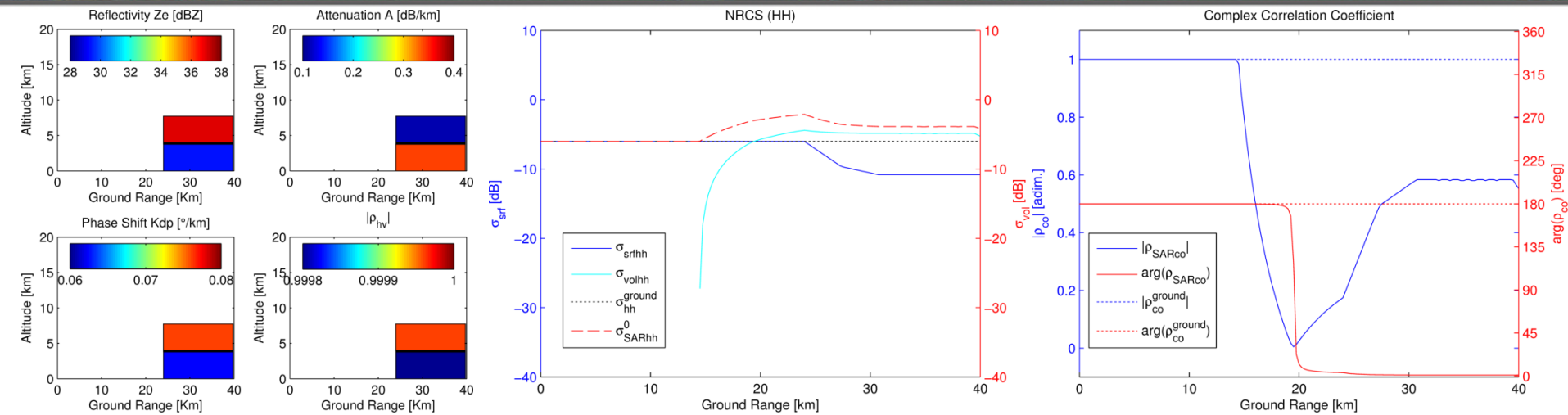
# Synthetic cloud 1: Rectangular cumulonimbus

- Parallelepiped structure
  - Horizontal uniform 10 km wide hydrometeor density
  - 2 vertical layers: raindrops  $0.3 \text{ g/m}^3$  and snowflakes  $0.5 \text{ g/m}^3$  water content, with 15 km overall height
  - E. M. parameters from HESS T-Matrix model [Marzano et al., 2007]
  - Homogenous background from bare soil model [Oh et al., 2002]: Rms Height ( $k_s$ ) 1.5 cm, Correlation Length ( $k_l$ ) 5.0 cm, Volumetric soil moisture content ( $m_v$ )  $0.25 \text{ [cm}^3/\text{cm}^3]$
- Zero °C @4.9 km
  - $\sigma^{\text{ground}} \cong -6 \text{ dB}$
  - $f = 35 \text{ GHz}$
  - $\theta = 40 \text{ deg}$

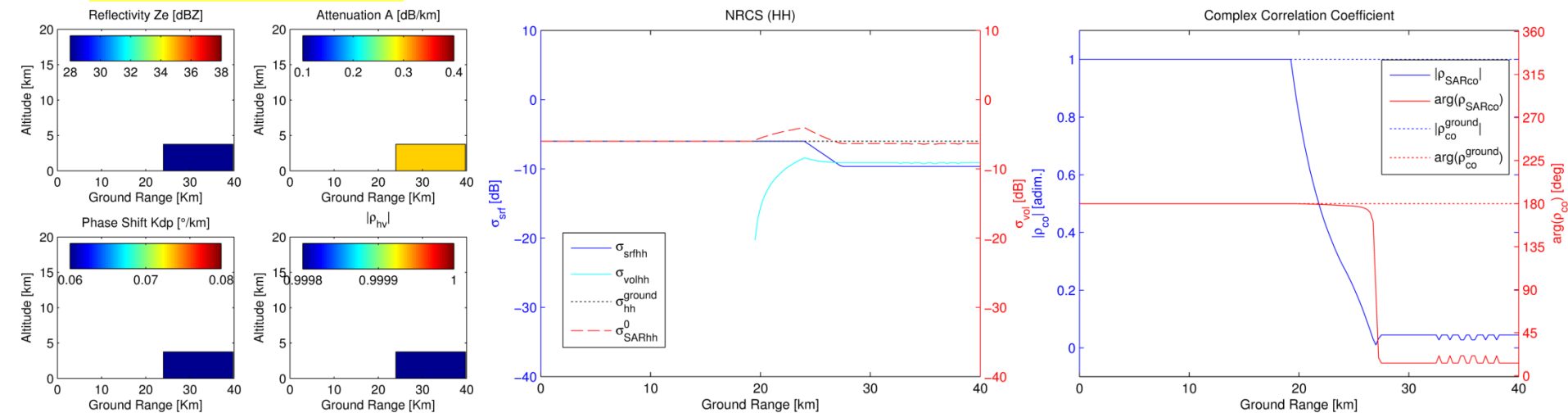




# Case 2: Stratiform cloud with and without ice



## Dihedrals – NRCS c.a. -6dB

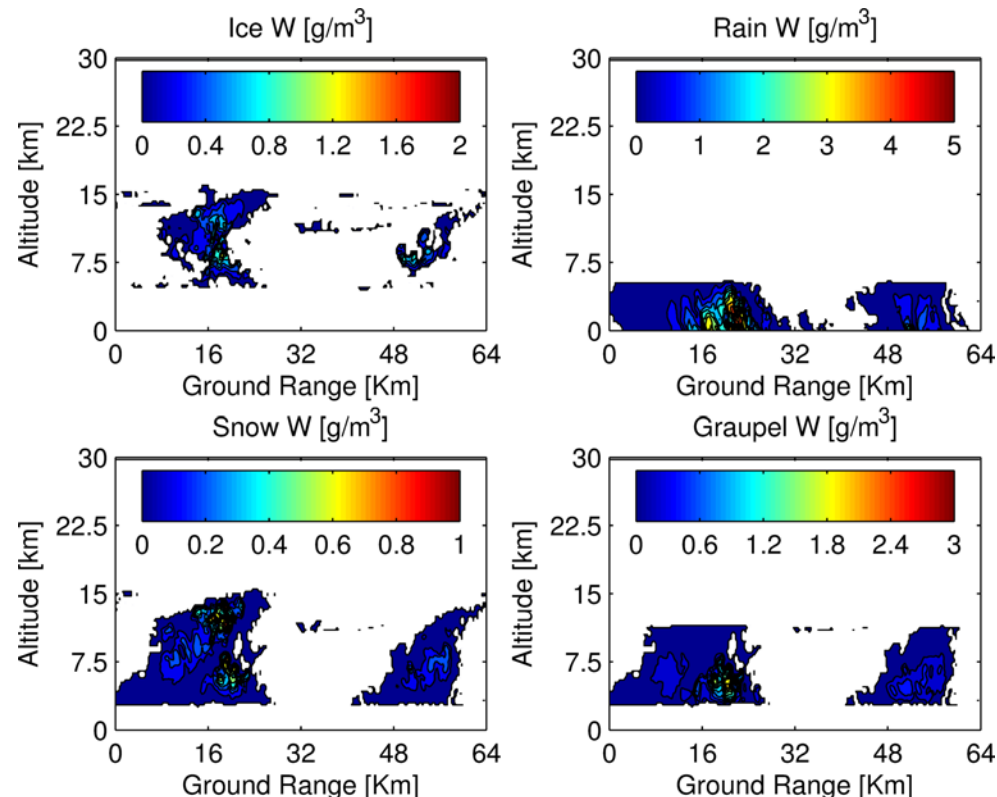


# High-resolution realistic clouds



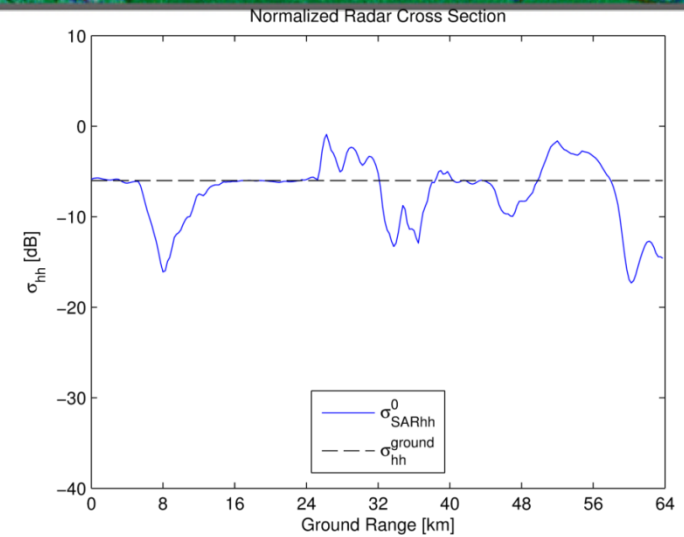
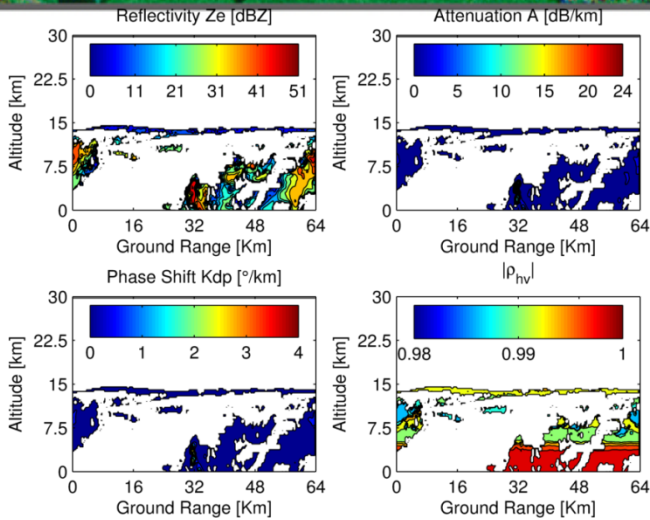
- The simulated **hydrometeors distributions** can be derived by simulated field through **3-D high-resolution mesoscale cloud-resolving models (CRMs)**
- The **System for Atmospheric Modeling (SAM) CRM** [Blossey et al., 2007] allows at simulating the distribution of Cloud, Rain, Ice, Snow, Graupel particles at 250 m resolution

Examples of horizontal and vertical distribution of densities of different hydrometeors predicted by model at a given epoch

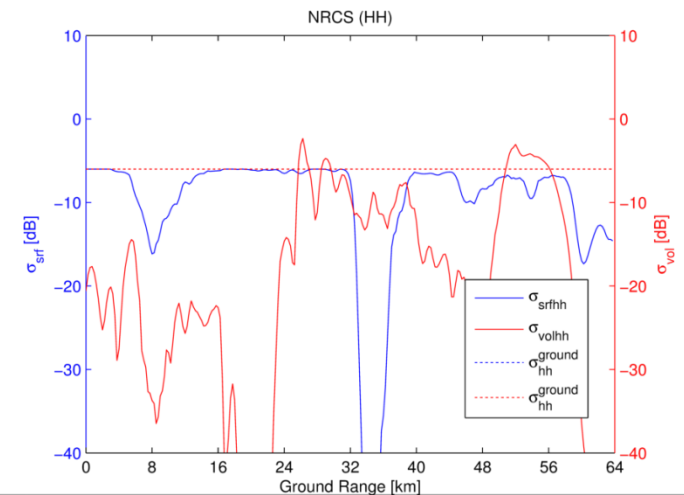
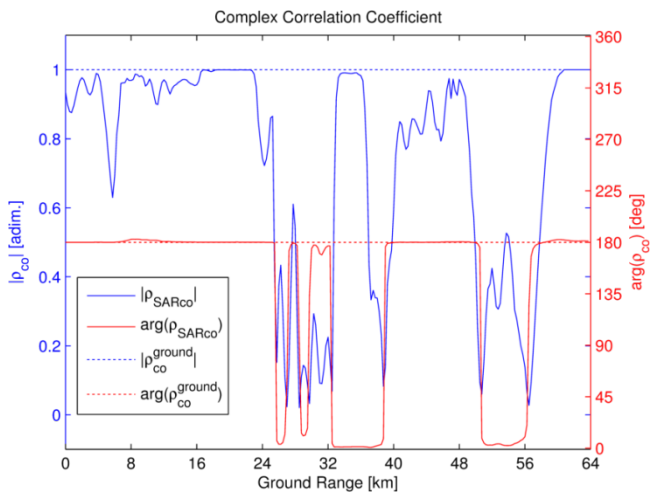




# Case 3. Realistic spread thin cloud



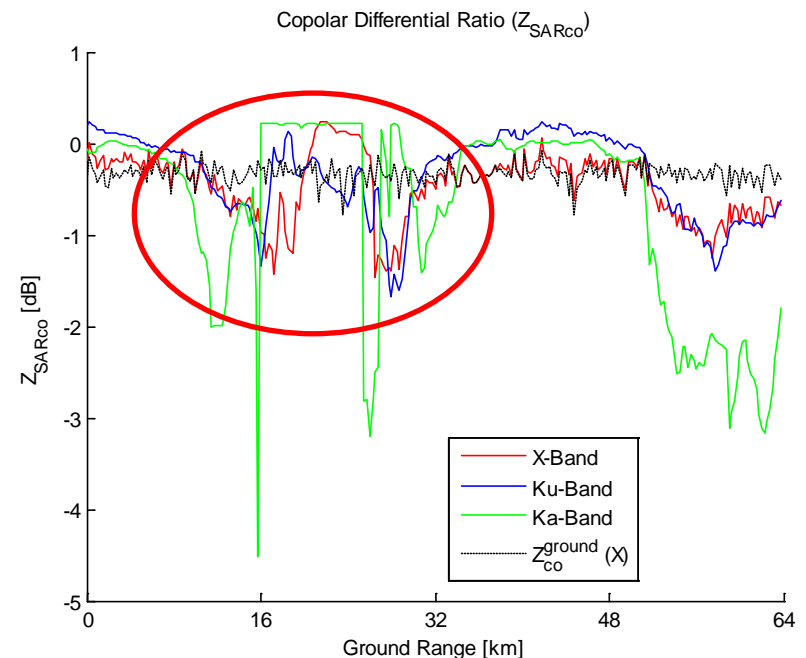
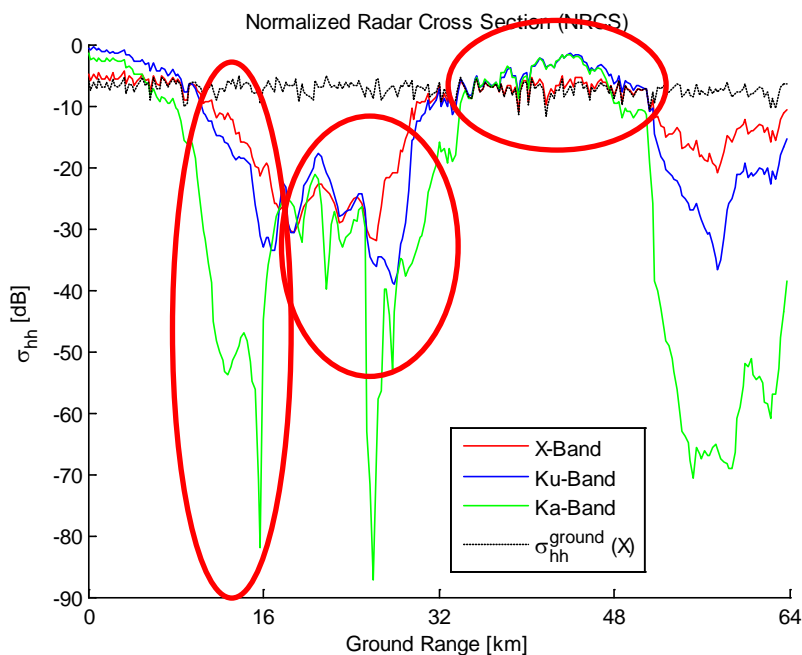
Dihedrals – NRCS c.a. -6dB



# Dependence on frequency



- **Both negative and positive peaks of NRCS increase with frequency**
  - Ka-band NRCS appears more sensitive to attenuation and less to reflectivity
  - X and Ku bands show a comparable dynamic range
- **Presence of dissimilar behavior**
  - Peak shifting, peak present only in one or two frequencies
- **$Z_{SARco}$  similar to NRCS**
  - Small positive values (presence of spherically-shaped frozen hydrometeors)
  - Large negative peaks (presence of intense precipitation with oblate raindrops)
  - Ka-band quite dissimilar from X and Ku ones

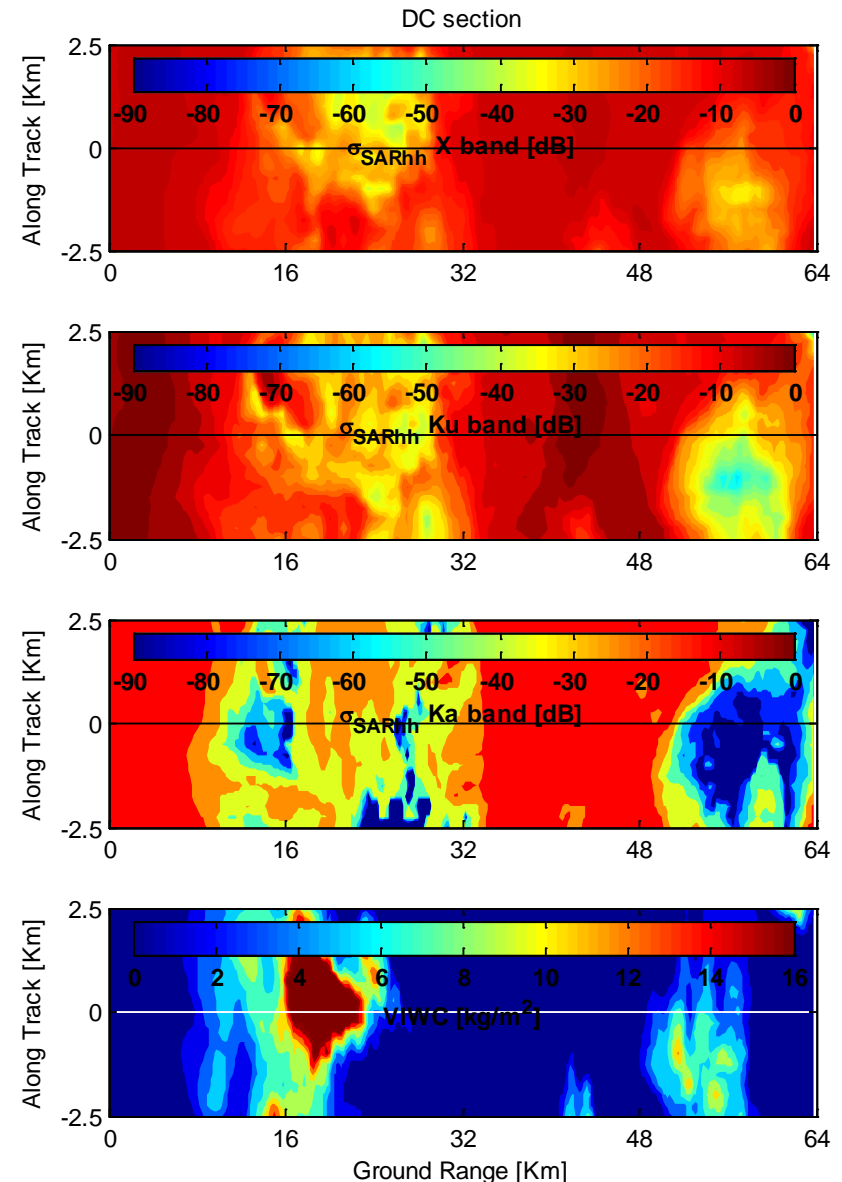




# Simulates NRCS image at X/Ku/Ka bands



- Ground plane  $\sigma^0_{SARhh}$  for X, Ku and Ka band,  $hh$  polarization plus the total vertically-integrated columnar content (VIWC).
- The ground response has been considered as constant (about -7 dB at X-band), and so the incident angle ( $40^\circ$ ).
- The images of the storm are simulated in ground range (the x-y plane), 5 km width, spread all the cross range dimension (64 km).



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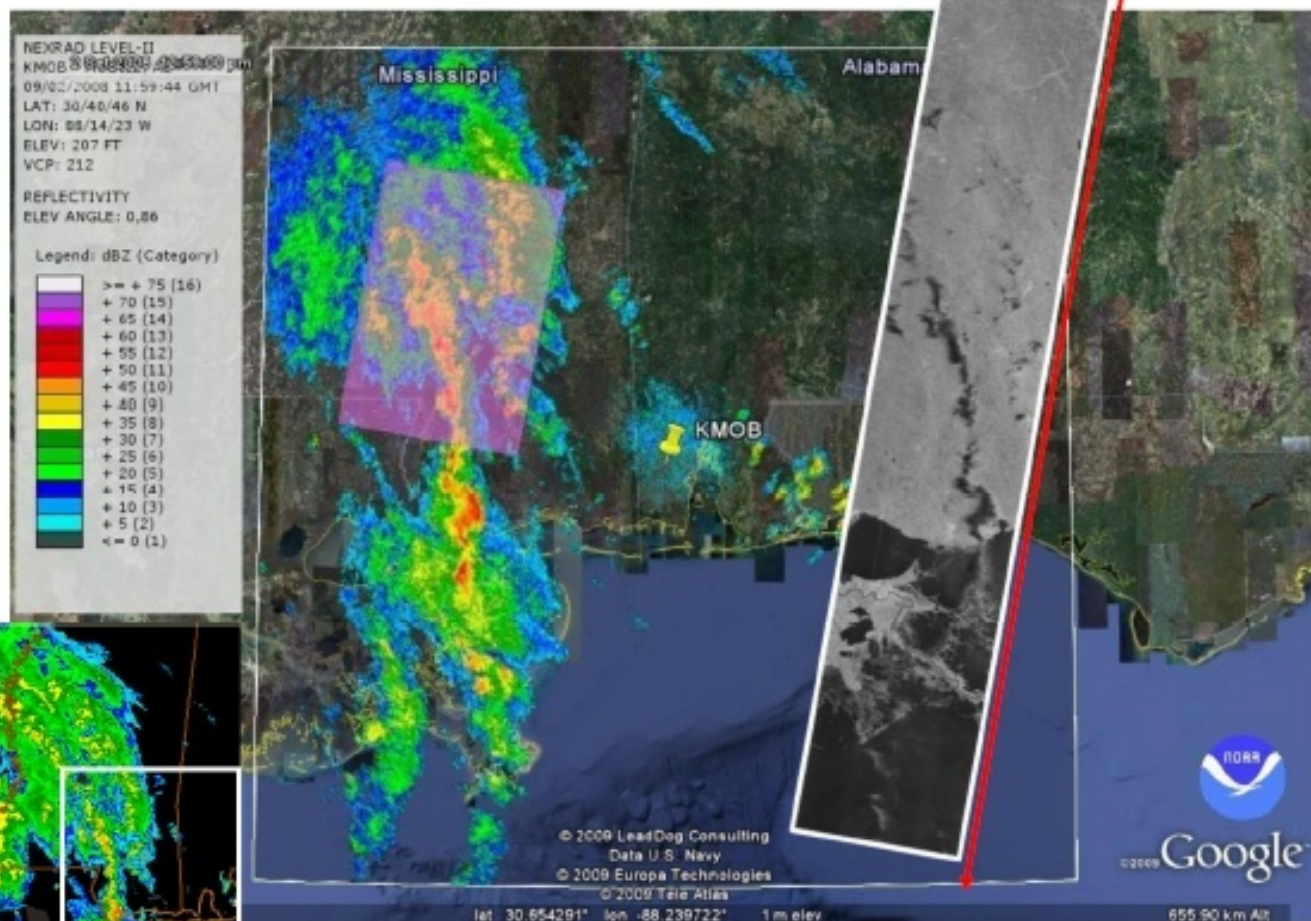
# X-SAR retrievals: Hurricane "GUSTAV" case



- South eastern Louisiana around  $30.5^{\circ}$  N x  $89.5^{\circ}$  W
- September 2, 2008 at 12:00 UTC

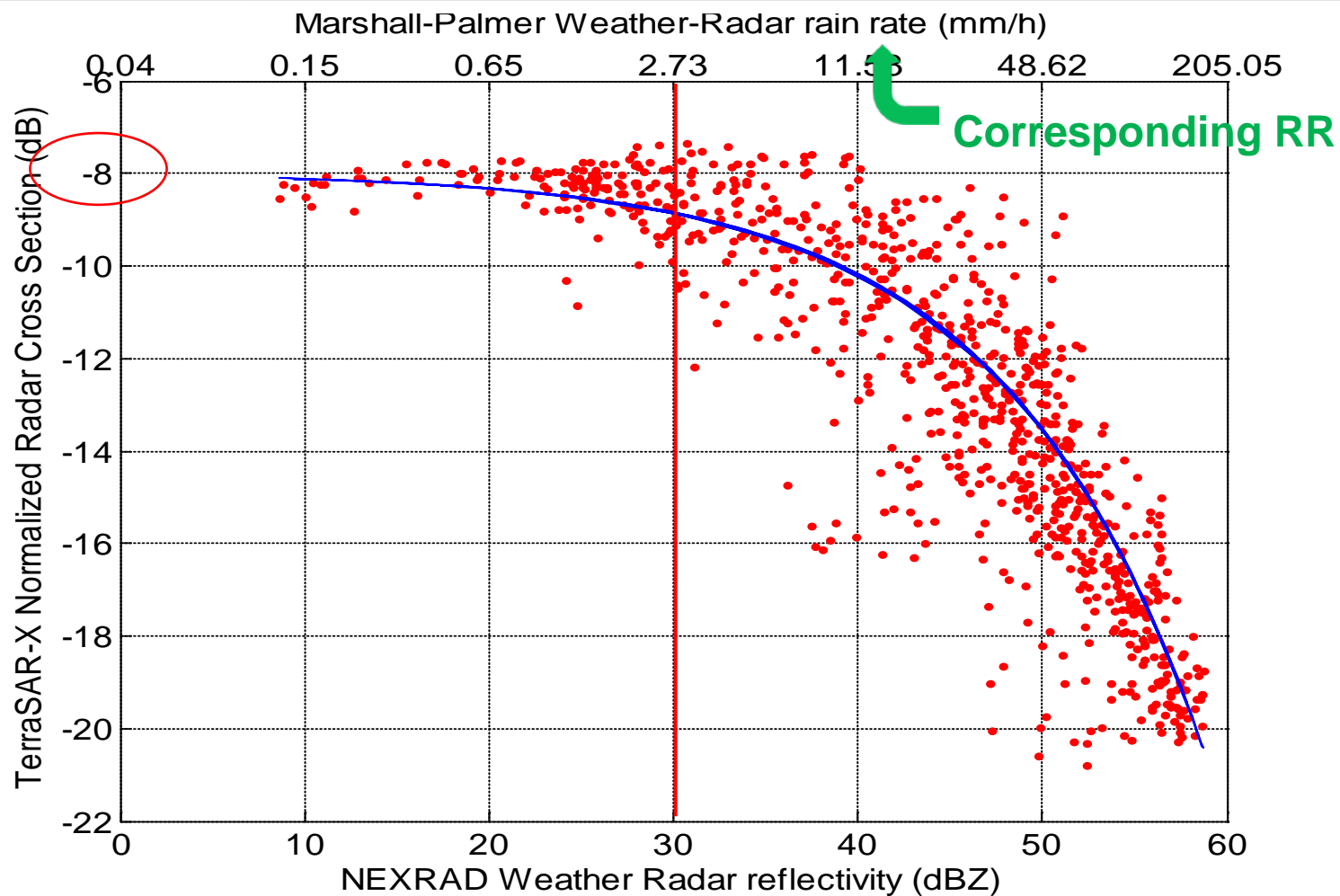
TerraSAR-X:  
HH pol data

NEXRAD:  
S-band reflectivity  
PPI at  $0.86^{\circ}$   
KMOB site



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# X-SAR retrievals : Hurricane "GUSTAV" case



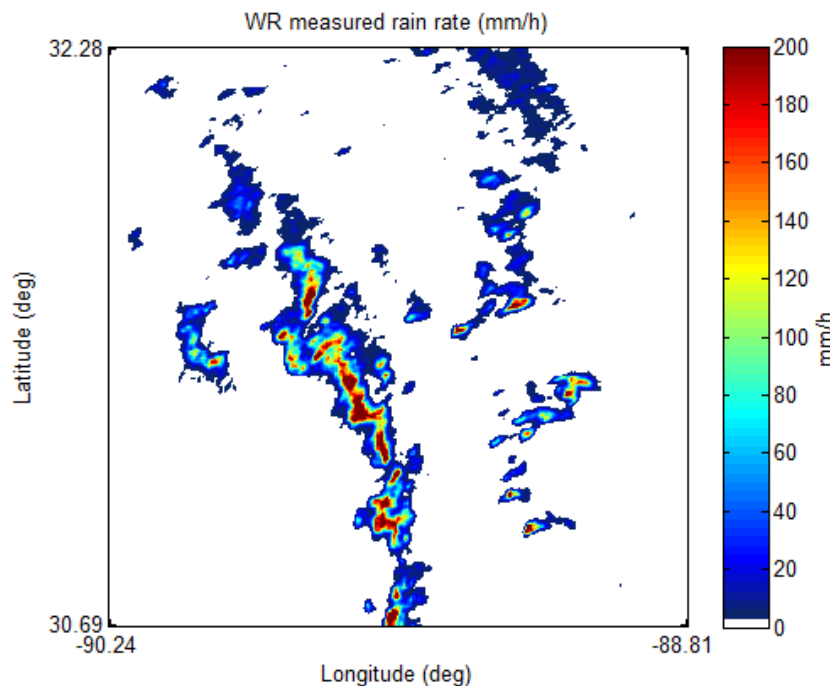
**Correlation between NRCS values  $\sigma_{SAR}$  against co-located and co-registered NEXRAD WR reflectivity  $Z$ , for the selected region of interest (ROI)**



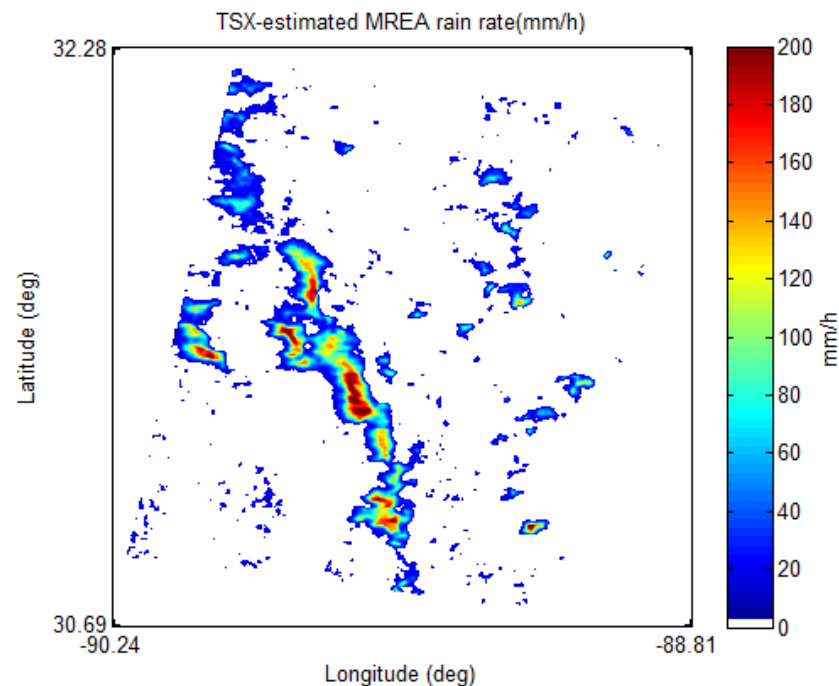
# X-SAR retrievals: Hurricane "GUSTAV" case



## Precipitation rate [mm/h]



Ground Weather Radar



Spaceborne X-SAR

- Corr = 0.75
- Bias = -0.66 mm/h
- RMSE = 22.28 mm/h
- FRMSE = 0.98

# Conclusions



- A 2D/3D simulator of the polarimetric response of a SAR in presence of precipitating clouds has been developed
- Considering simple synthetic clouds as well as realistic clouds predicted by a mesoscale model we found that:
  - **Bad news**
    - The impact of clouds on  $\alpha$  and especially Ka band SAR can be significant, even for clouds with relatively high occurrence
    - The polarimetric signature of the ground target can be significantly modified or even completely masked
  - **Good news**
    - Both copolar ratio and differential phase exhibit a good correlation with columnar contents.
    - SAR remote sensing of clouds has spatial resolution useful for water budget, water management and hydrological model initialization

## □ Acknowledgments

- This work has been partially funded by:
  - ESA/ESTEC contract N. 4000109477/13/nl/lvh) “Ka-Band SAR Backscatter Analysis in Support of Future Applications”
  - European Union project Earth2Observe “Global Earth Observation for Integrated Water Resource Assessment” (<http://www.earth2observe.eu>)

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Marzano et al., HESS, 2011  
Marzano et al., TGRS, 2012  
Pulvirenti et al., TGRS 2014