



Tropical drivers of the Atmospheric Growth Rate 2011-2010





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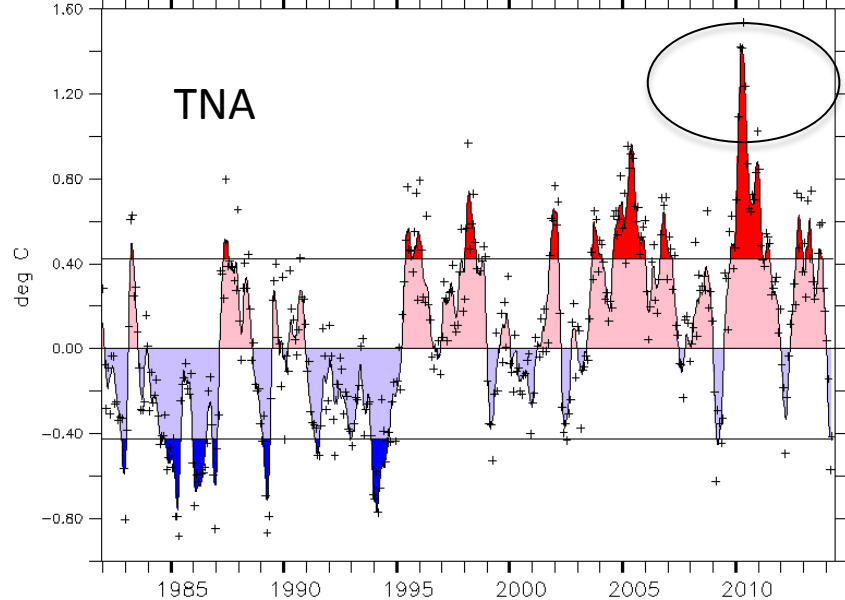
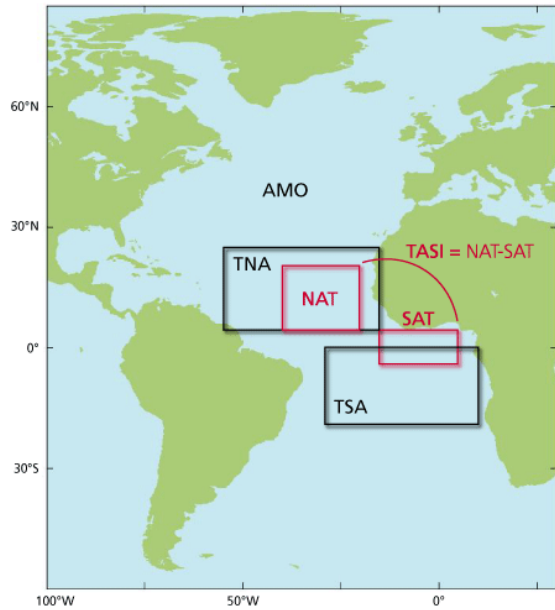
Kevin Gurney

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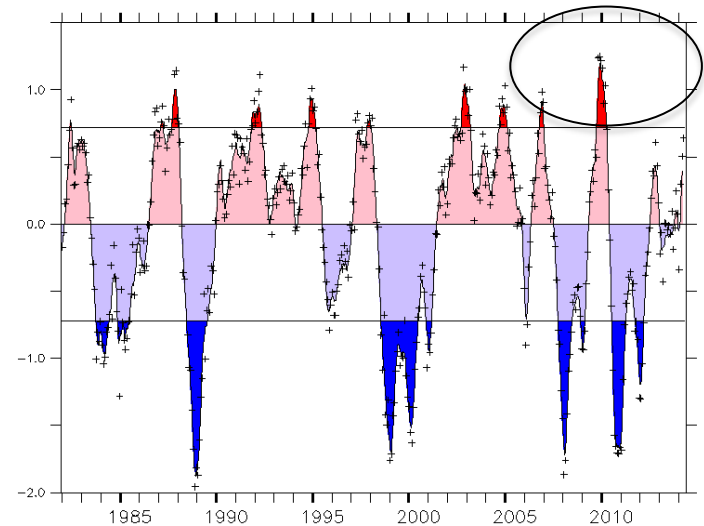
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University of Colorado, Boulder



Drought and the Tropical North Atlantic



El Nino 4.0



The tropical North Atlantic experienced a record-breaking SST anomaly $>1.5^{\circ}\text{C}$.

Consequence of a strong Central-Pacific “Modoki” El Nino amplified by a strong and persistent negative phase of the North Atlantic Oscillation (NAO) (Hu *et al*, J. Climate 2011)

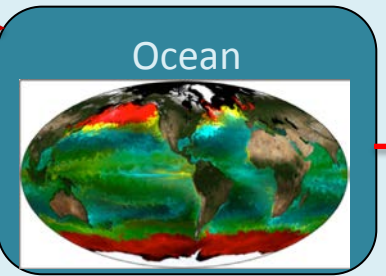
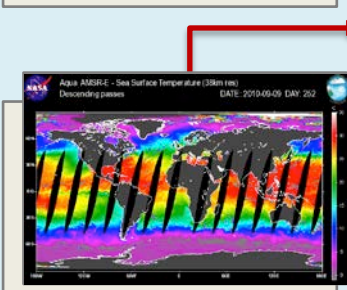
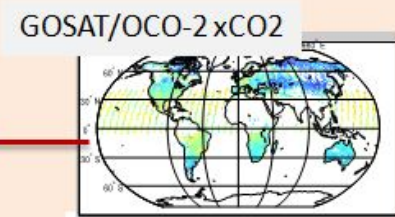
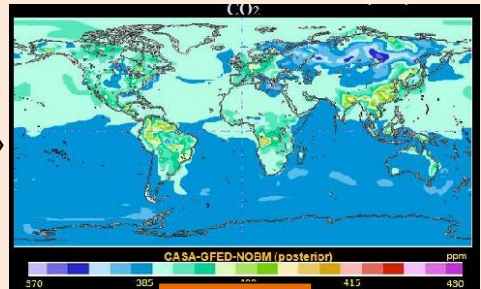
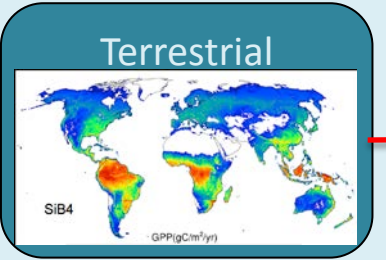
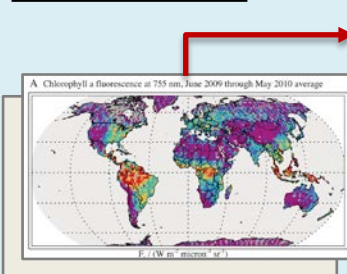
Shifted ITCZ north leading to a dramatic Amazonian drought (Lewis *et al*, Science, 2010)

Surface Satellite Data

Carbon Cycle Models

Transport Model

Atmospheric Satellite Data



Forecast

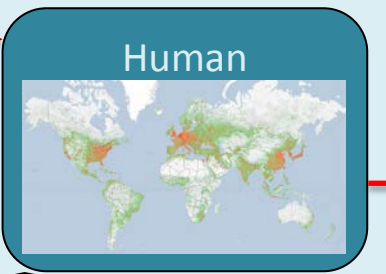
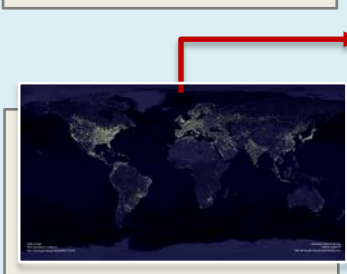
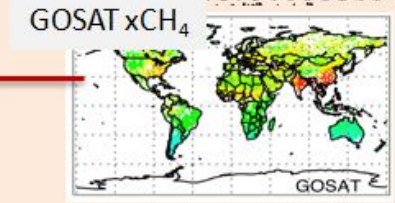
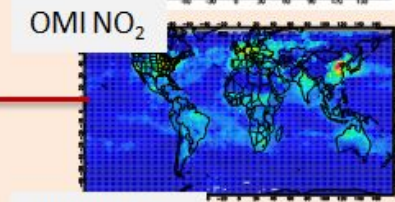
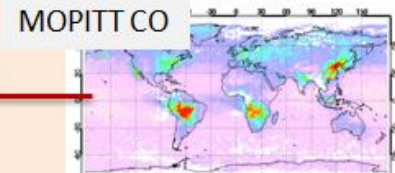
$$C_{\text{prior}}(\mathbf{x}) = \|\mathbf{x} - \mathbf{x}_a\|_{S_a}^2$$

$$C_{\text{obs}}(\mathbf{x}) = \|\mathbf{y} - \mathbf{F}(\mathbf{x})\|_{S_n}^2$$

$$\min_{\mathbf{x}} C(\mathbf{x}) = C_{\text{obs}}(\mathbf{x}) + C_{\text{prior}}(\mathbf{x})$$

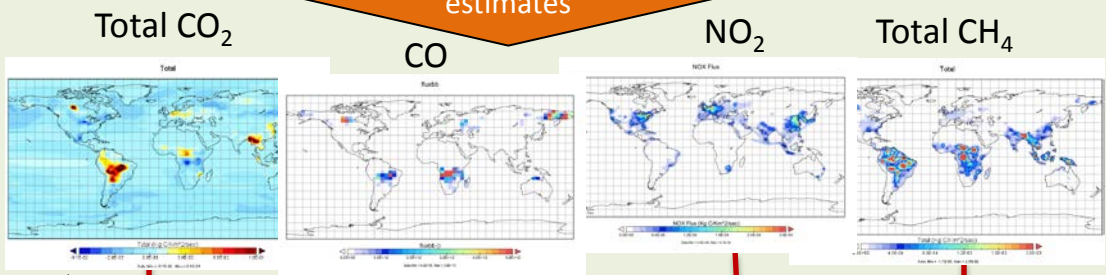
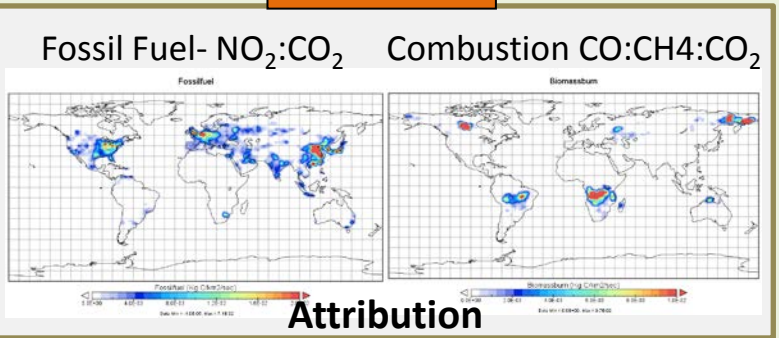
Inverse modeling

Observations



Reconciliation

Top-down estimates



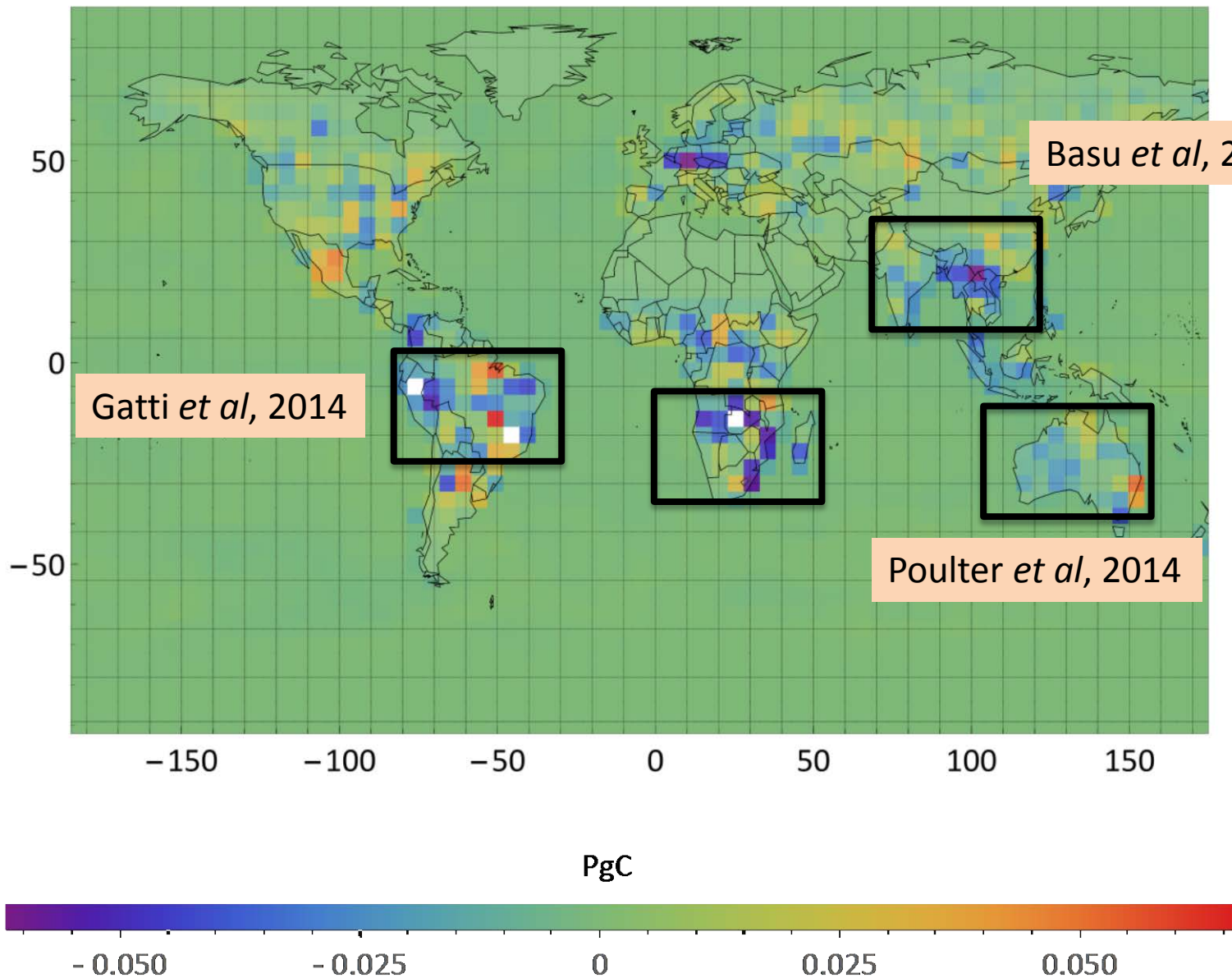
Posterior fluxes and uncertainties

Surface fluxes and uncertainties



Changes in total flux 2011-2010

Based upon ACOS v3.5



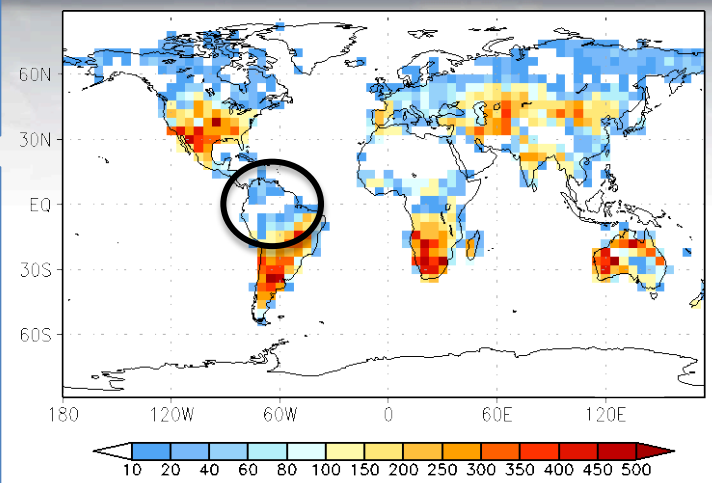


Cloud Canopy: Brazilian Challenge

Amazon is a critical but challenging region for satellite observations because of persistent cloud cover

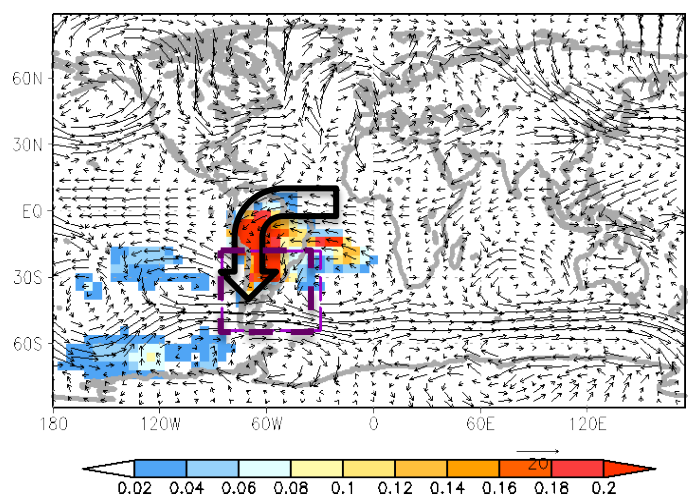
Adjoint analysis (Liu, Bowman, and Henze, JGR 2015) shows that mid-latitude S. America obs. can help infer Amazonian fluxes:

- The fluxes over Amazonia contribute more than 40% to the X_{CO_2} variability over mid-lat. S. America (rectangle area) due to the persistent wind direction (Fig. 2).
- The impact of GOSAT over mid-lat. S. America on Amazonia flux estimation is between 30% and 59% of full observation impact (right figure).

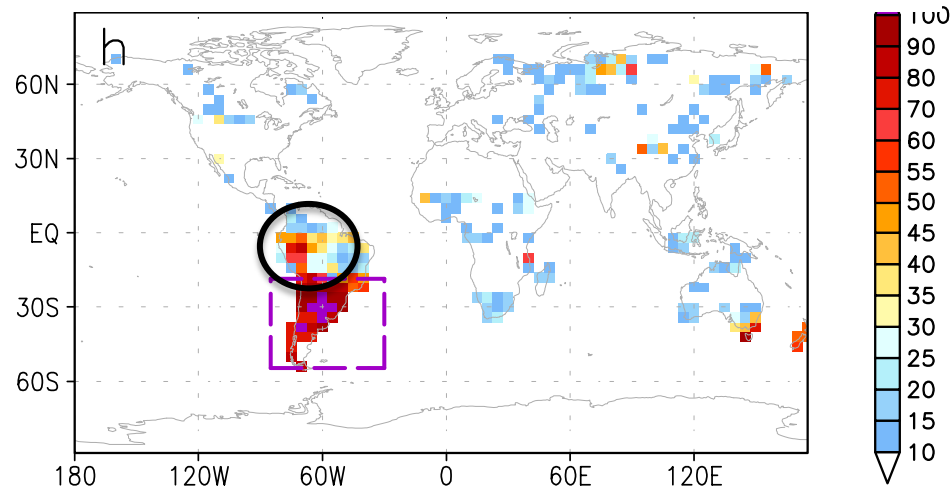


Annual total number of good quality nadir ACOS-GOSAT v3.5 obs at 4° x 5° resolution

700 hPa wind vector and sensitivity of X_{CO_2} in S-S- America (rectangle) at Jan 31 to surface fluxes on Jan 23rd



The lower bound impact of GOSAT X_{CO_2} within S-S- America (rectangle) on local and nonlocal flux estimation

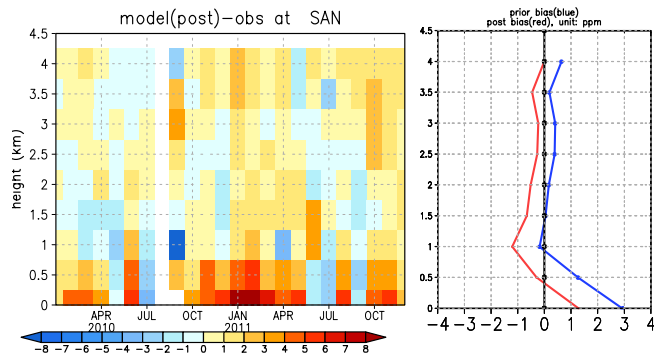
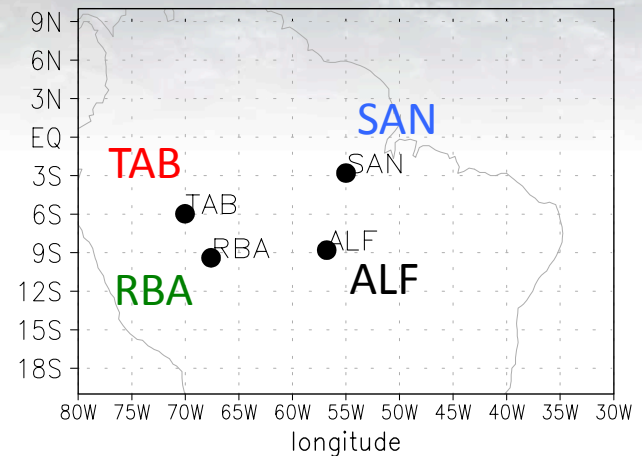




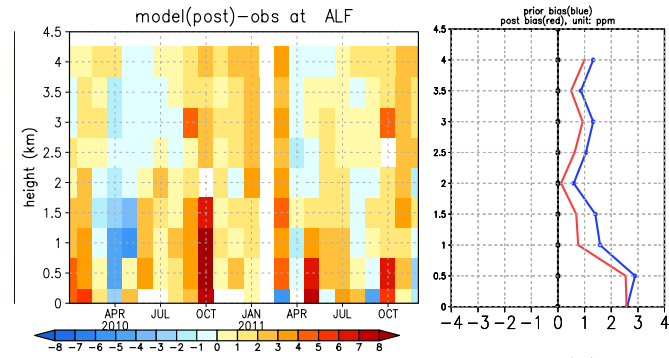
Comparison against aircraft obs.

- Aircraft spirals (from Gatti *et al.*) taken twice per month from 2010-2011.
- The mean posterior CO₂ bias is less than 1 ppm above 1 km.
- The bias in posterior CO₂ is smaller than prior CO₂ for TAB, ALF, RBA, but not SAN
- TAB region has strongest impact on midlatitude S. American obs.

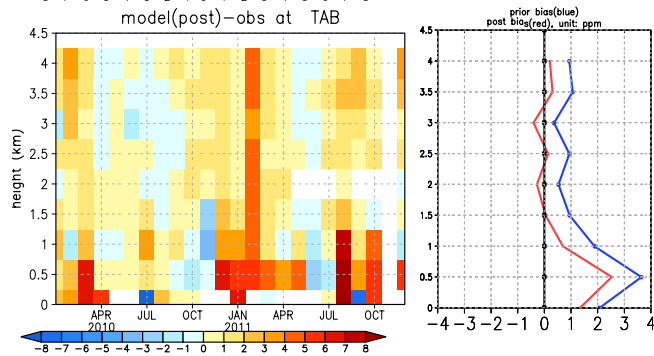
aircraft locations over Amazon



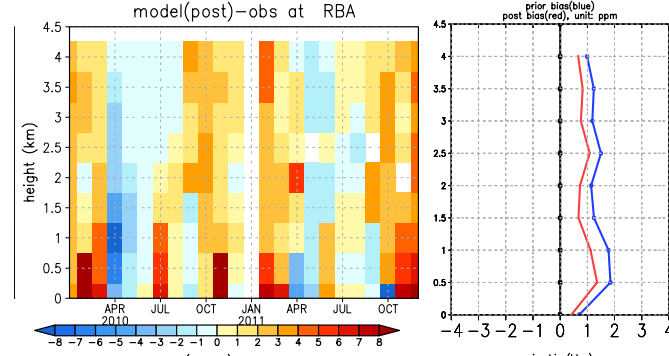
SAN



ALF



TAB



RBA



Flux Decomposition

Net flux into the atmosphere is positive

$$F^{\uparrow}(x, y, t) = F_F + F_O + F_{BB} - F_{NEP} + F_{chem}$$

Fossil Fuel Ocean Biomass burning NEP Chemical Source

The total flux inferred from CMS-Flux can be decomposed into a sum of terms representing key processes within the carbon cycle.

The interannual change in total flux can be similarly decomposed:

$$\delta F_{2011-2010}^{\uparrow}(x, y, t) = \delta F_F + \delta F_O + \delta F_{BB} - \delta F_{NEP} + \delta F_{chem}$$

Strategy: Use multiple satellite observations to constrain individual terms and infer the others as a residual.



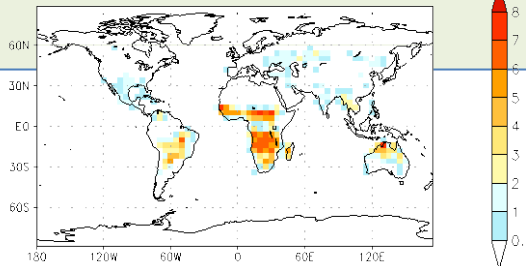
Respiration: combustion

MOPITT Measurements of Pollution in the Atmosphere

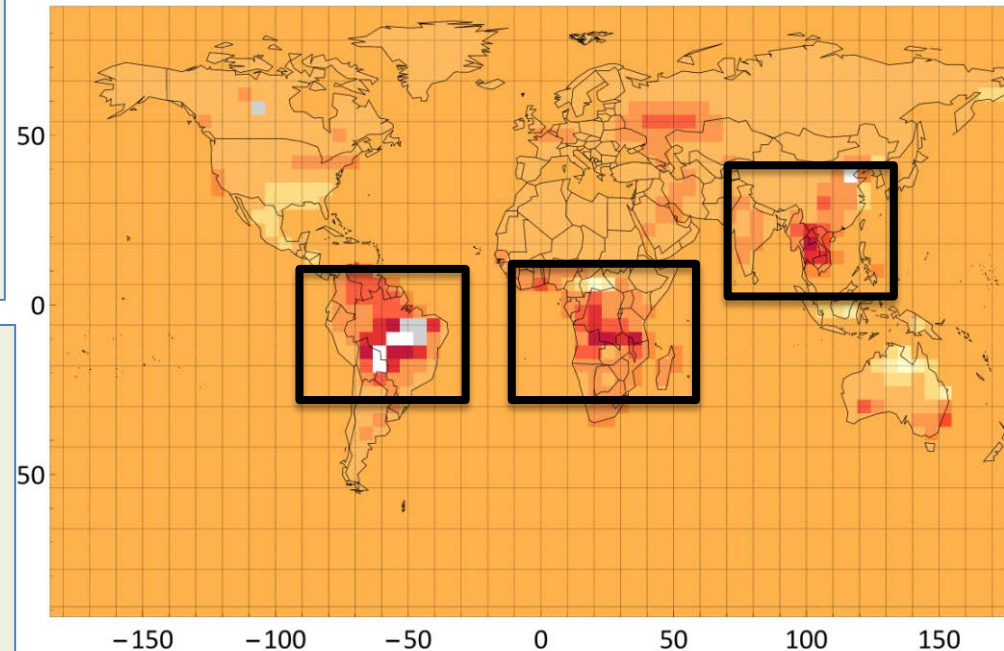


- Carbon monoxide is a by-product of incomplete combustion.
- MOPITT provides CO verticals with near surface sensitivity.
- CMS-Flux estimates CO from MOPITT and converts to CO₂

- CO₂ from biomass burning is calculated from CO/CO₂ ratios (Andreae and Merlet, GBC, 2001)
- Emission factors are a function of dry mass (given) and burning efficiency, which is a function of plant function type.



δBB (2011-2010)



Emission factors Incorporation of Bloom et al uncertainties underway



Flux decomposition: NEP

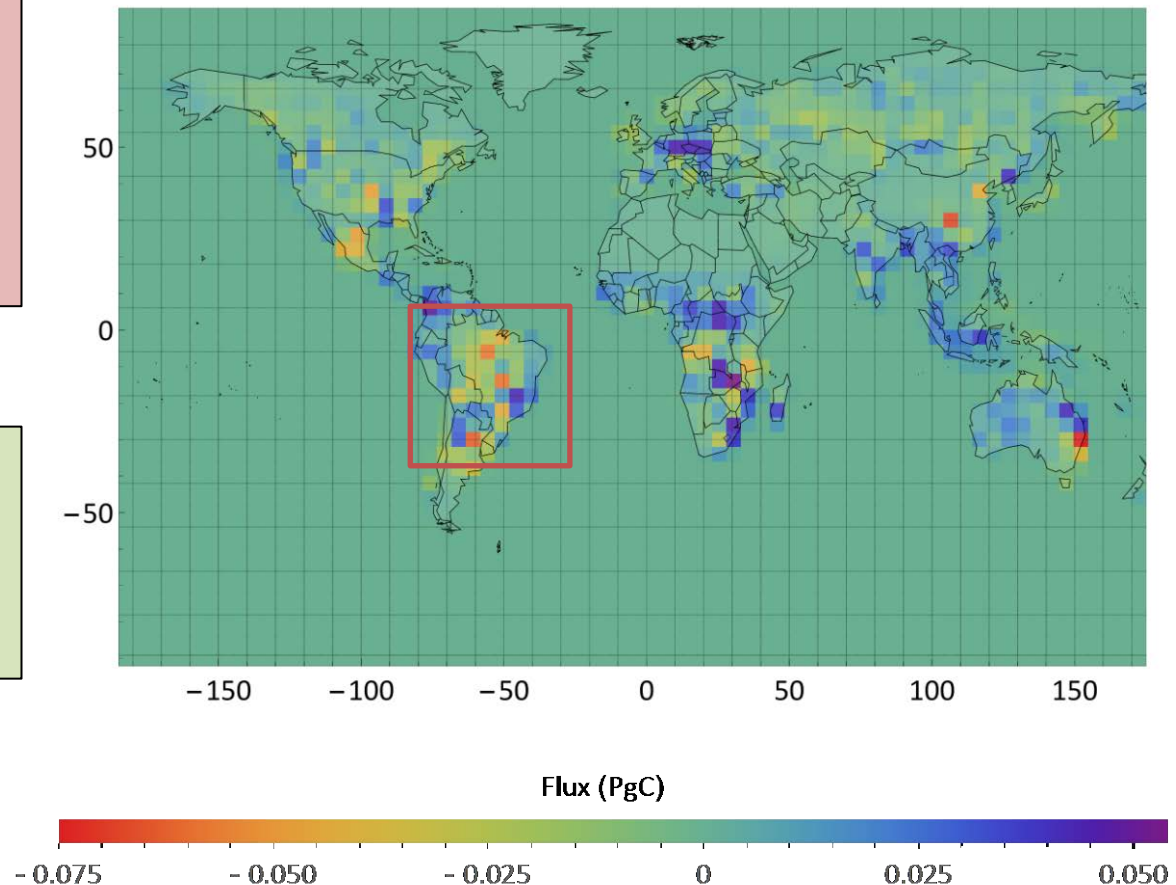
$$\delta F_{2011-2010}^{\uparrow}(x, y, t) = \delta F_F + \delta F_O + \delta F_{BB} - \delta F_{NEP} + \delta F_{chem}$$

Net flux into the atmosphere is positive

Net ecosystem production (NEP) can be inferred as a residual from the posterior total flux, biomass burning, fossil fuel, and chemical production.

The inferred NEP was greater in 2011 in Africa, Southeast Asia, and Australia while lower in Brazil and Mexico.

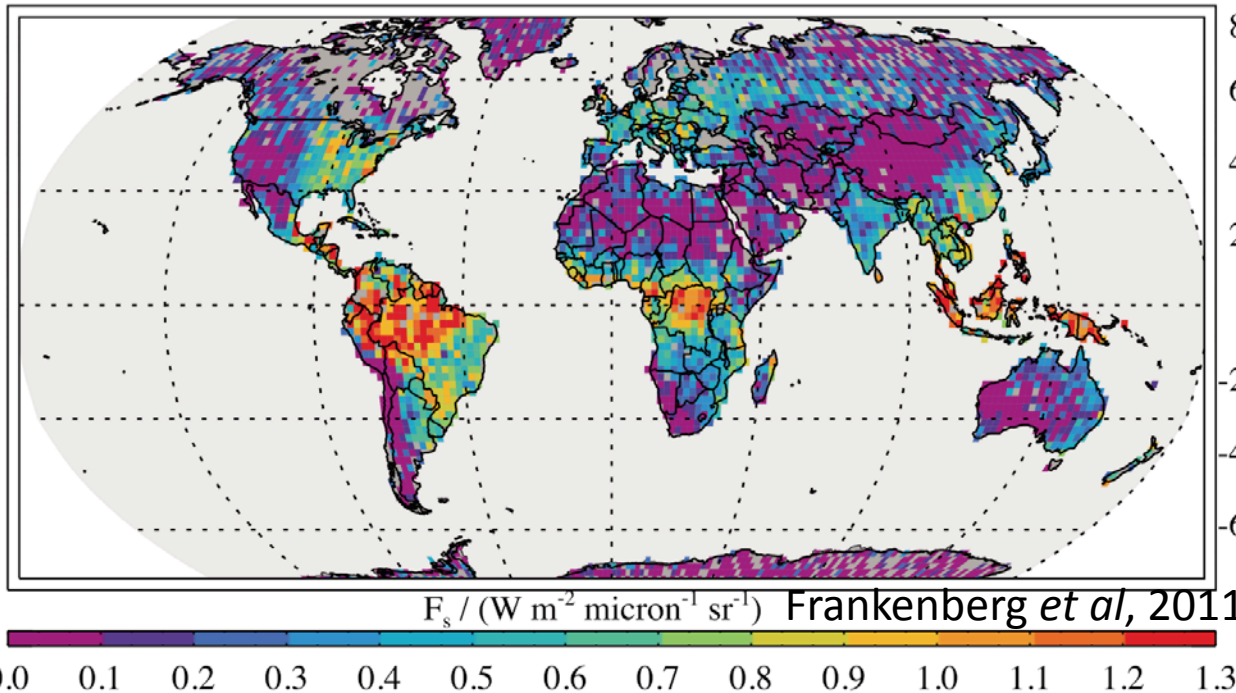
2011-2010 NEP Change



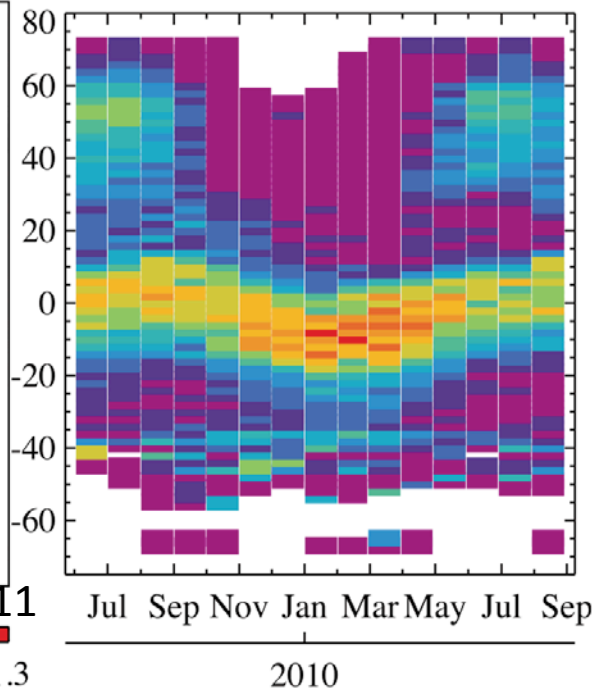


GPP inferred from solar induced fluorescence

A Chlorophyll a fluorescence at 755 nm, June 2009 through May 2010 average



B Timeseries



Optimal estimation provides a framework to determine GPP that accounts for uncertainty in the fluorescence, prior uncertainty in GPP, satellite coverage and timing.

$$\hat{\mathbf{x}} = \min_{\mathbf{x}} C(\mathbf{x}) = \min(\|\mathbf{y} - \mathbf{F}(\mathbf{x})\|_{\mathbf{S}_n}^2 + \|\mathbf{x} - \mathbf{x}_a\|_{\mathbf{S}_a}^2)$$

\mathbf{x}_a : mean Trendy GPP

\mathbf{y} : GOSAT SIF at time $\{t_i\}$

$\mathbf{F}(\mathbf{x})$: Observation operator: GPP to GOSAT overpass

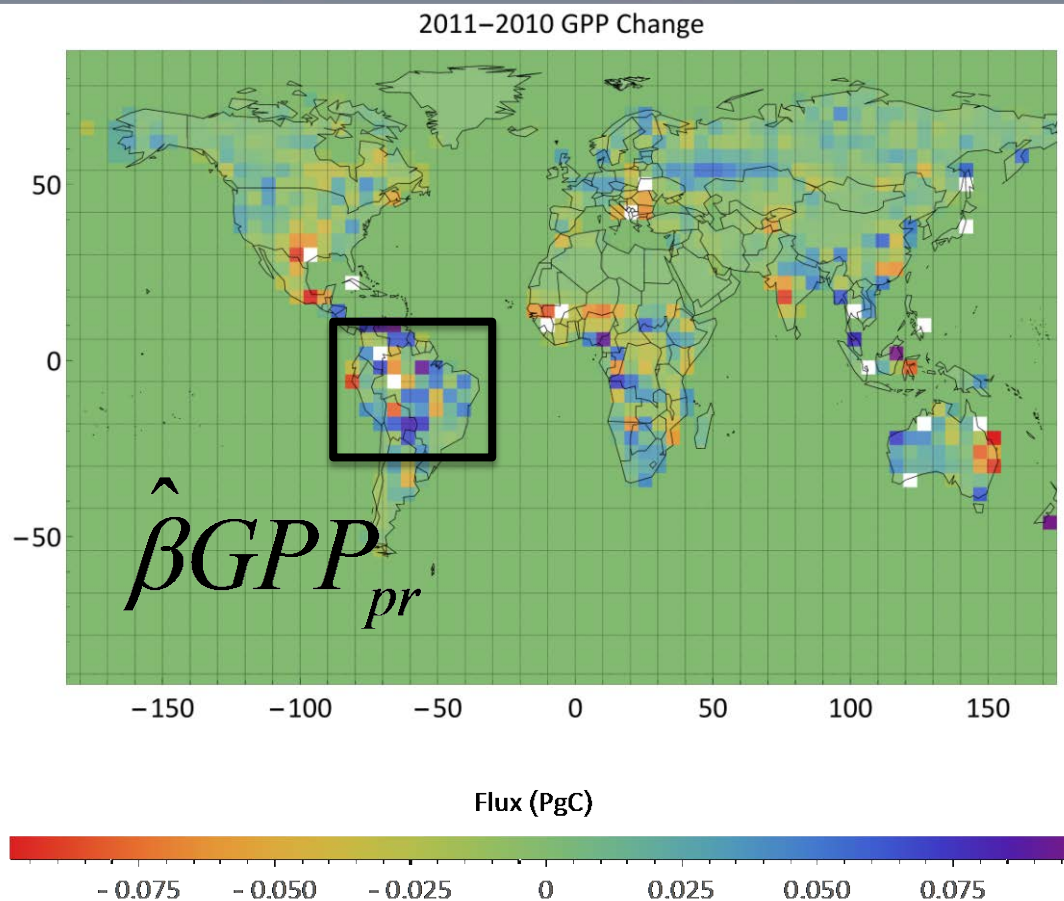
\mathbf{S}_n : Error in GOSAT SIF, \mathbf{S}_a : Ensemble Trendy spread

$$\hat{\mathbf{S}} = (\mathbf{K}^\top \mathbf{S}_n^{-1} \mathbf{K} + \mathbf{S}_a^{-1})^{-1}$$

Parazoo et al, 2013



GPP-Solar induced fluorescence



- Solar induced fluorescence (SIF) is a direct measure of photosynthetic activity (e.g., Porcar-Castell *et al*, 2014)
- Parazoo *et al*, GCB, 2014 introduced an optimal estimation method for GPP balancing satellite SIF and carbon cycle models (TRENDY and Casa-GFED)
- Brazil shows increased productivity in 2011

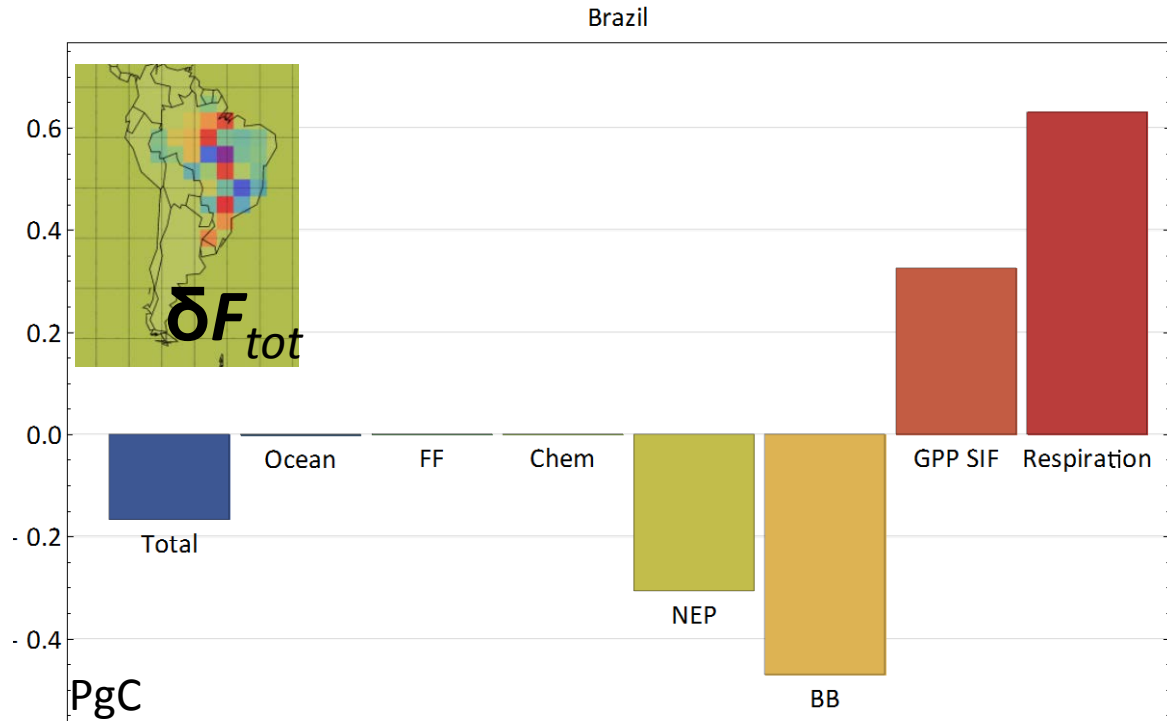


Brazil

$$F^{\uparrow}(x, y, t) = F_F + F_O + F_{BB} + \underbrace{(R - GPP_{SIF})}_{-F_{NEP}} + F_{chem}$$

Net flux into the atmosphere is positive

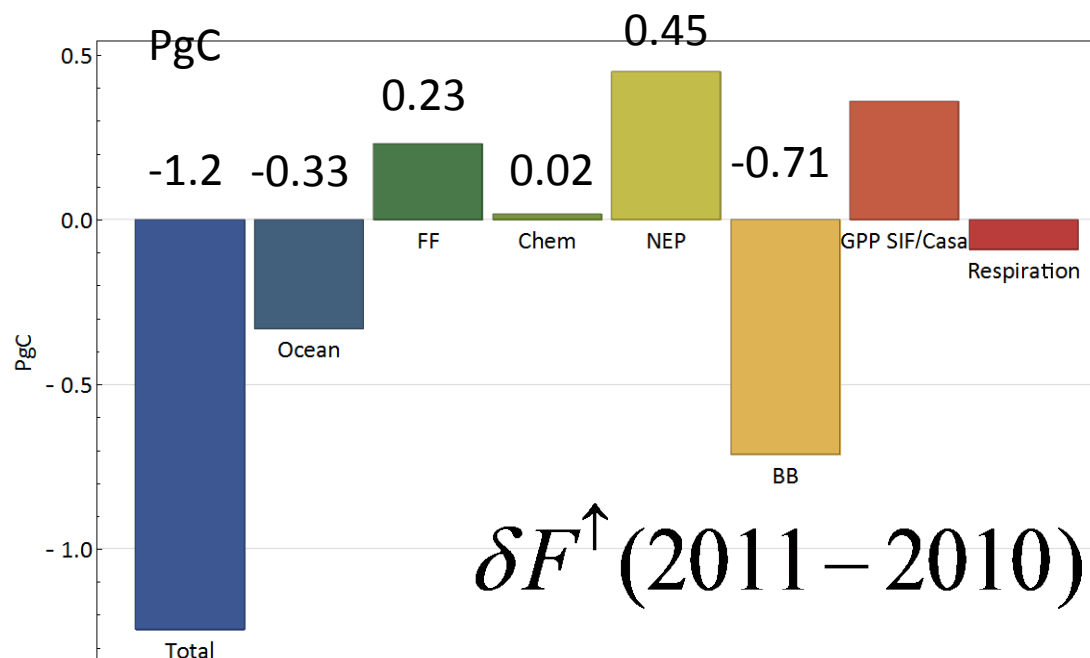
- Weak total flux tendency and strong biomass burning tendency implies negative NEP.
- GPP is positive → R must be higher to compensate.
- GPP → fast carbon pool?
- R → slow carbon pool?
 - Mortality (Saatchi et al, PNAS, 2013)?
 - High precipitation in 2011 could stimulate response





Complex global carbon cycle response to historic Central Pacific El Nino

- Globally, change in CO₂ growth rate (GR) driven by biomass burning (~50%)
 - Brazil was primary contributor
- Global NEP (30% CO₂ GR) was driven by GPP
 - Implies fast carbon pools readjusted.
- Brazilian NEP was driven by respiration
 - Slow carbon pools possibly impacted
- Increased Ocean uptake accounts for 20% CO₂ GR
 - Climatologies, e.g., Takahashi atlas, would miss process



More frequent extreme events like the 2010 mega-drought on Brazil could favor respiration processes, which could accelerate the CO₂ growth rate.

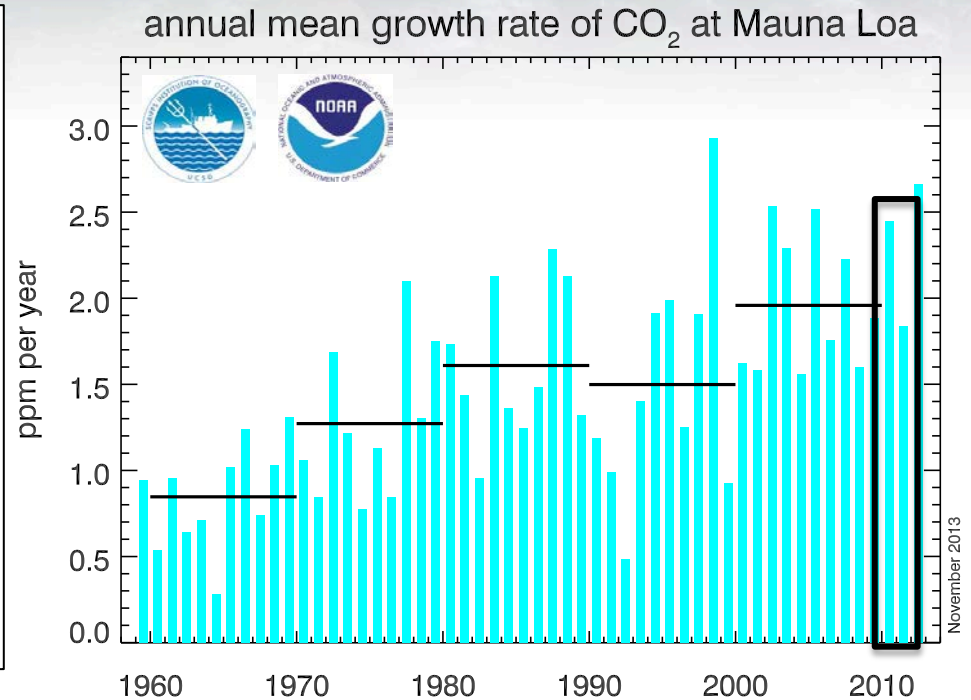
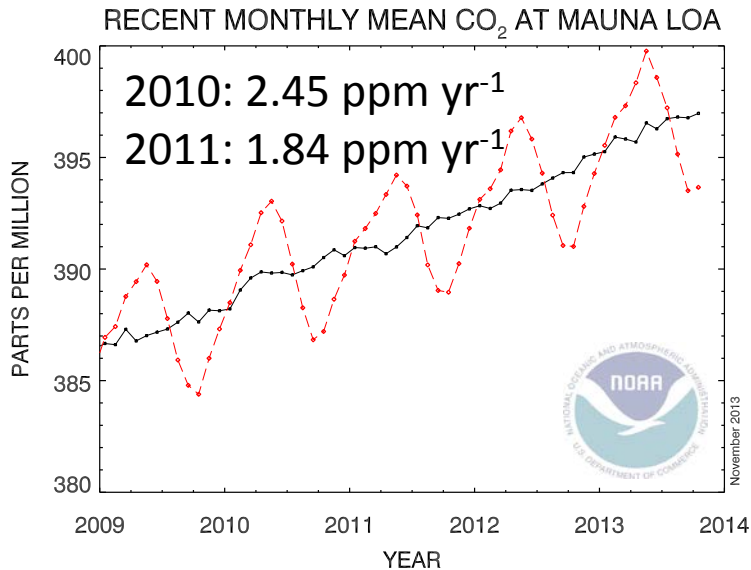
Long-term observations of CO₂, CO, and SIF from OCO-2, CrIS, TROPOMI will help disentangle slow and fast carbon pools and their response to climate. See Dejian Fu *et al* poster



Backup



Atmospheric CO₂ Growth Rate



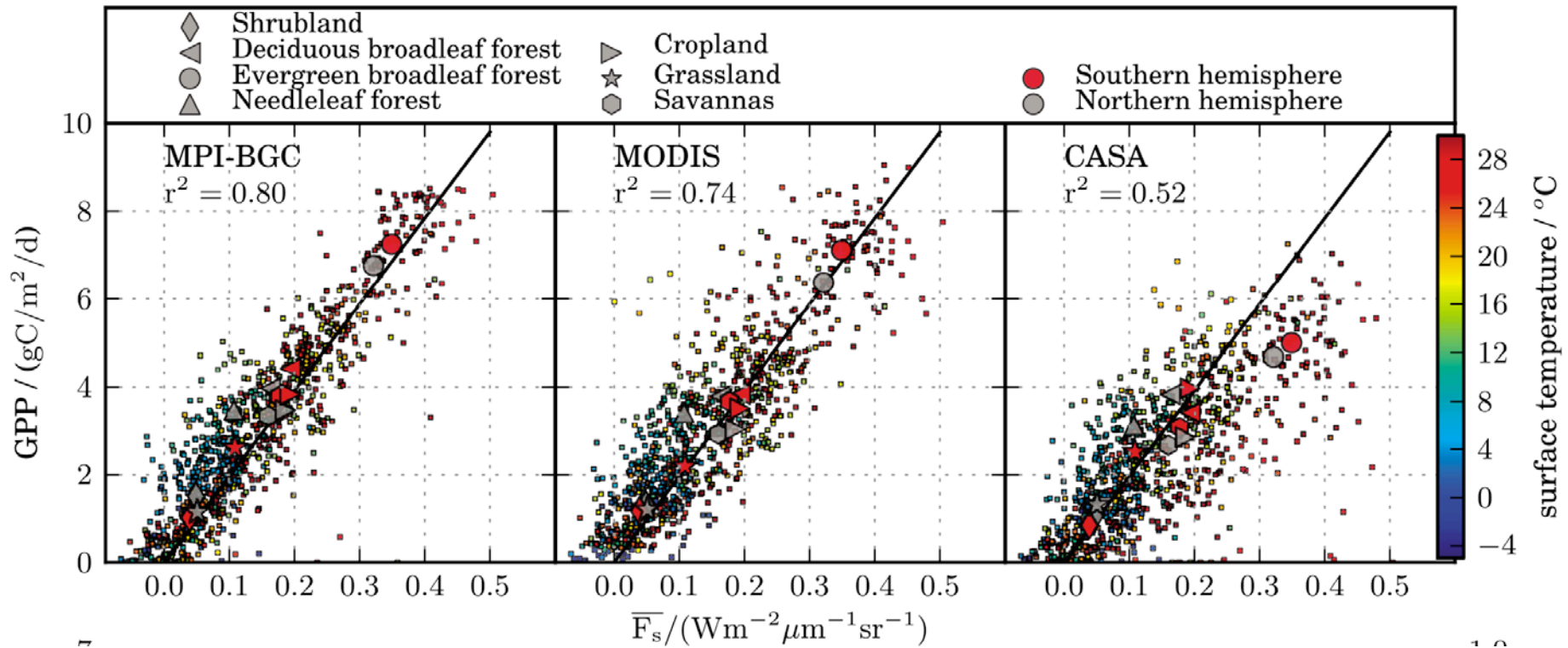
The atmospheric growth rate of CO₂ in 2010 was about 33% higher than 2011.

Why?

What are the processes driving those differences?
What are the spatial drivers of the CO₂ growth rate?



Covariation with GPP



Frankenberg *et al*, 2011

Comparisons with MPI-BGC point to the proportionality of GPP to fluorescence.



Flux decomposition

$$F^{\uparrow}(x, y, t) = F_F + F_O + F_{BB} - F_{NEP} + F_{chem}$$

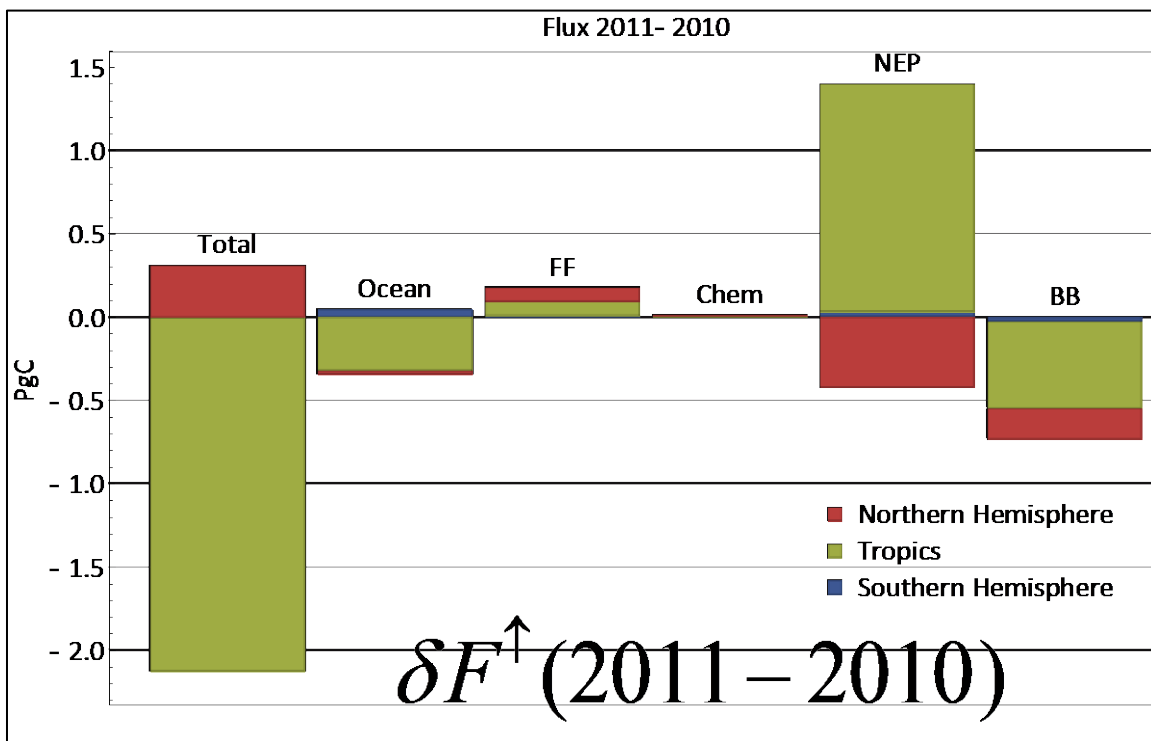
Net flux into the atmosphere is positive

Atmospheric growth rate was driven by changes in tropical flux.

Tropical net ecosystem production (NEP) change was offset by reductions in the NH.

Biomass burning dropped in both hemispheres.

The combination of reduced biomass burning and increased NEP reduced the atmospheric growth rate in 2011 relative to 2010.

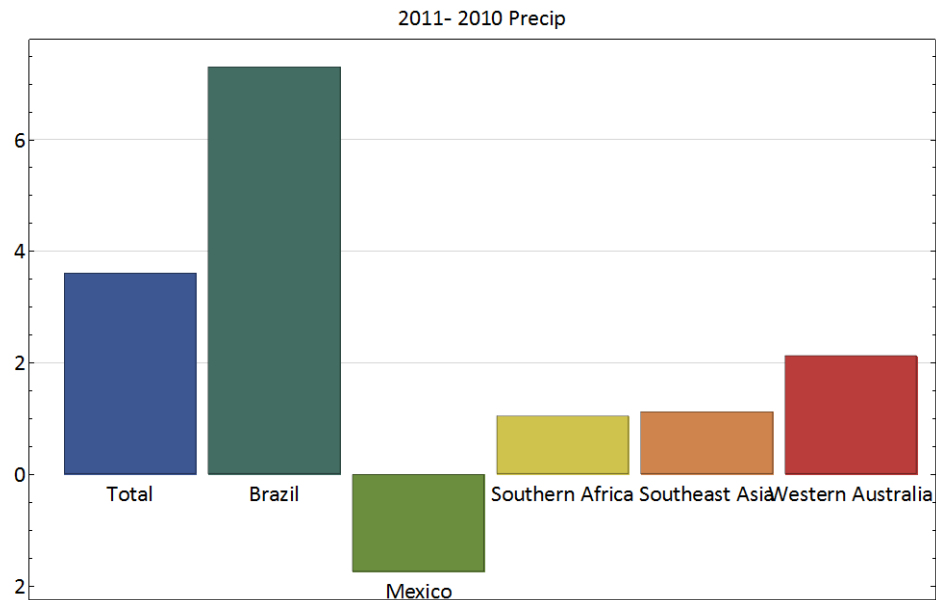
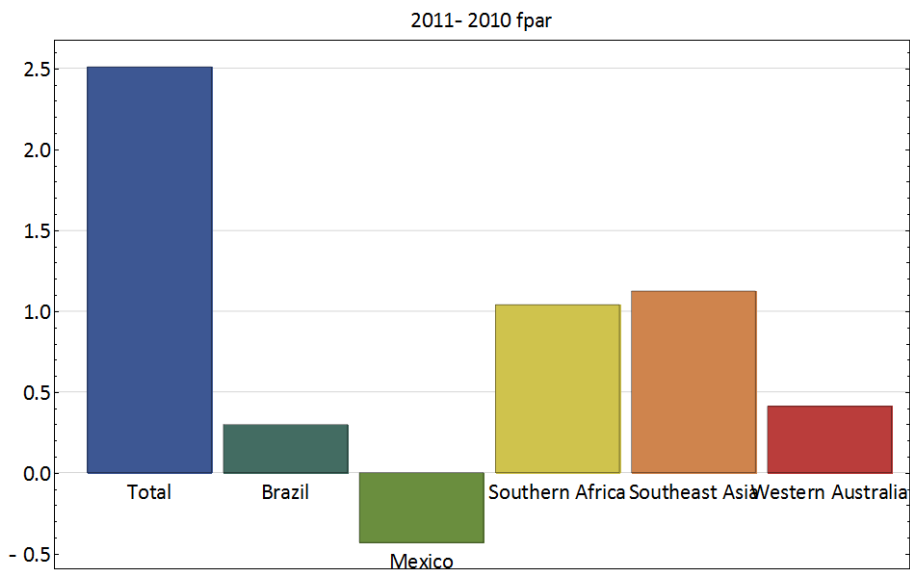
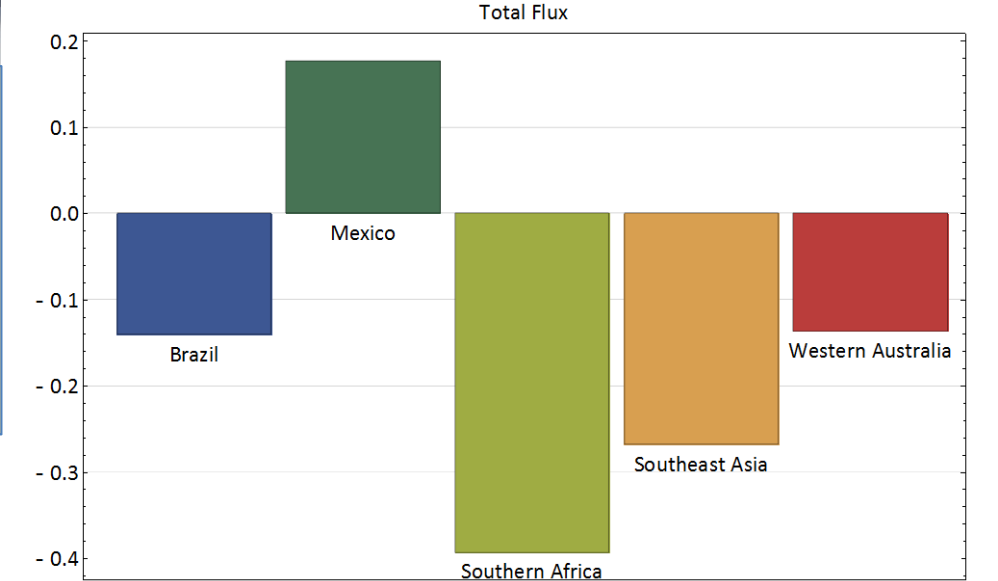


What drove NEP?



Regional drivers

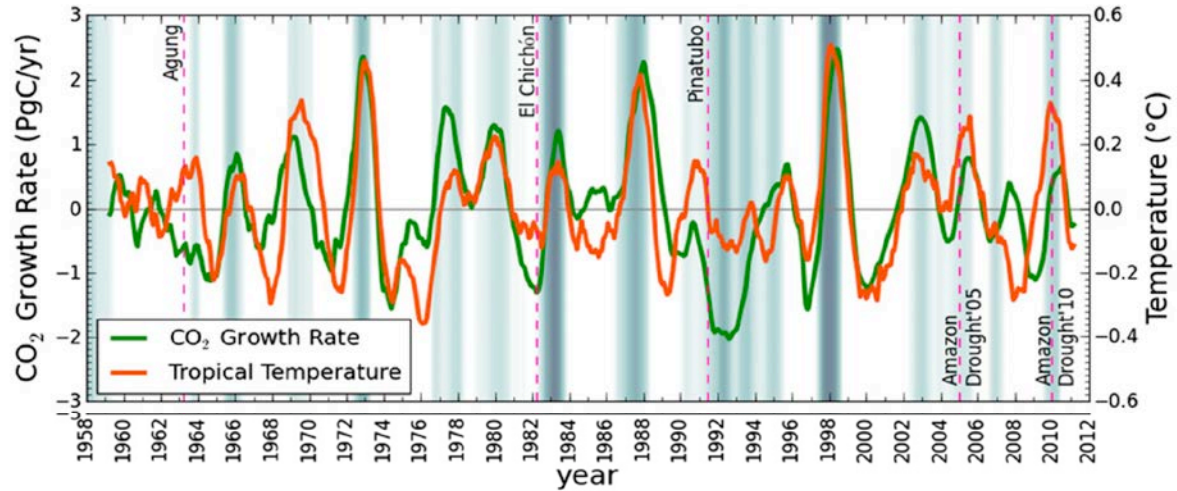
- Southern Africa dominated growth reduction
- Southeast Asia response dominated by NPB (consistent with Basu *et al*, 2013)
- Sign of Fpar and precipitation consistent with total flux response



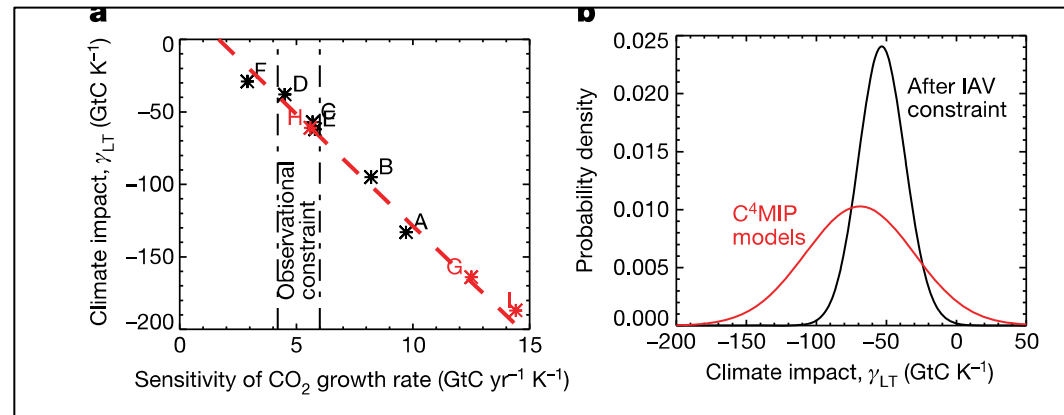


Link between tropical temperatures and CO₂ growth rate

Wang et al, 2013 (PNAS) showed that 1 °C tropical temperature anomaly leads to a $3.5 \pm 0.6 \text{ PgC yr}^{-1}$ CO₂ growth rate anomaly (1959-2011) ($r^2 \approx 0.5$).



Cox et al, 2013 (Nature) used that same relationship to apply an observational constraint on carbon-cycle climate feedback parameter $\rightarrow \gamma \text{ (GtC K}^{-1}\text{)}$ in C⁴MIP models



Spatially explicit drivers of atmospheric CO₂ growth rate and ancillary species could provide better process scale constraints

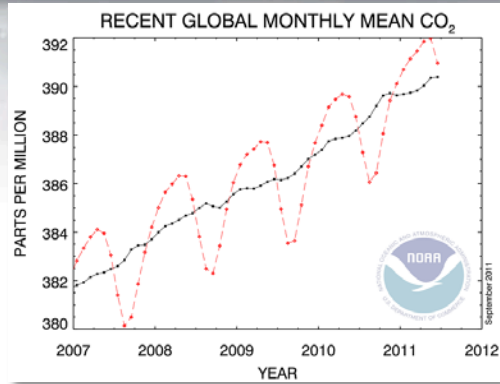


Attribution Strategy

A priori carbon cycle budgets are constructed to be consistent with atmospheric CO₂ growth

9.25 PgC yr⁻¹ (2010)
9.50 PgC yr⁻¹ (2011)

CO₂ accumulation
2010 2.44 ppm yr⁻¹
2011 1.84 ppm yr⁻¹



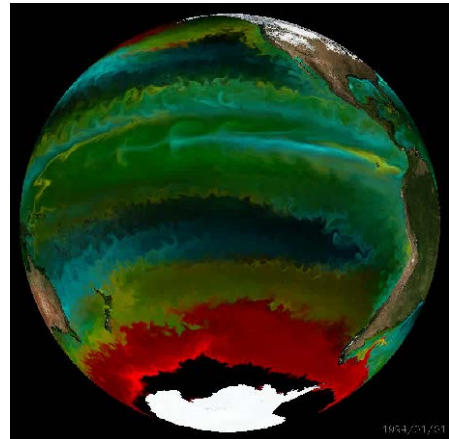
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4.85 - 2.92 PgC yr⁻¹(2010)
5.79 - 2.24 PgC yr⁻¹(2011)



Fossil Fuel (FFDAS)
Incorporates separate chemical CO₂ production, aviation, bunker, and shipping (Nassar et al, 2010)

2.38 (2010) PgC yr⁻¹
2.71 (2011) PgC yr⁻¹
Derived from ECCO-2 ocean flux estimate



Land-use change and river export are not included



GPP-Solar induced fluorescence

$$C_{\beta,j} = \frac{1}{2} [\mathbf{y}_j - \mathbf{f}_j(\beta_j)]^T \mathbf{R}_j^{-1} [\mathbf{y}_j - \mathbf{f}_j(\beta_j)] + \frac{1}{2} [\beta_j - \beta_b]^T \mathbf{P}_j^{-1} [\beta_j - \beta_b]$$

- Solar induced fluorescence is a direct measure of photosynthetic activity (e.g., Porcar-Castell *et al*, 2014)
- Parazoo *et al*, GCB, 2014 developed an inverse method to estimate GPP from satellite SIF and TRENDY models

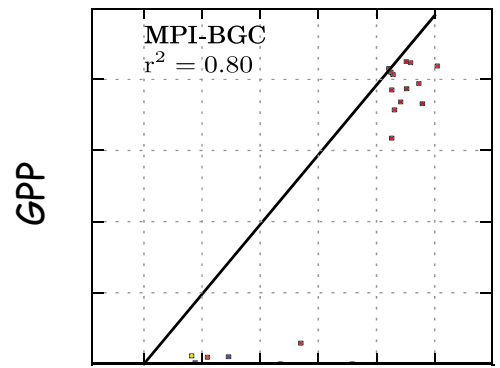
$$\mathbf{f}(\boldsymbol{\beta}) = \text{GPP}(x,y,t) = \beta_{\text{GPP}}(x,y,t) * \text{GPP}_{\text{pr}}(x,y,t)$$

y: Inferred GPP (Fluorescence scaled to MPI)

R: GPP Uncertainty (Measurement Error)

P: Prior Uncertainty (TRENDY)

β_b Prior Scale Factor (CASA or JULES)



Frankenberg *et al* 2011

Fluorescence

2011-2010 GPP Change

