

Future sea level projections

Tamsin Edwards
Open University



Ice Flows climate change game puts the fate of Antarctic penguins in your hands

WIRED

RONNE

FILCHNER



SANAE

FUJI

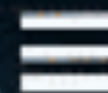
CHILLY

VOSTOK

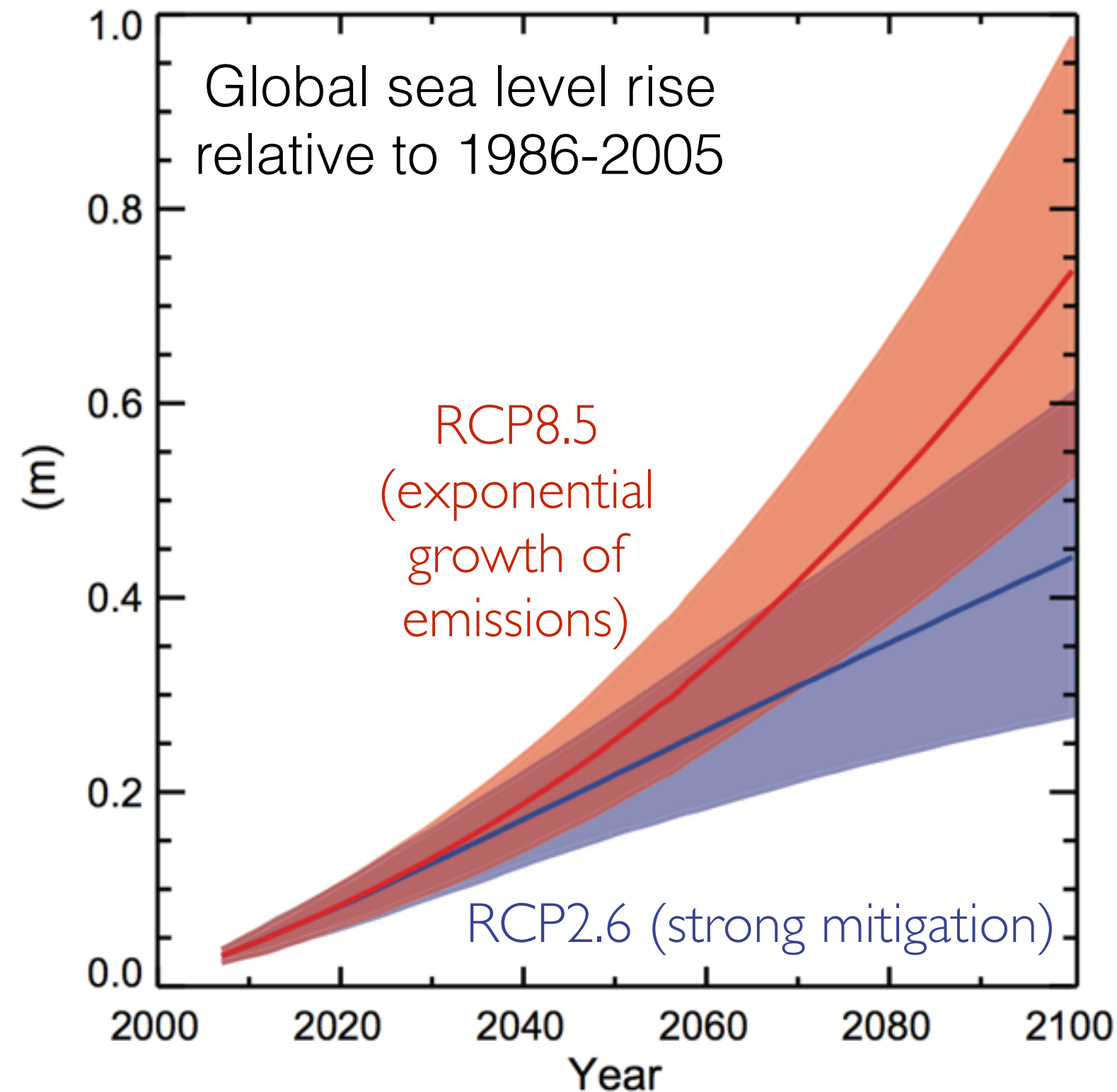


SCORE: 02

🕒 49



IPCC projections



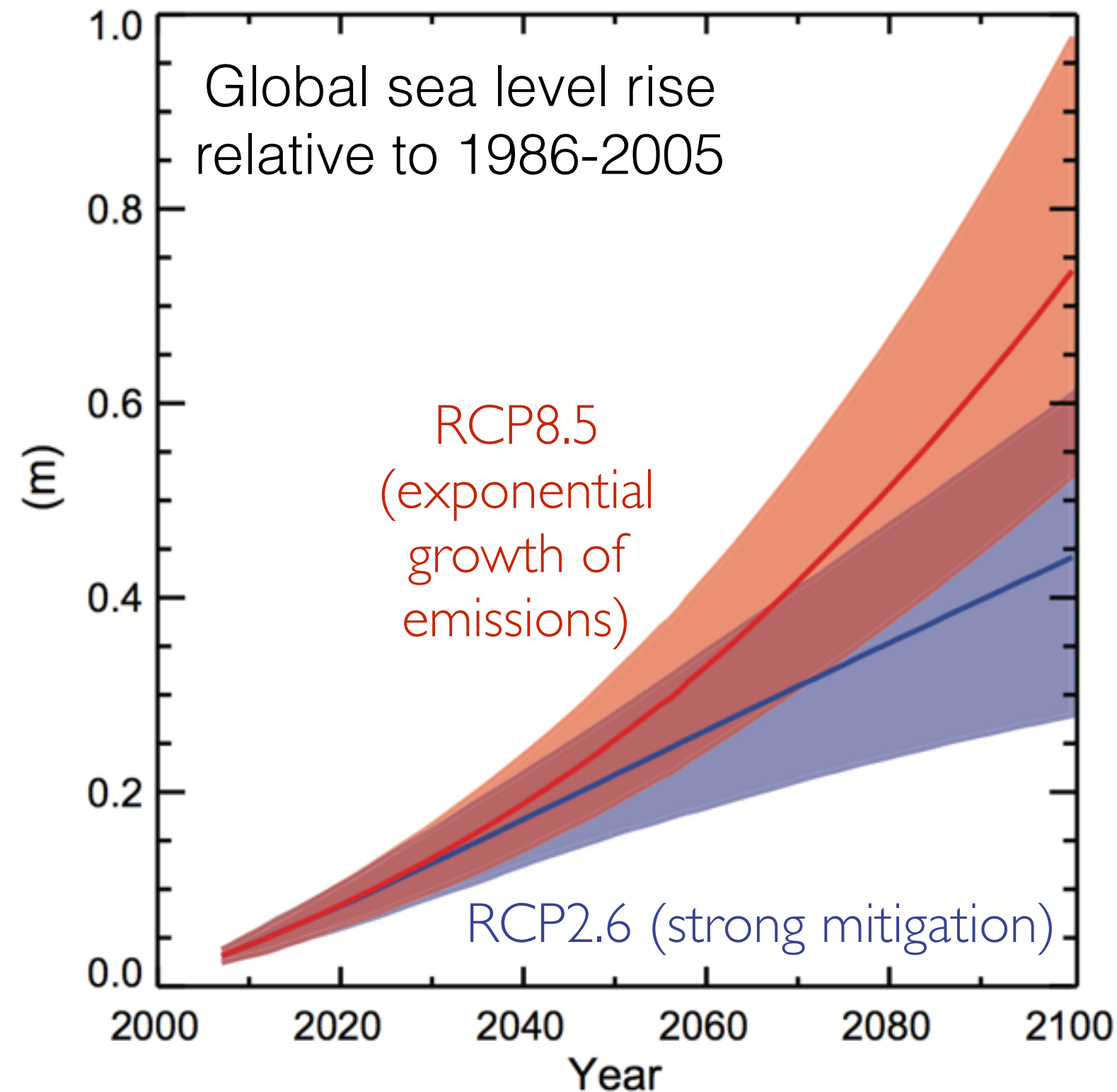
Median [$\geq 66\%$]

74 [52, 98] cm

44 [28, 61] cm

- Substantial sea level rise no matter what
- Large uncertainties

IPCC projections



Median [$\geq 66\%$]

74 [52, 98] cm

44 [28, 61] cm

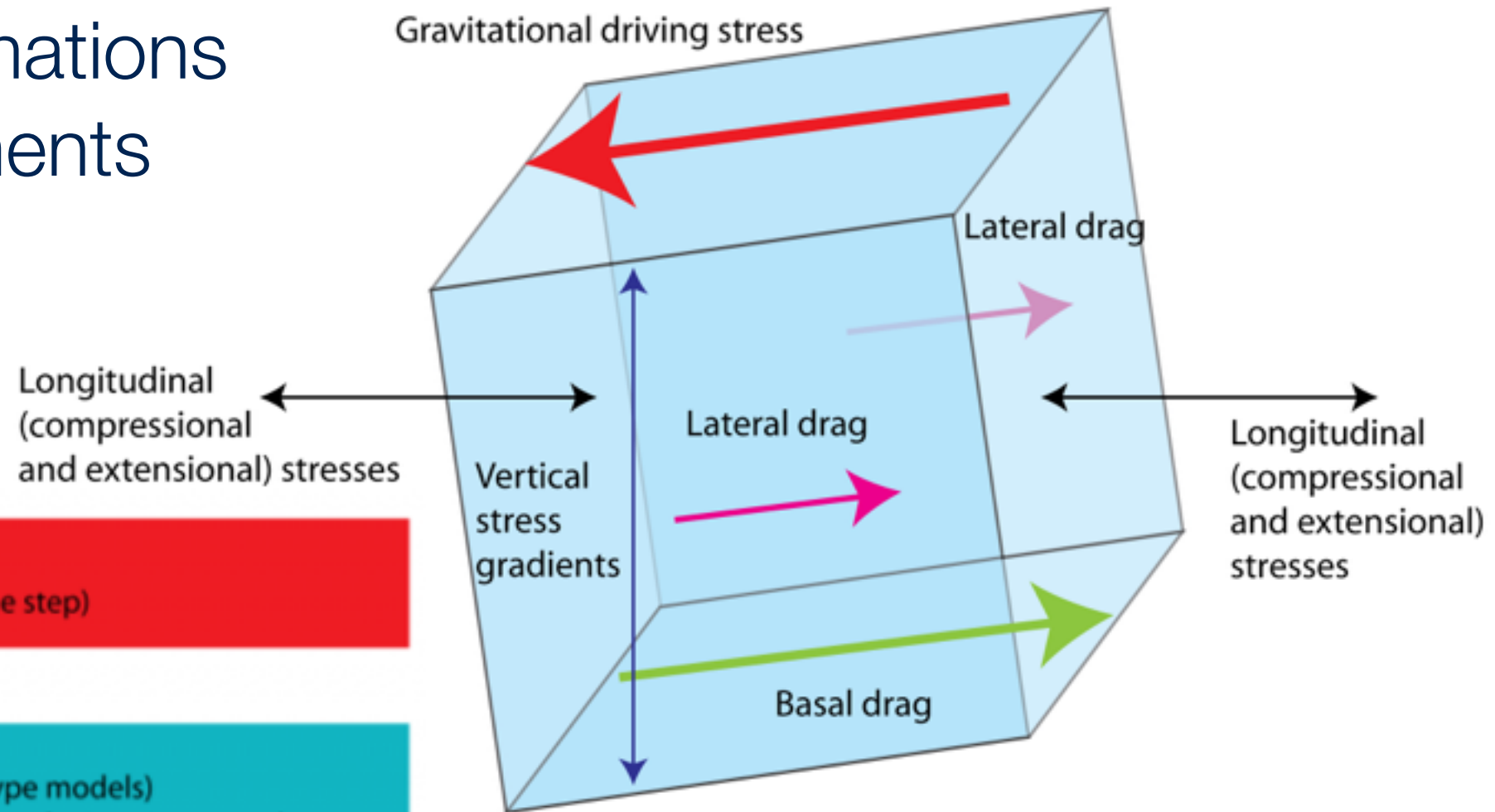
- Substantial sea level rise no matter what
- Large uncertainties
- Ice sheets 1/4 or more
- Antarctica the largest uncertainty: 7 [-1, 16] cm
- Very poorly-constrained upper tail: 50 to 100 cm

Ice sheet models

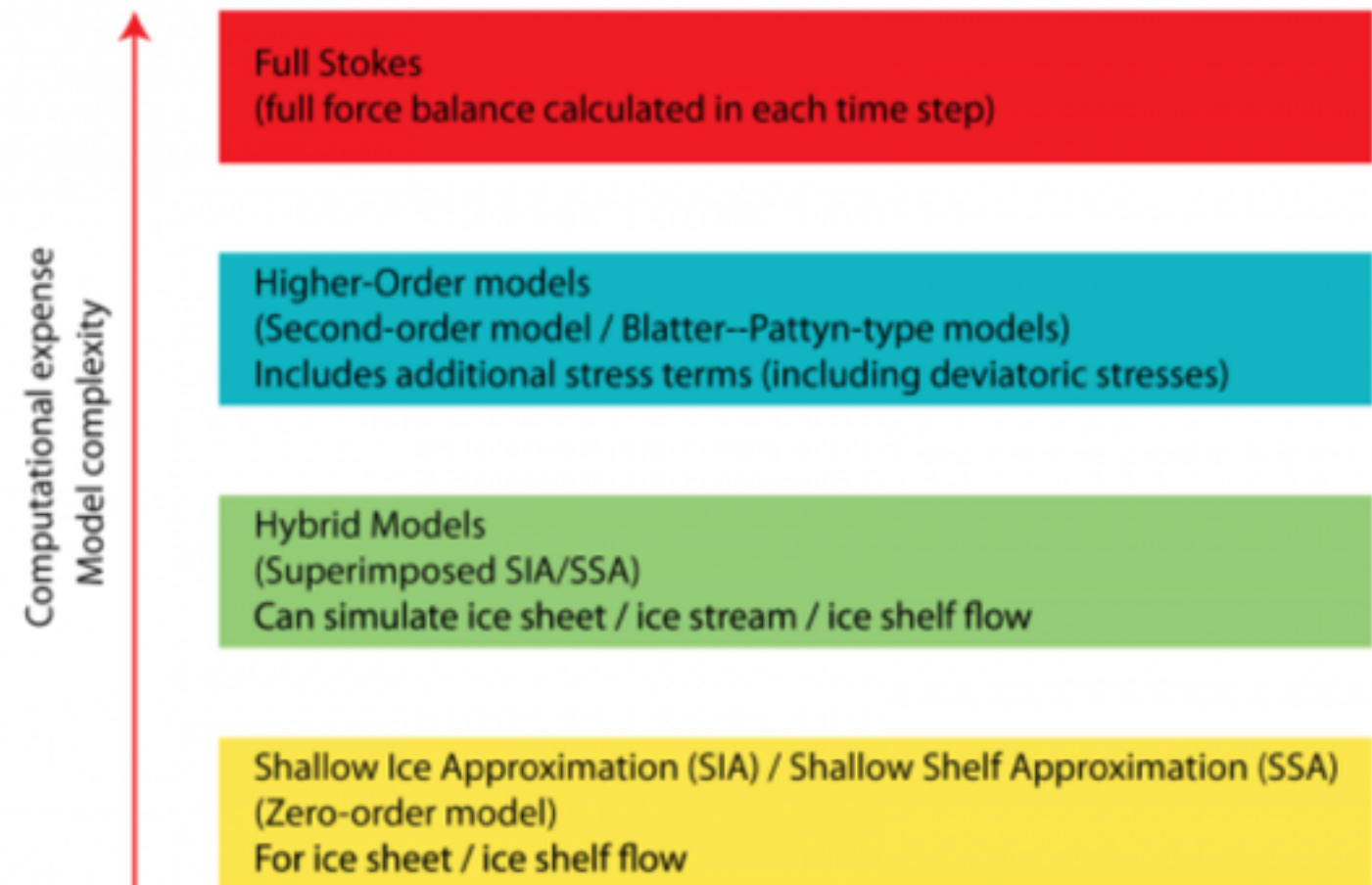
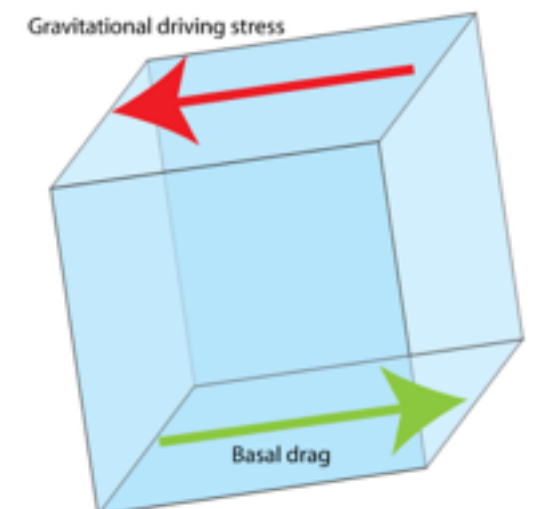
- flow of ice under its own weight
- different approximations of stress components

AntarcticGlaciers.org

Full Stokes

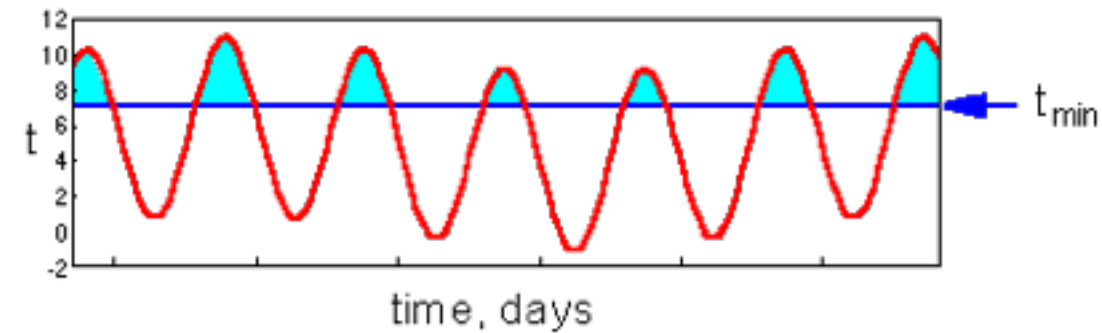


Shallow Ice Approximation

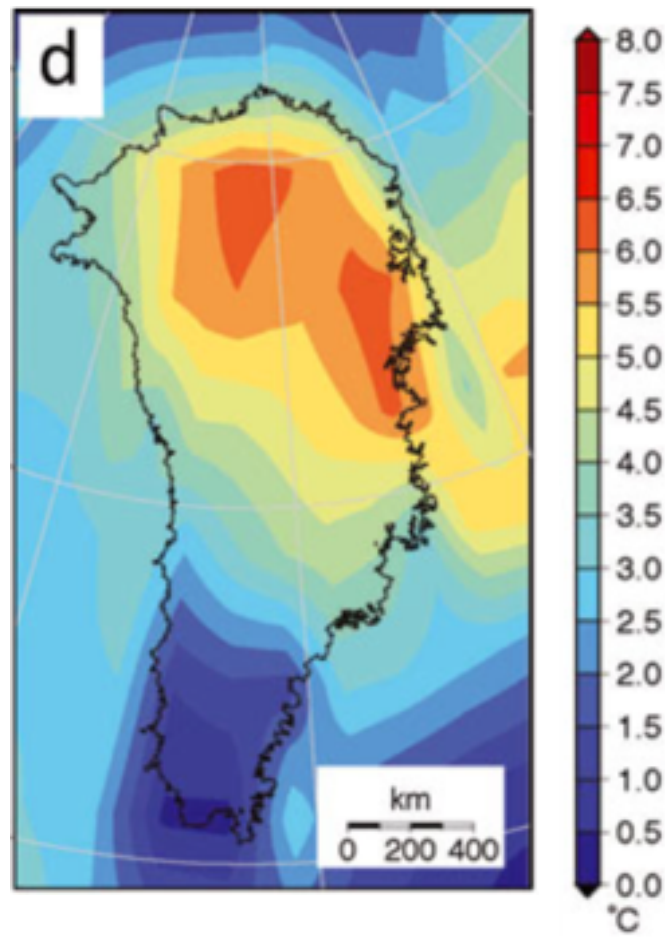


Ice sheet models

- can calculate surface mass balance from atmosphere
 - degree day models
 - energy balance models
- and/or basal mass balance from ocean
 - melt parameterisation

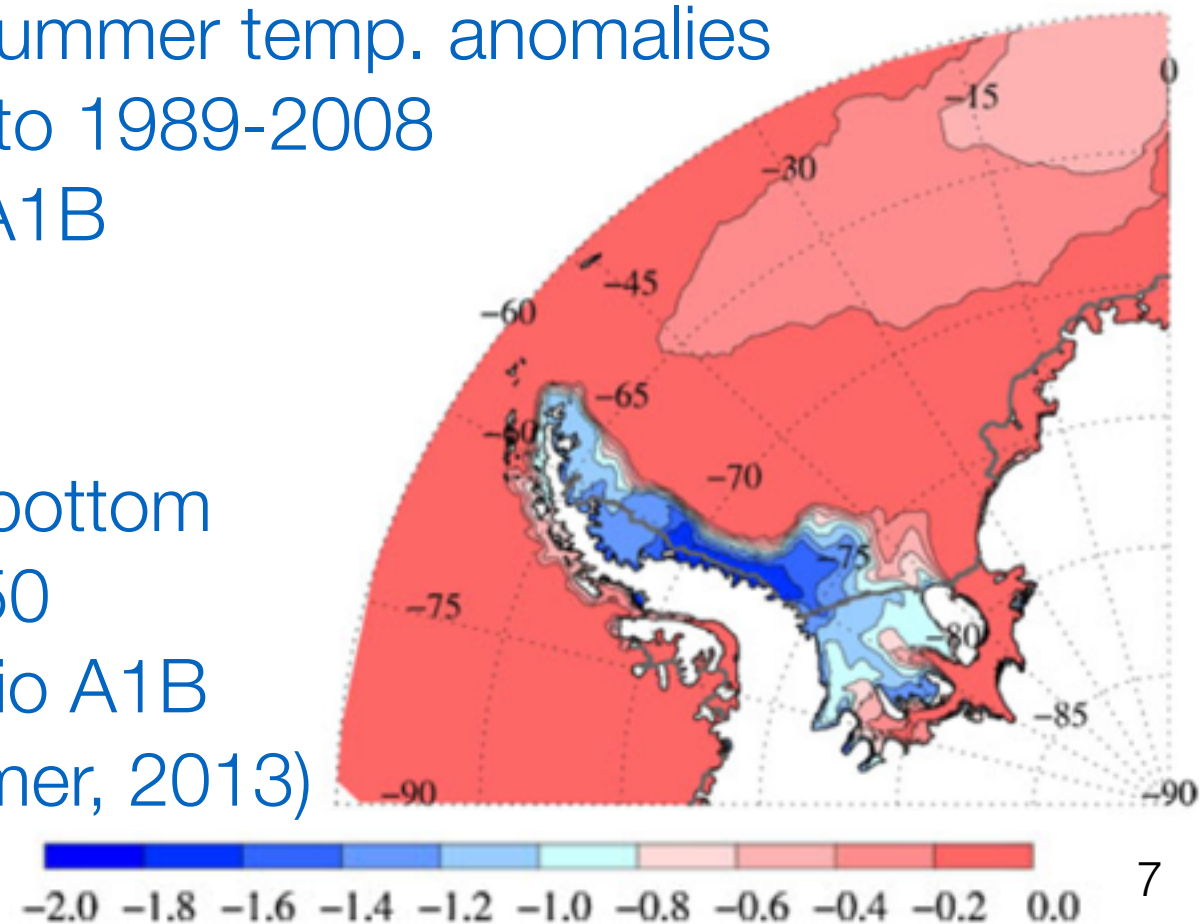


Source: Alexei Sharov

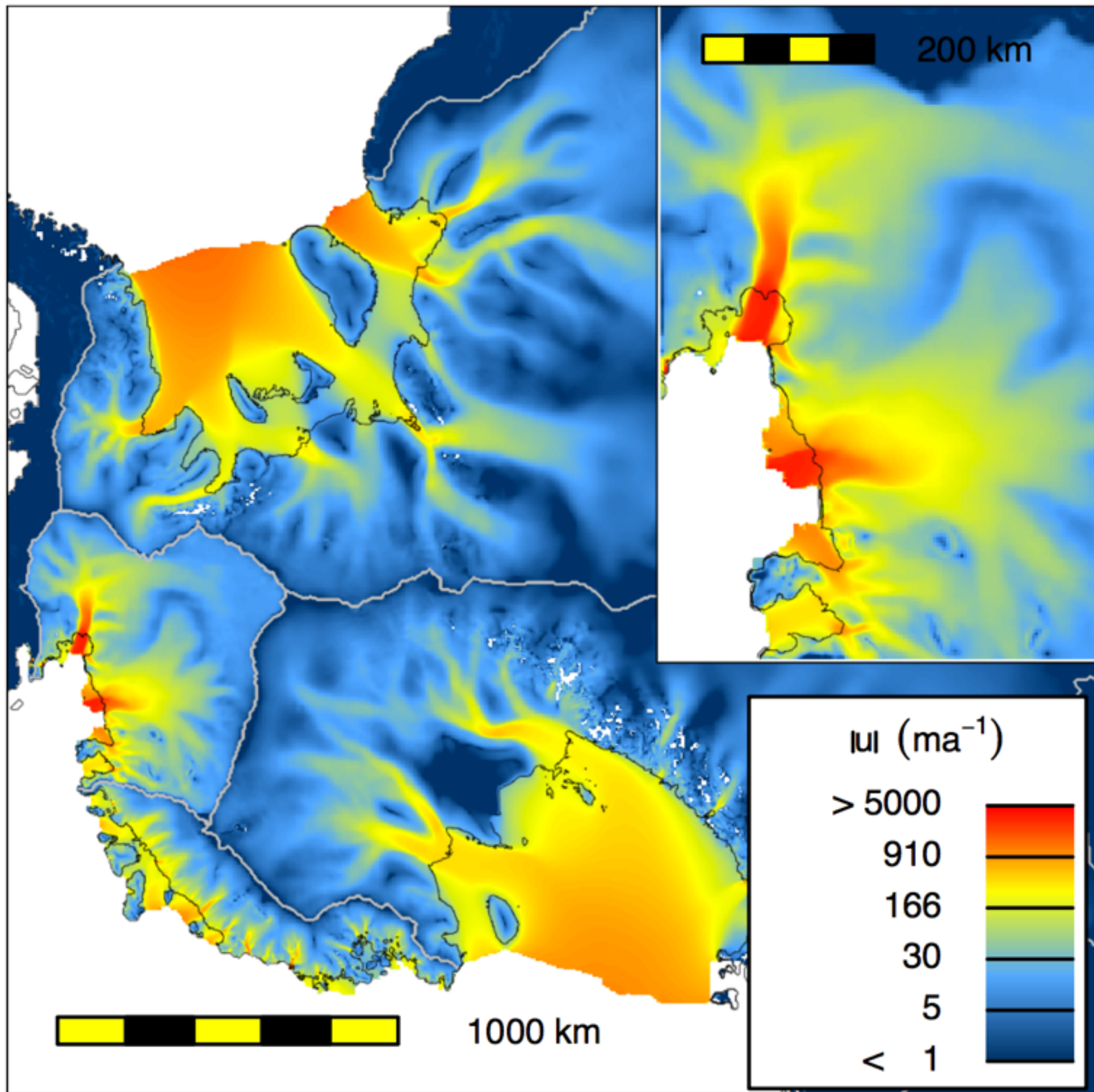


Global climate model summer temp. anomalies for 2091-2100 relative to 1989-2008 under SRES scenario A1B (Goelzer et al., 2013)

Ocean model near-bottom temperatures in 2150 under SRES scenario A1B (Timmerman & Hellmer, 2013)

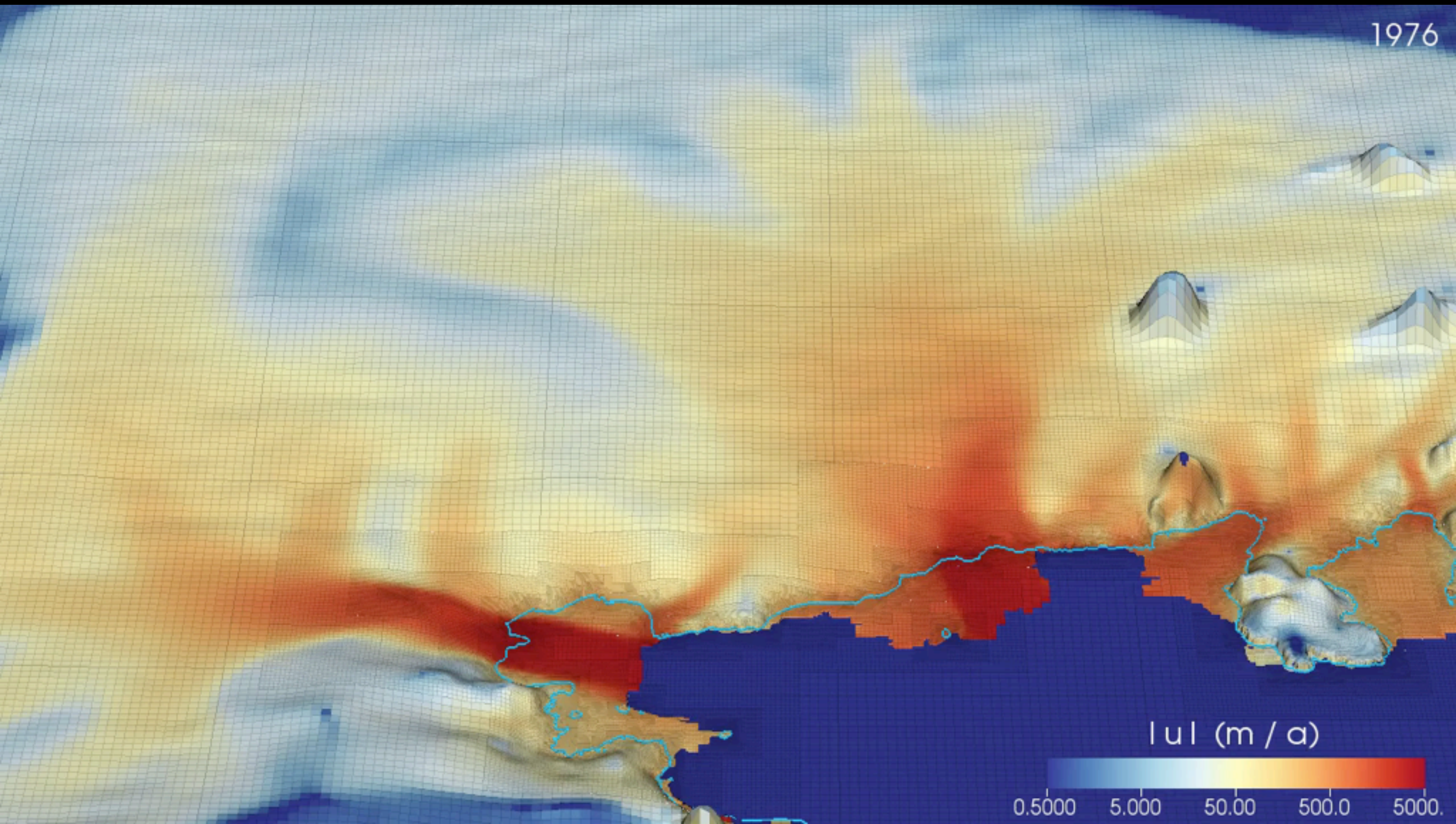


BISICLES



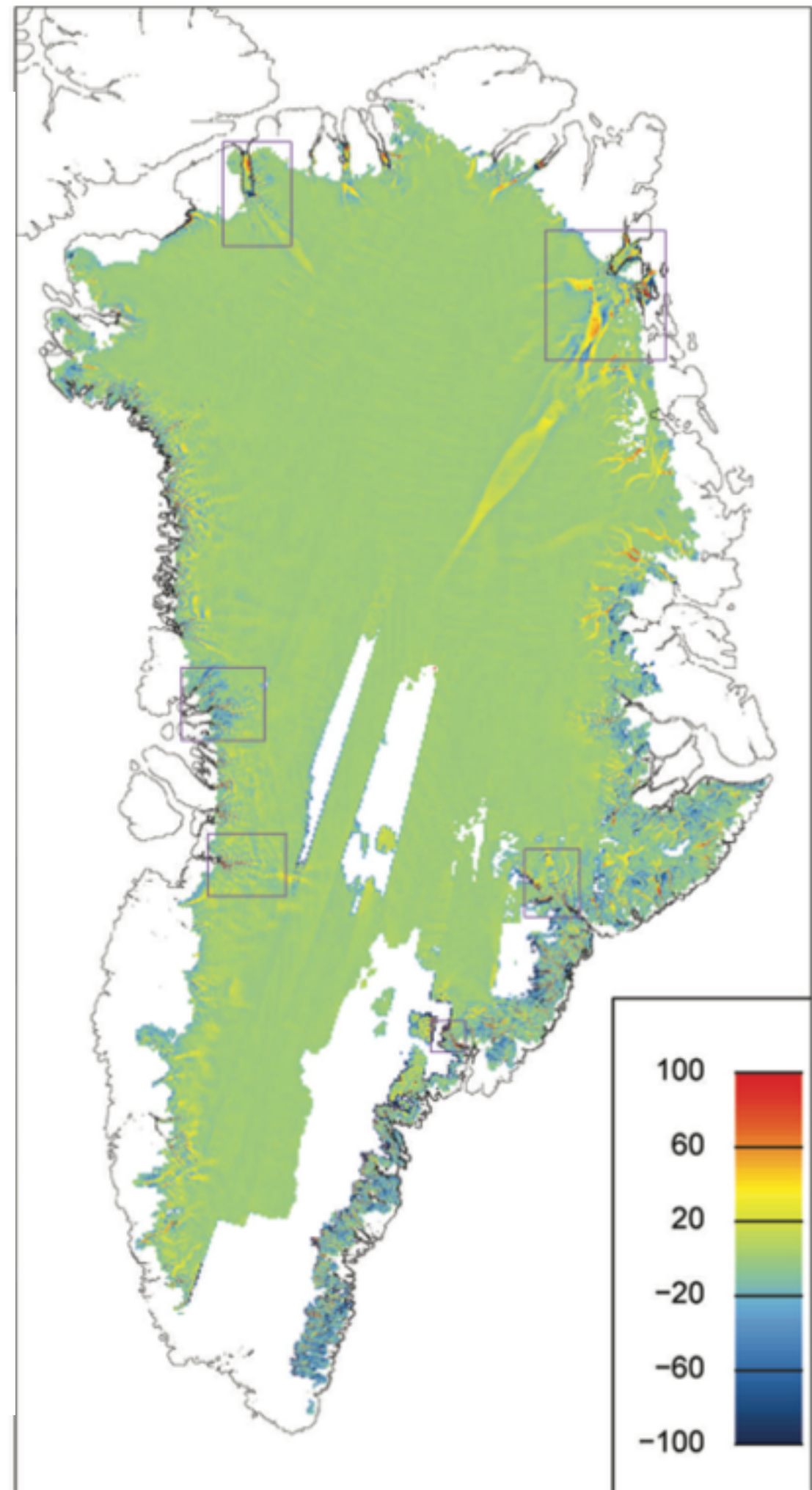
initial ice flow speed
(Cornford et al.,
2015)

BISICLES

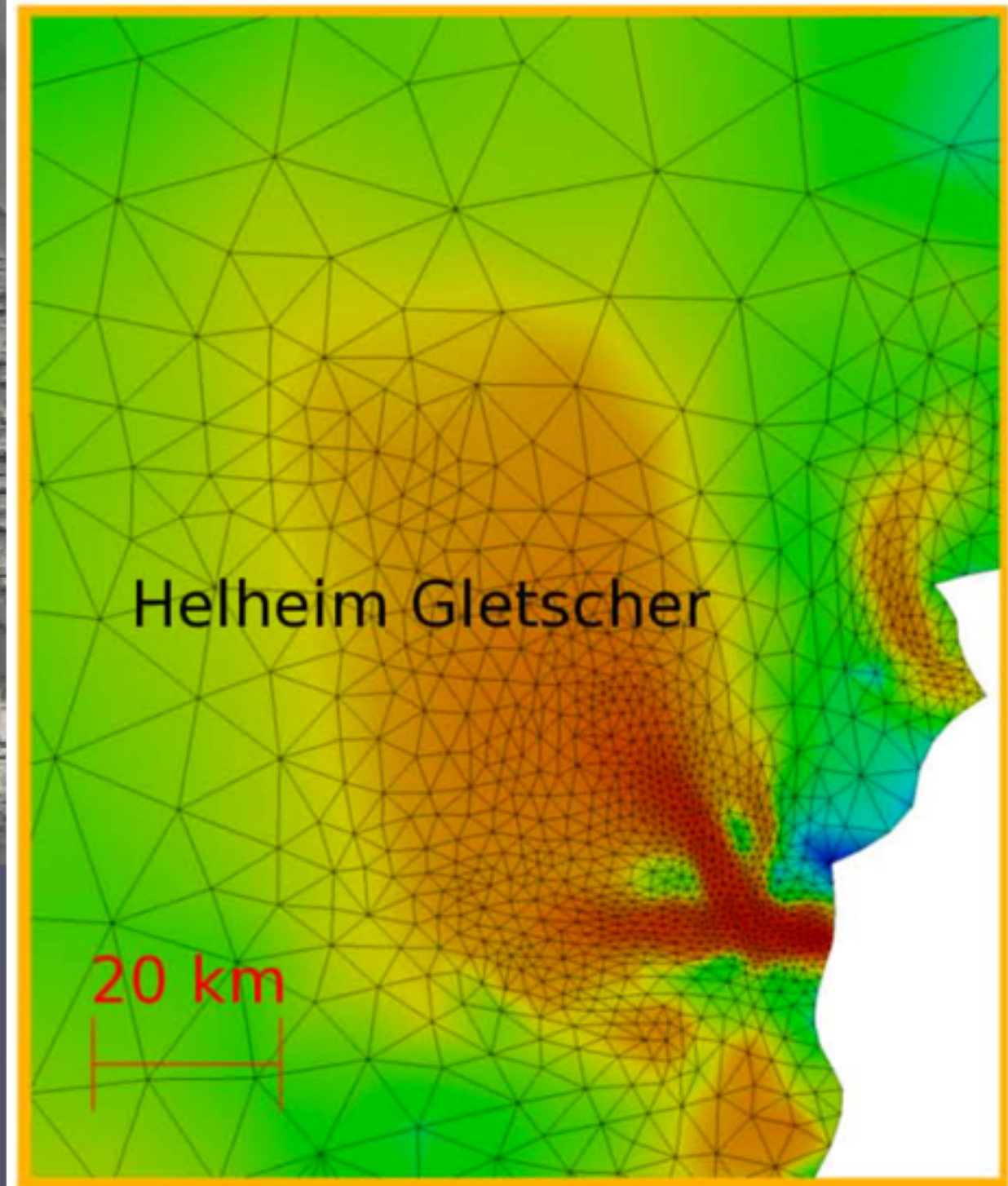
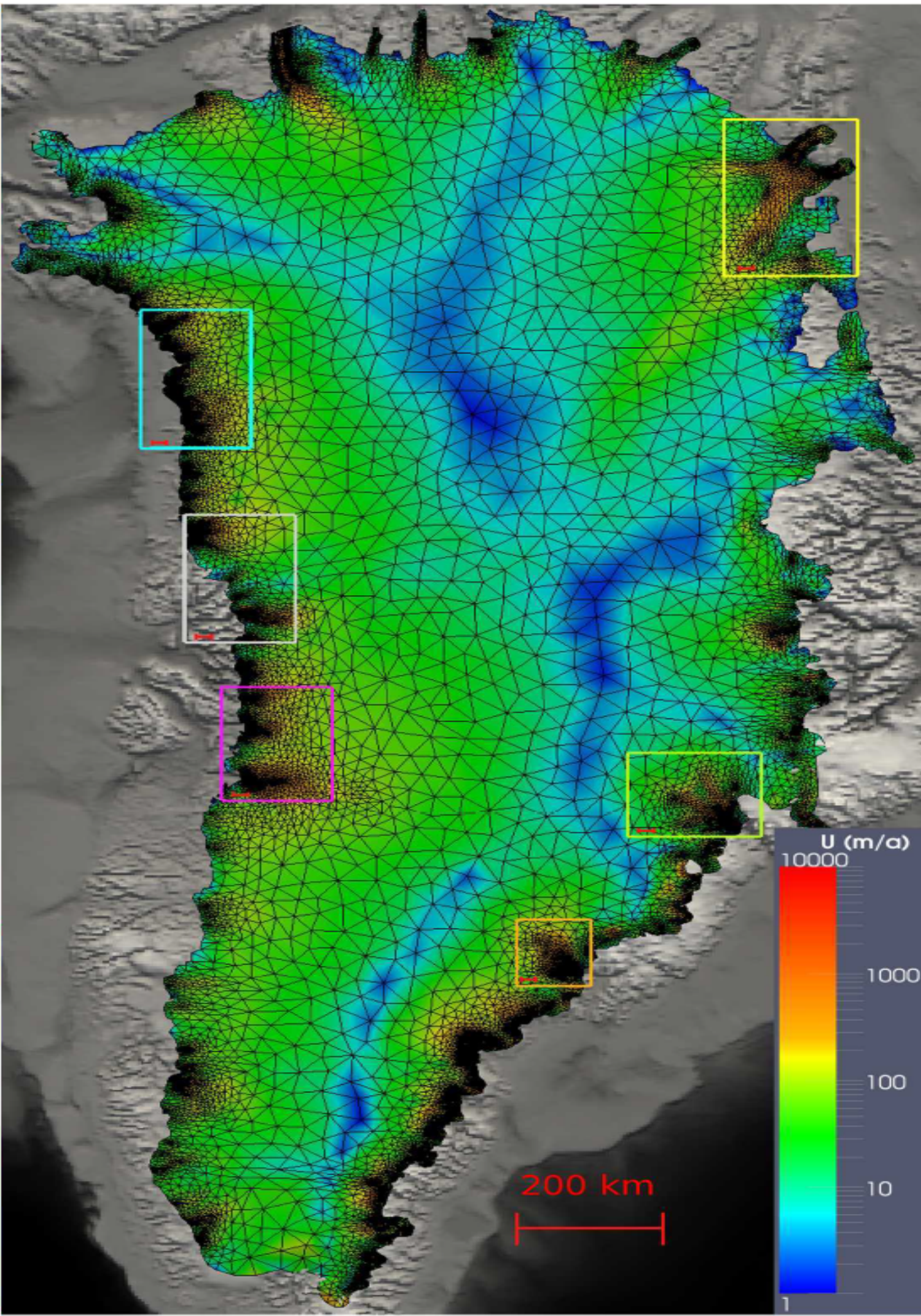


ice flow speed forced by ocean under SRES scenario A1B
(Cornford et al., 2015)

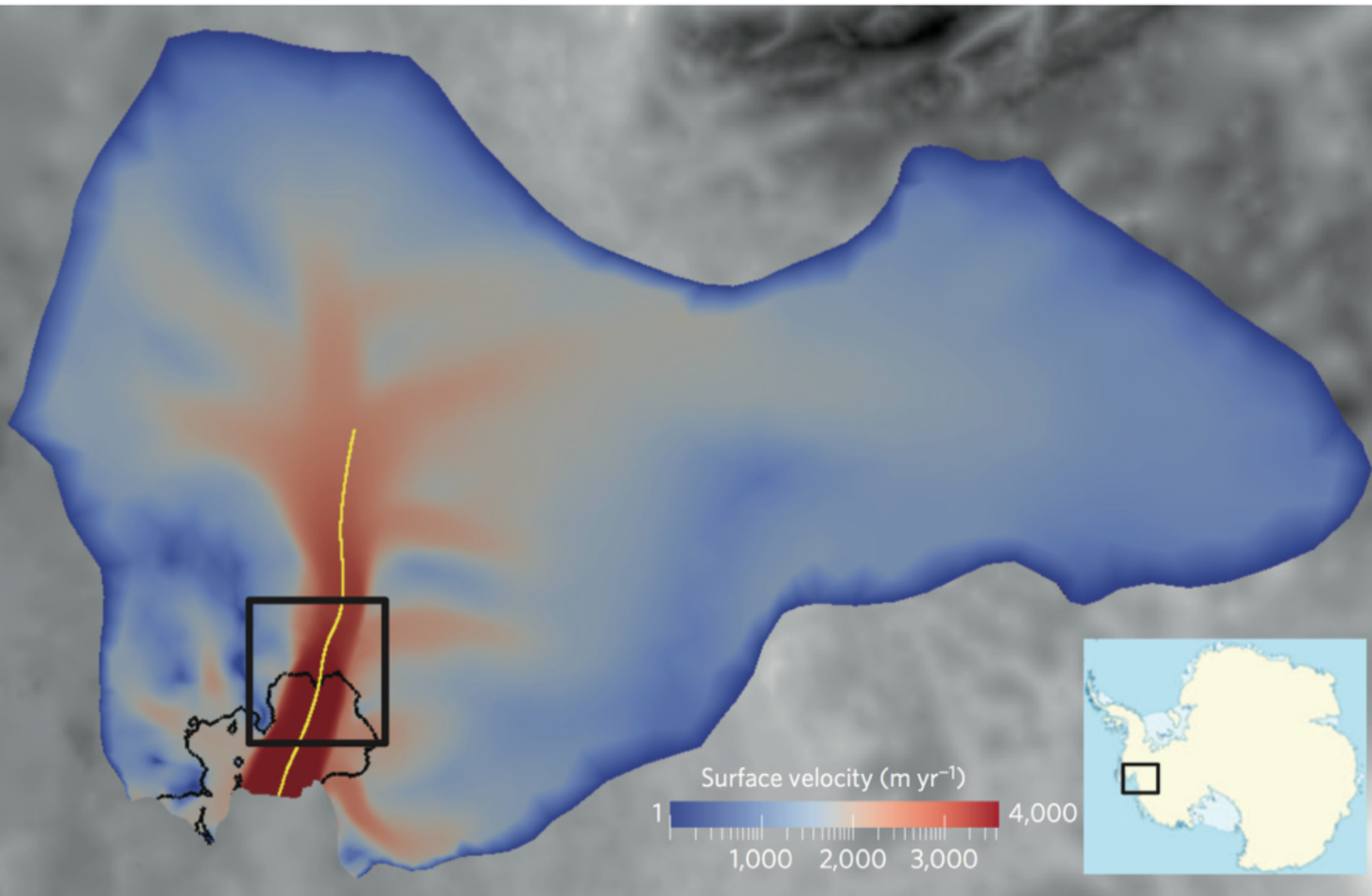
BISICLES



% mismatch between initial ice speed and observations (Lee et al., 2015)

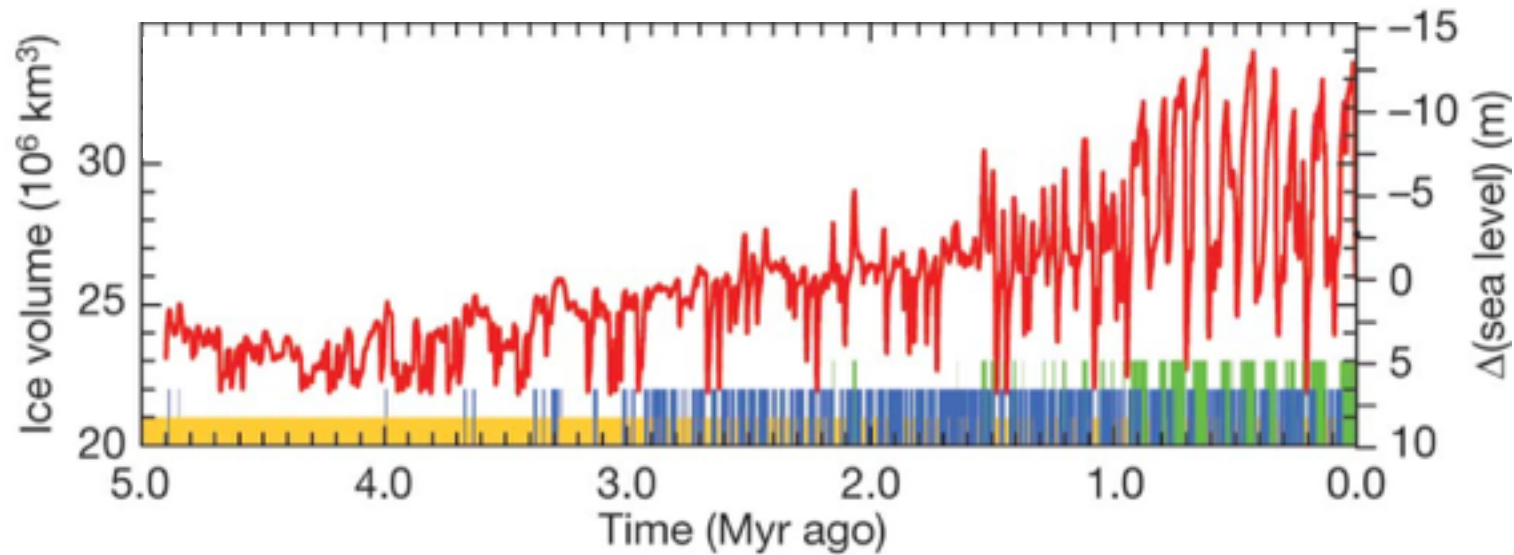


initial surface velocities
(Gillet-Chaulet et al., 2012)



What is an ice sheet model used for?

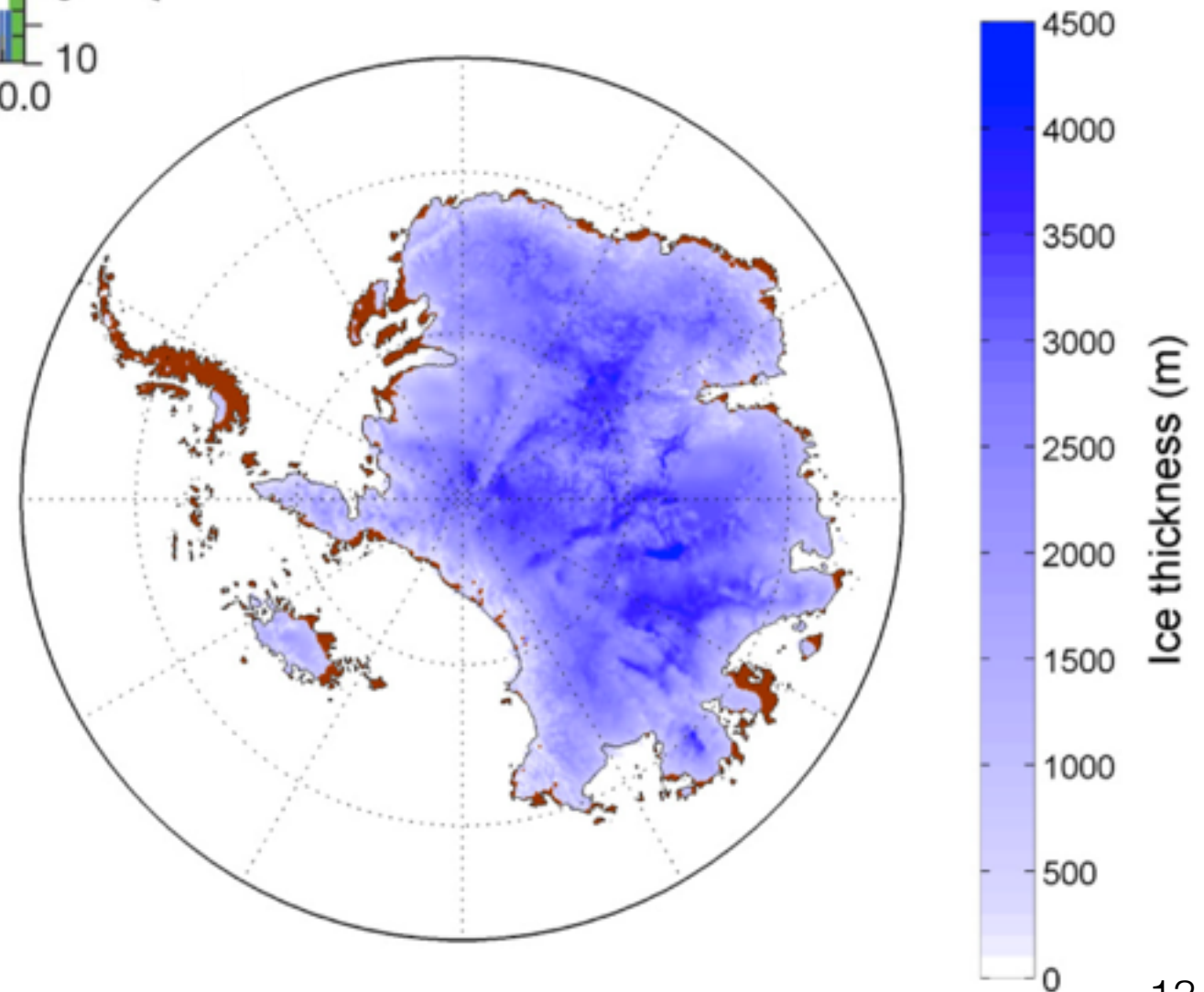
- Understanding palaeoclimates



Antarctic ice volume and sea level equivalent (Pollard and DeConto, 2009)

Pliocene ice sheet simulation DeConto and Pollard (2016)

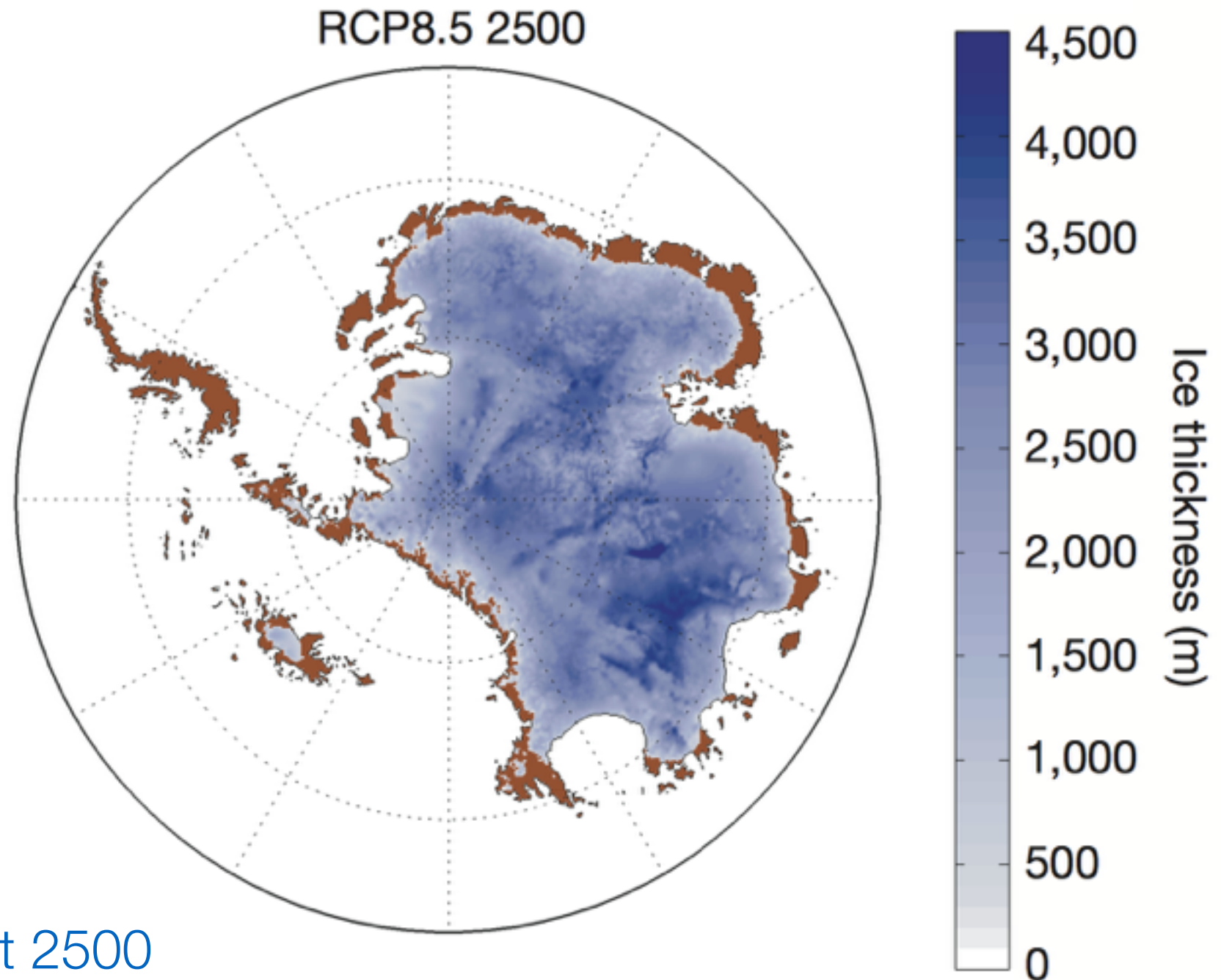
Pliocene (3Ma BP)



What is an ice sheet model used for?

- Predicting the long-term future

“Antarctica has the potential to contribute ...more than 15 m [of sea-level rise] by 2500, if emissions continue unabated. ...prolonged ocean warming will delay its recovery for thousands of years.”



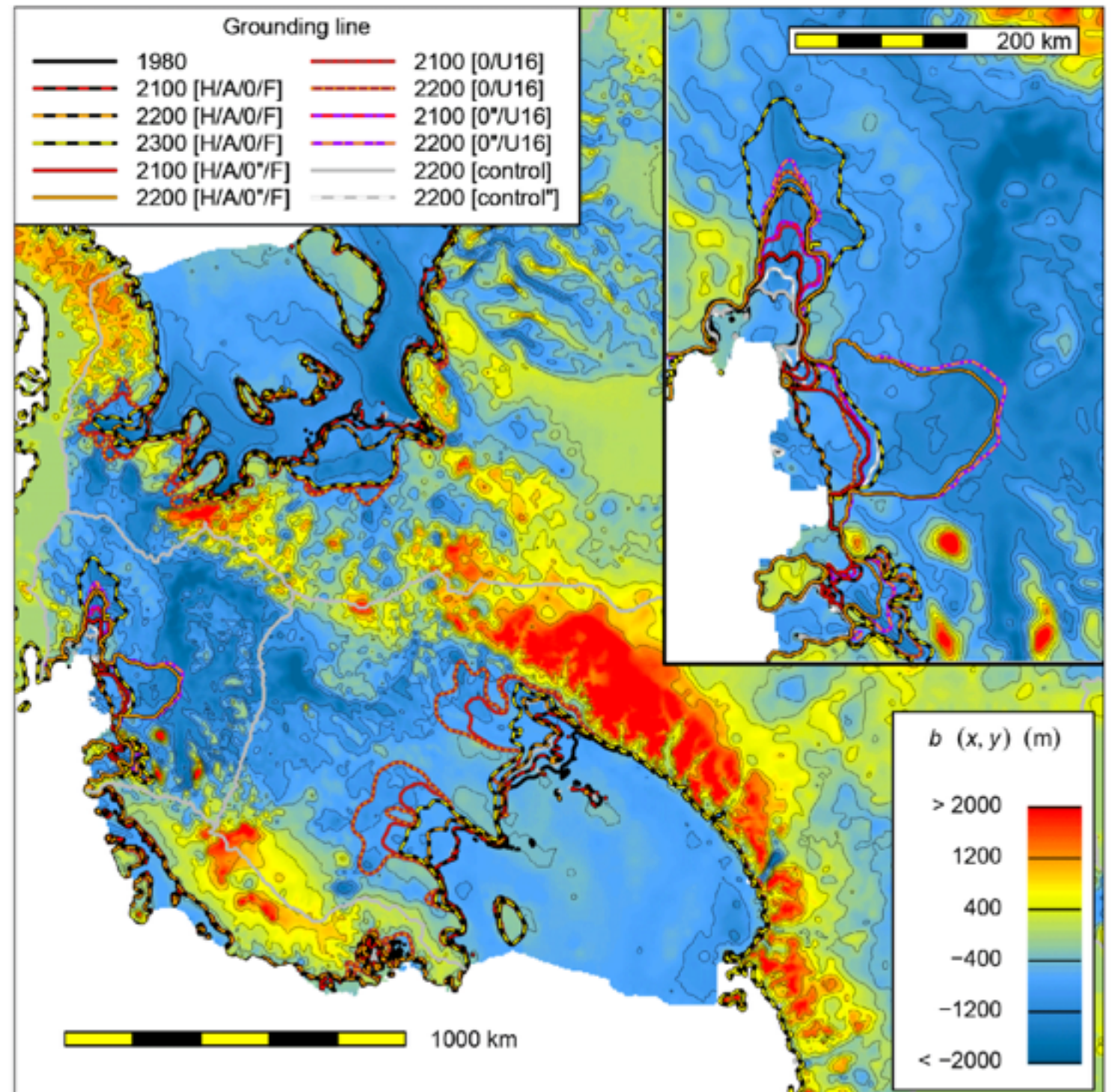
RCP8.5 ice sheet prediction at 2500
DeConto and Pollard (2016)

What is an ice sheet model used for?

- Predicting the short(ish)-term future

“Given sufficient melt rates, we compute grounding line retreat over hundreds of kilometers in every major ice stream, but the ocean models do not predict such melt rates outside of the Amundsen Sea Embayment until after 2100.”

grounding line migration in ocean-forced simulations (Cornford et al., 2015)



Ice sheet predictions for policymakers

- Ice Sheet Model Intercomparison Project for CMIP6
 - CMIP = Coupled Model Intercomparison Project Phase 6



- IPCC Sixth Assessment Report
 - Published from 2020
- Range of complexities and computational expense
 - Physics: full Stokes, various approximations
 - Spatial resolution and domain
- **Standalone** and coupled with climate models

How to use observations with models?

- Combining and comparing observations with models
 - Both are **imperfect**
 - Different spatial resolution, domains, variables
- To obtain best possible estimates of:
 - system state: past, present and future ice sheet
 - model parameters



All models
are WRONG

...but some are
USEFUL

How to use observations with models?

- Formal methods often derived from **Bayes Theorem**

1. model simulation(s) of state
2. compare with observations
3. update estimate of state and/or parameters



- e.g.

- Data Assimilation (state)
- Bayesian calibration (parameters)

- incorporating obs into simulations
- estimating parameters from obs with Bayesian inference

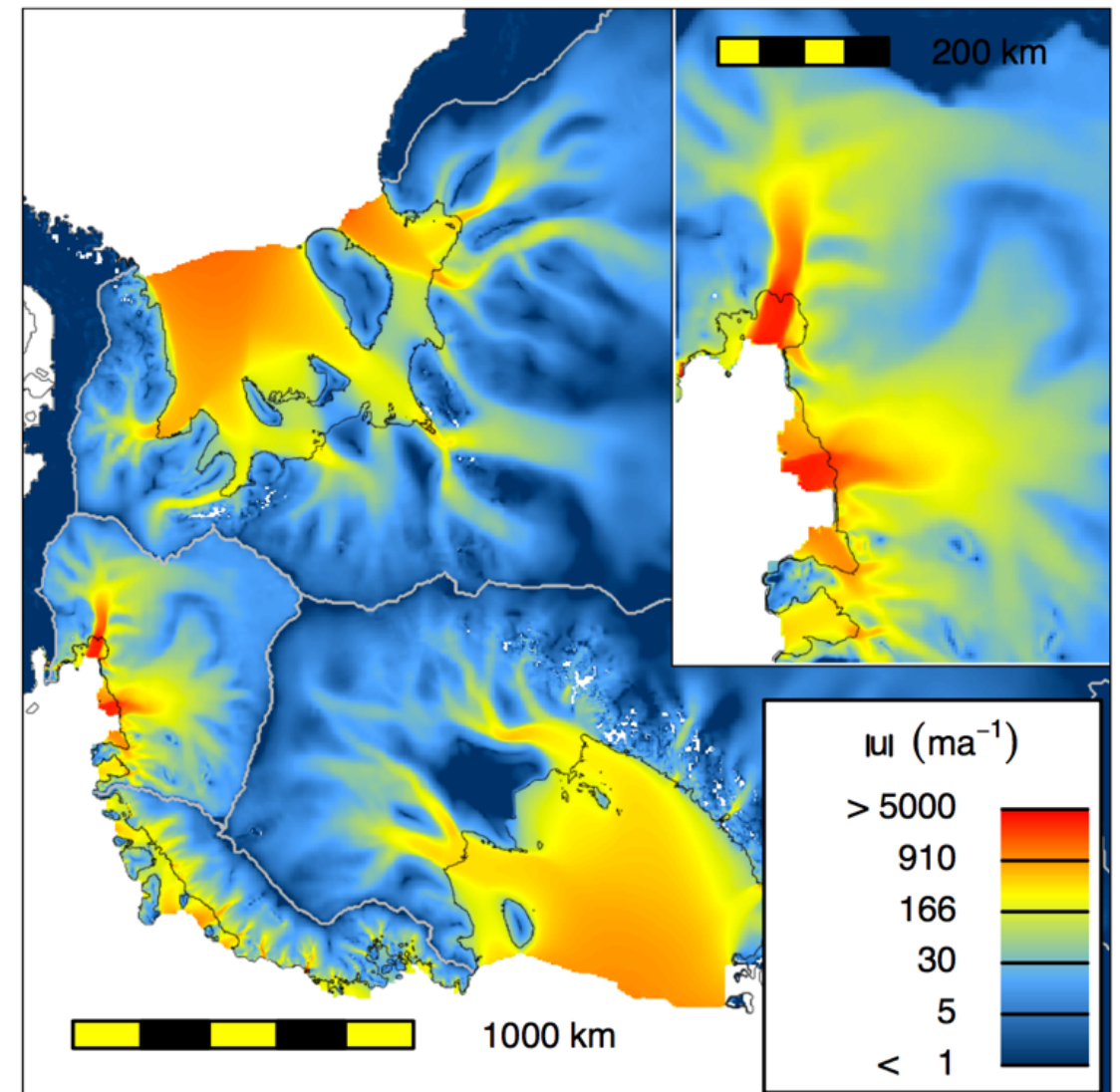
- Less formal methods also used...

- 'nudging'
- 'relaxation'
- hand tuning

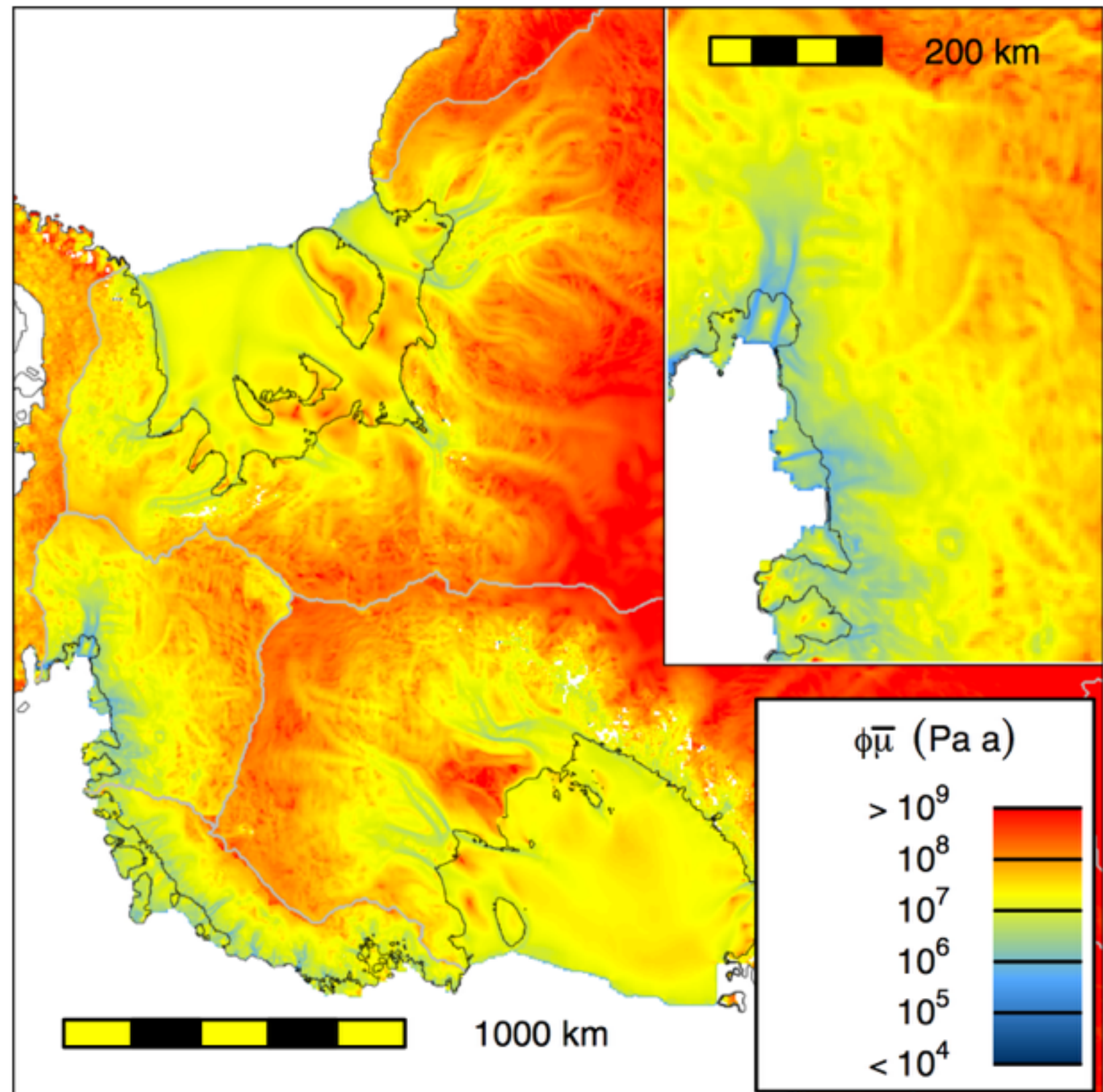
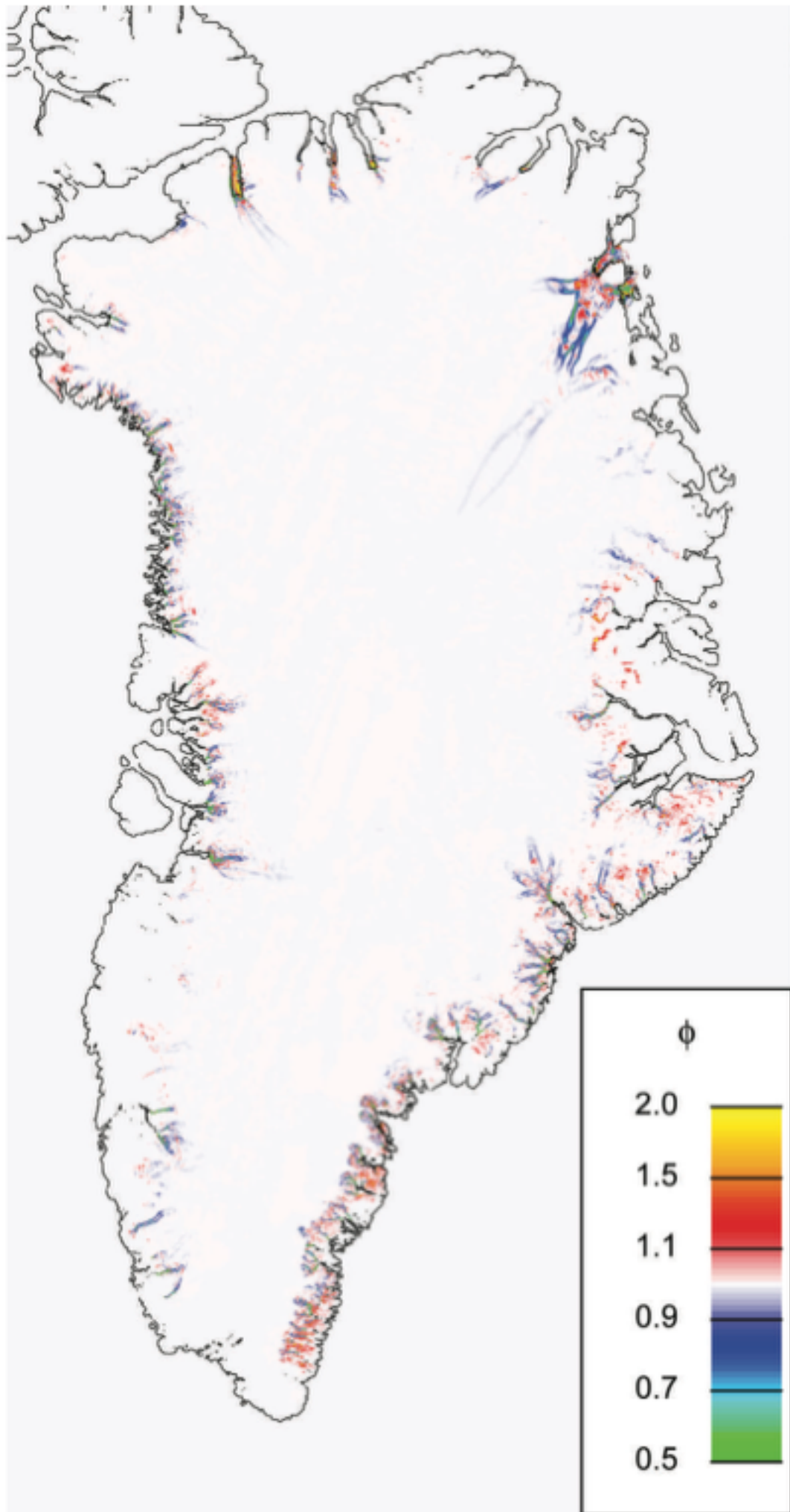
What does an ice sheet model need?

1. Initial state

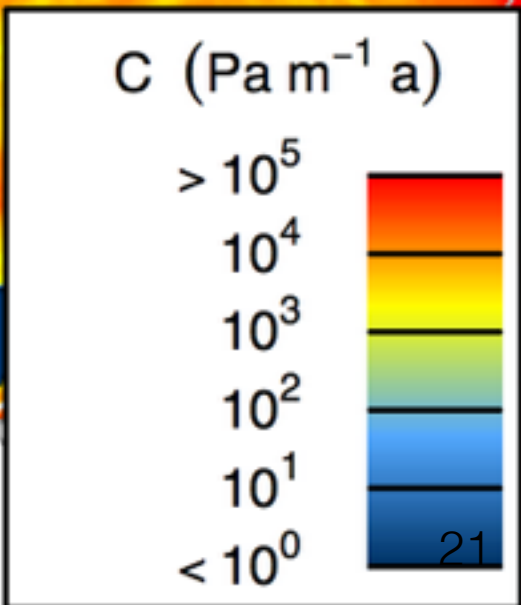
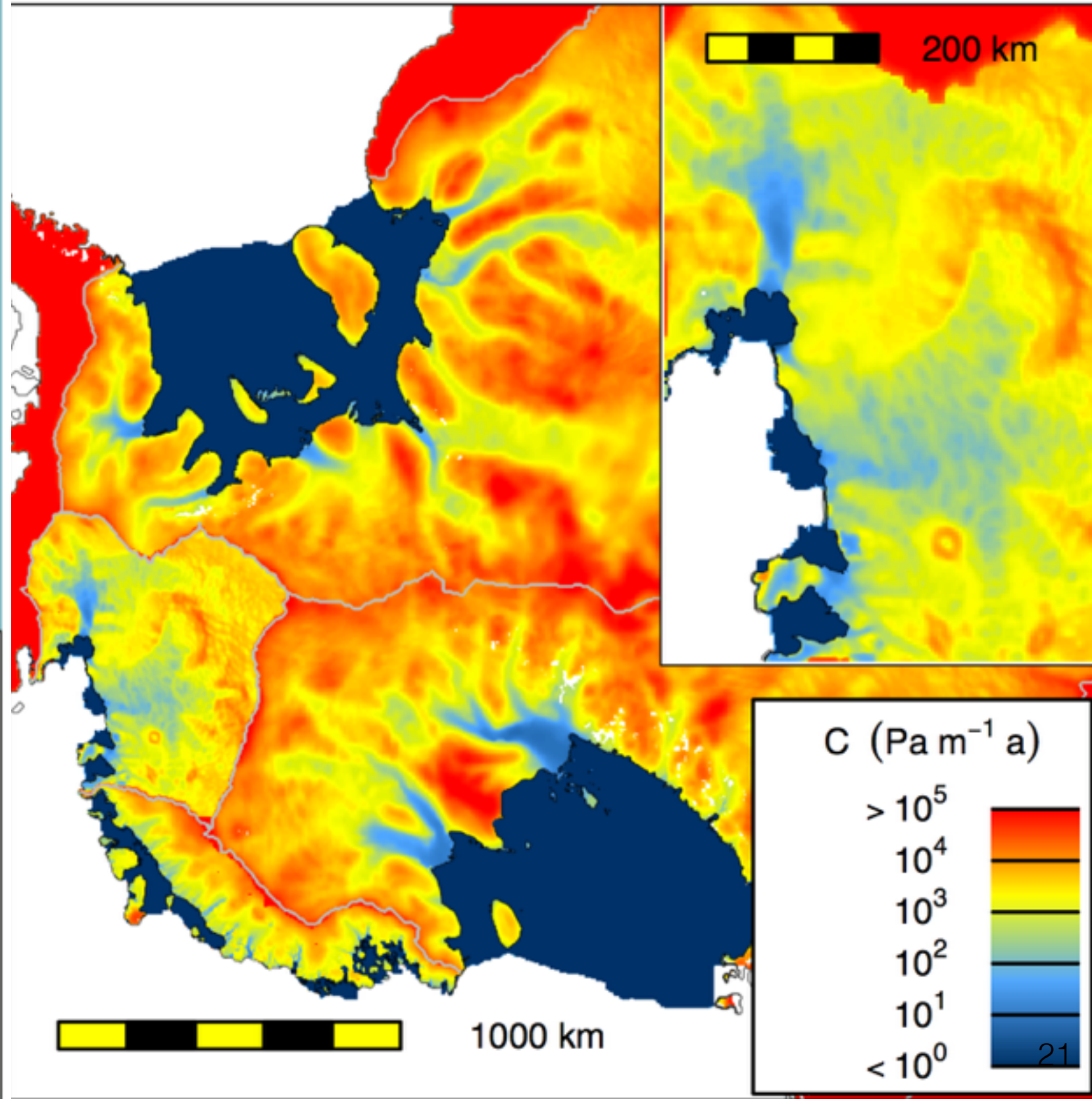
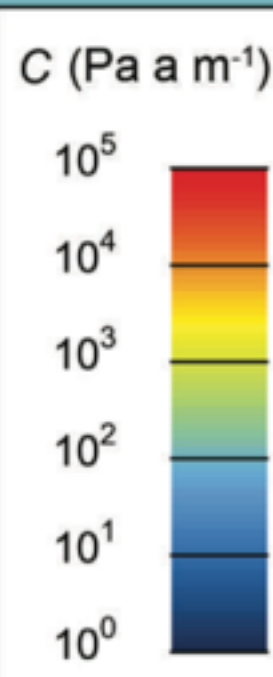
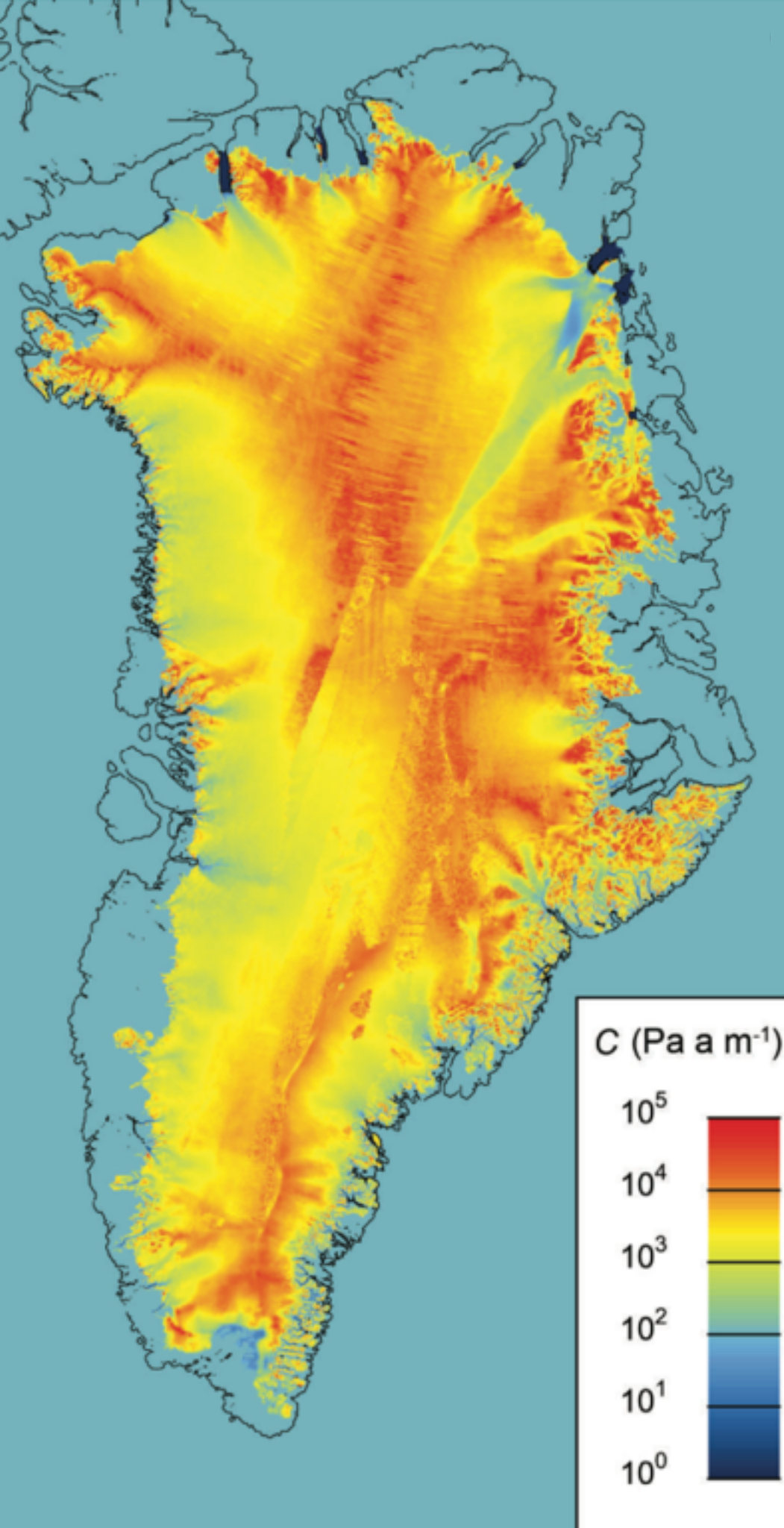
- ice sheet geometry
- ice velocity
- internal ice temperature
- basal traction coefficient
- maybe others, e.g.:
 - ▶ enhancement factor
 - ▶ effective viscosity, stiffness
 - ▶ bedrock topography corrections due to obs uncertainties
 - ▶ mass balance corrections to prevent artefacts/drift



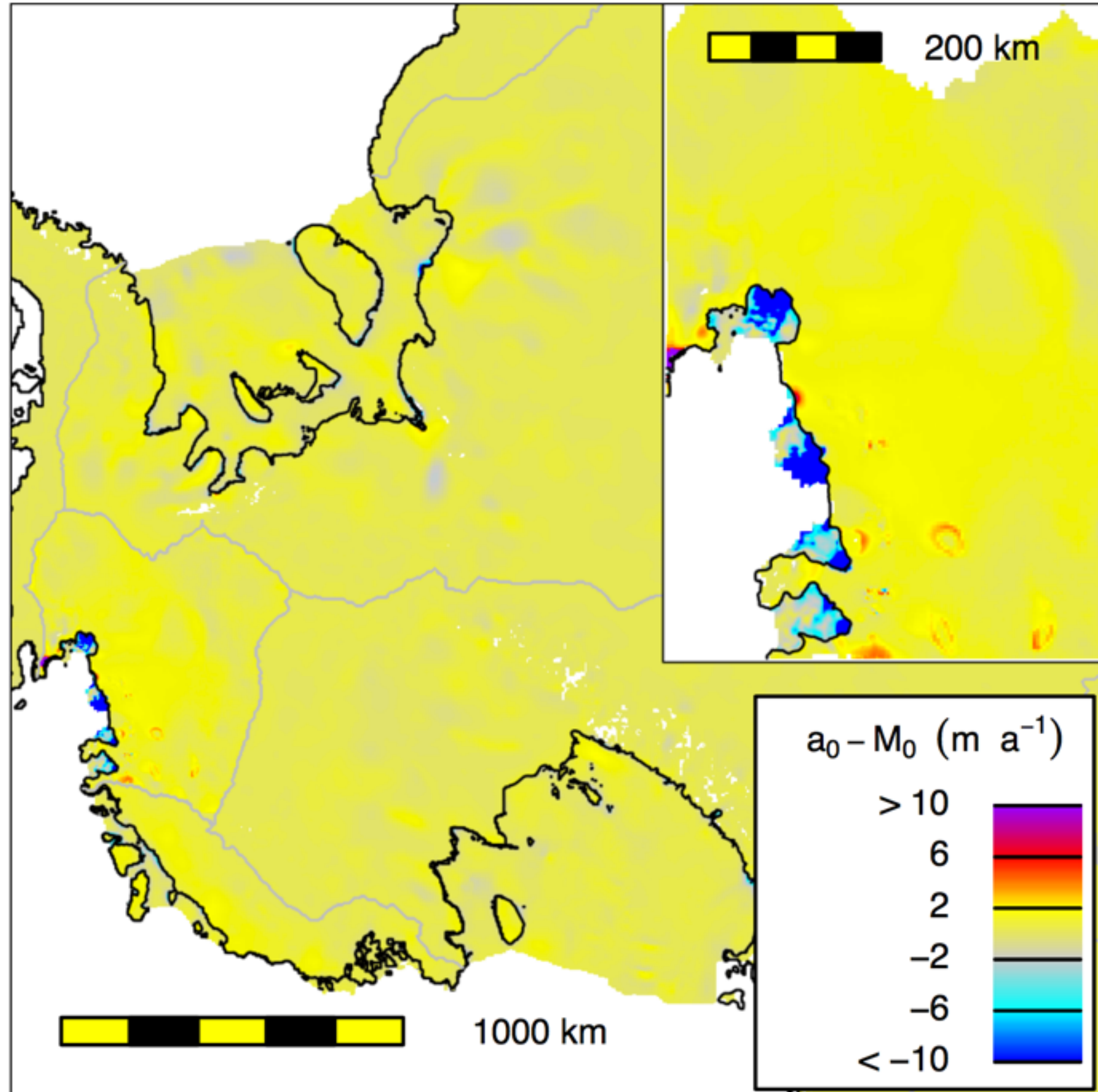
initial ice flow speed again
(Cornford et al., 2015)



basal traction coefficient
(Lee et al., 2015; Cornford et al., 2015)



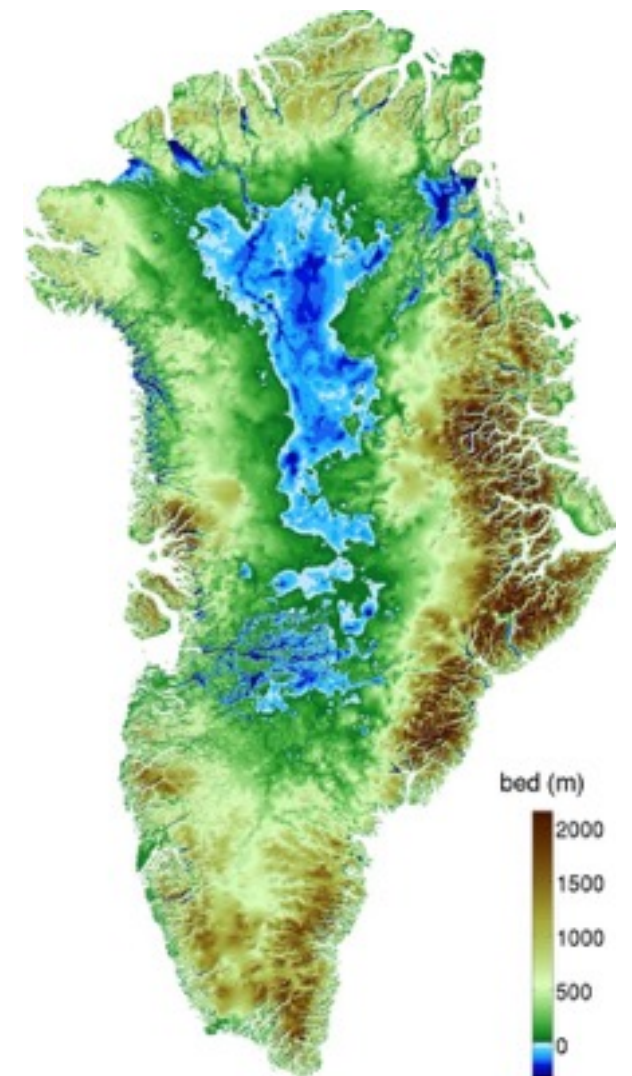
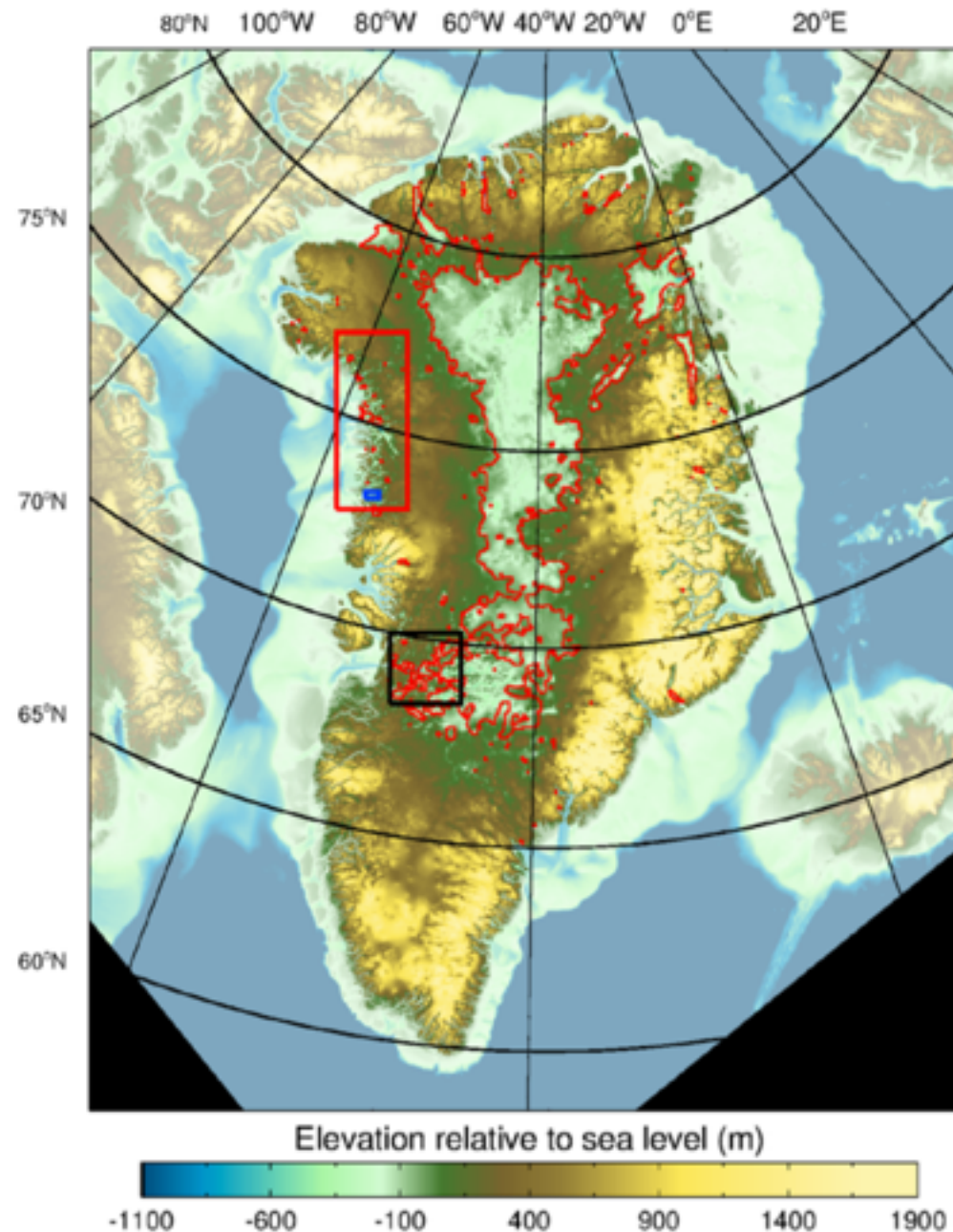
synthetic mass balance
(Cornford et al., 2015)



What does an ice sheet model need?

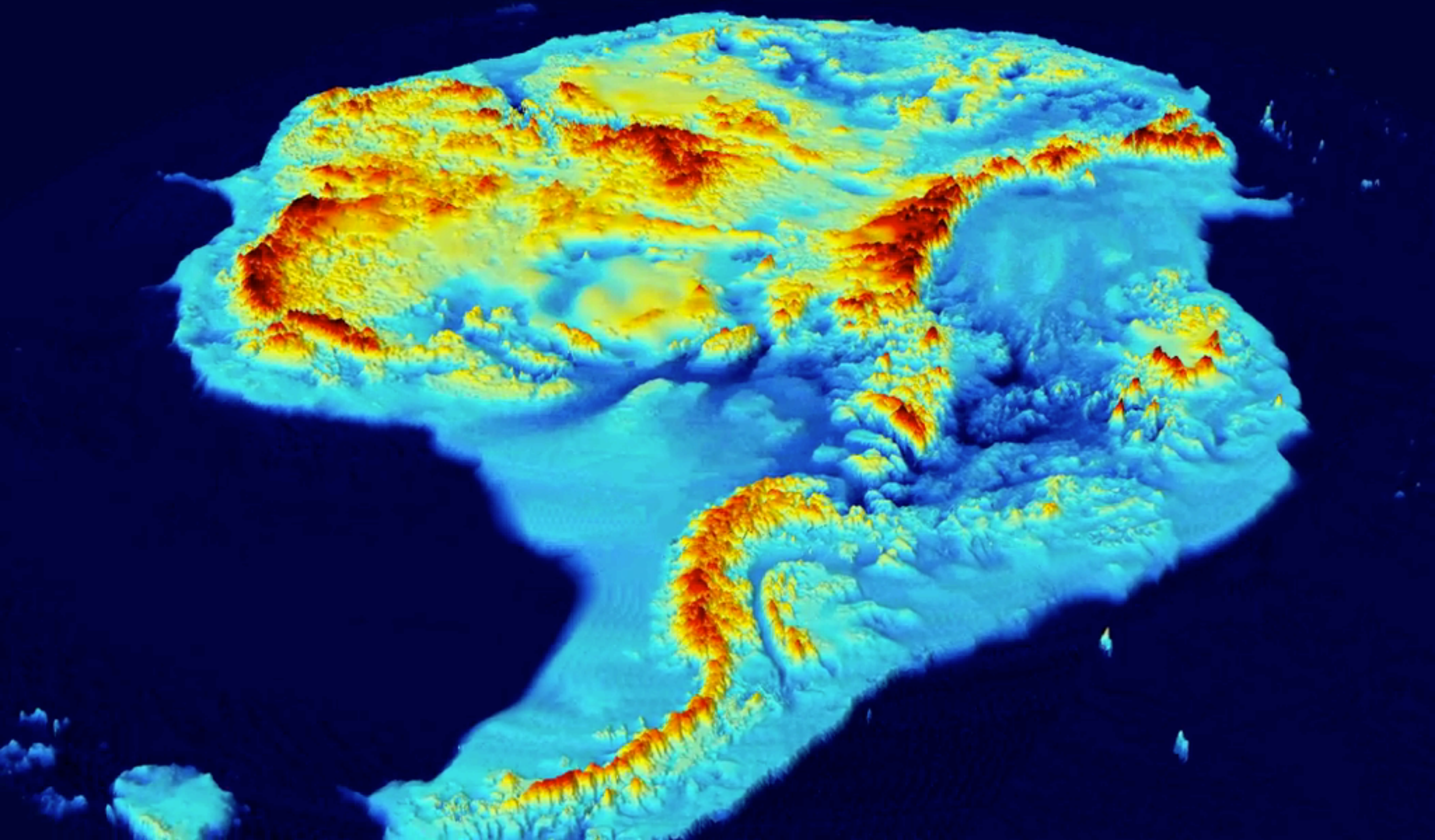
2. Boundary conditions

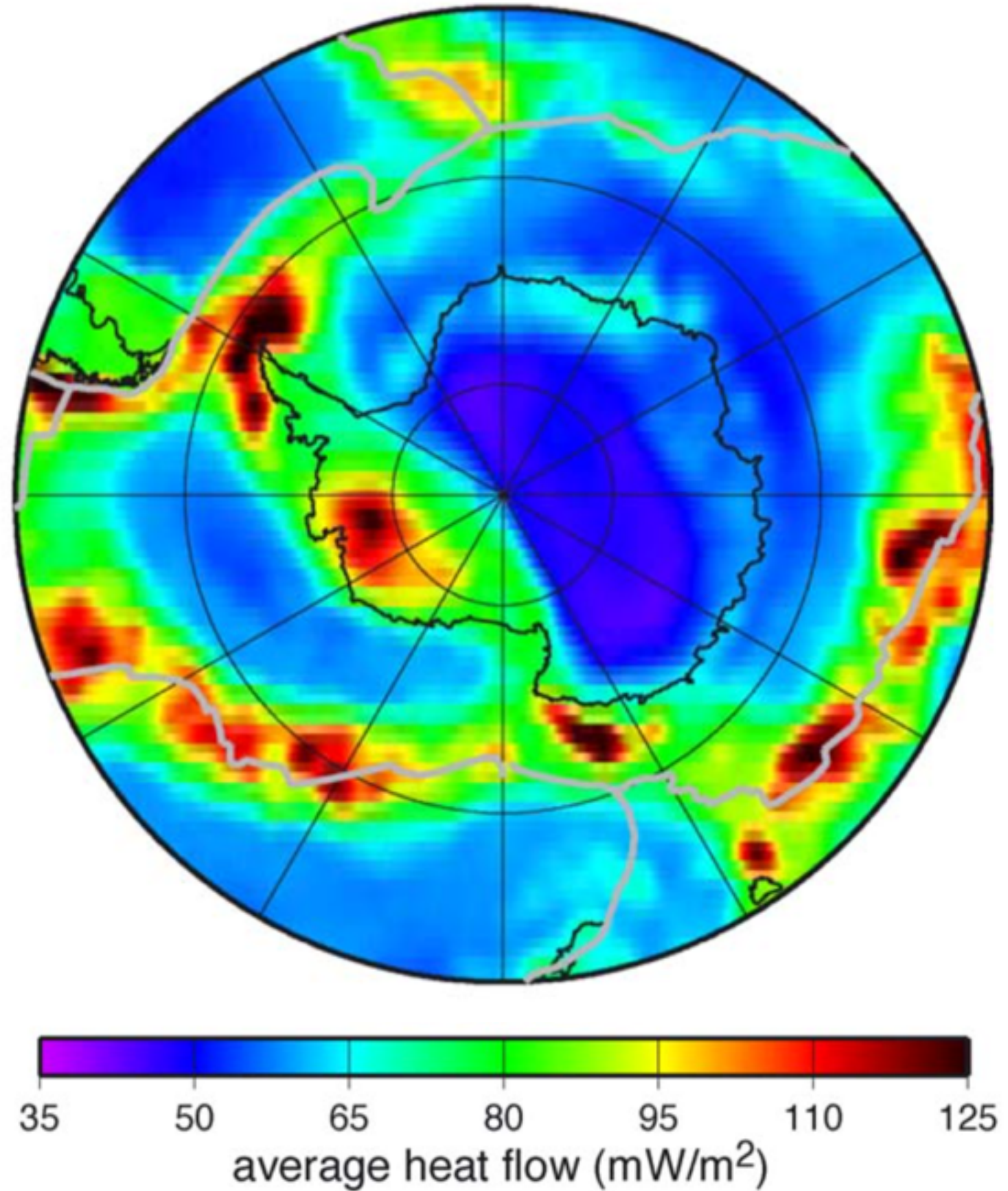
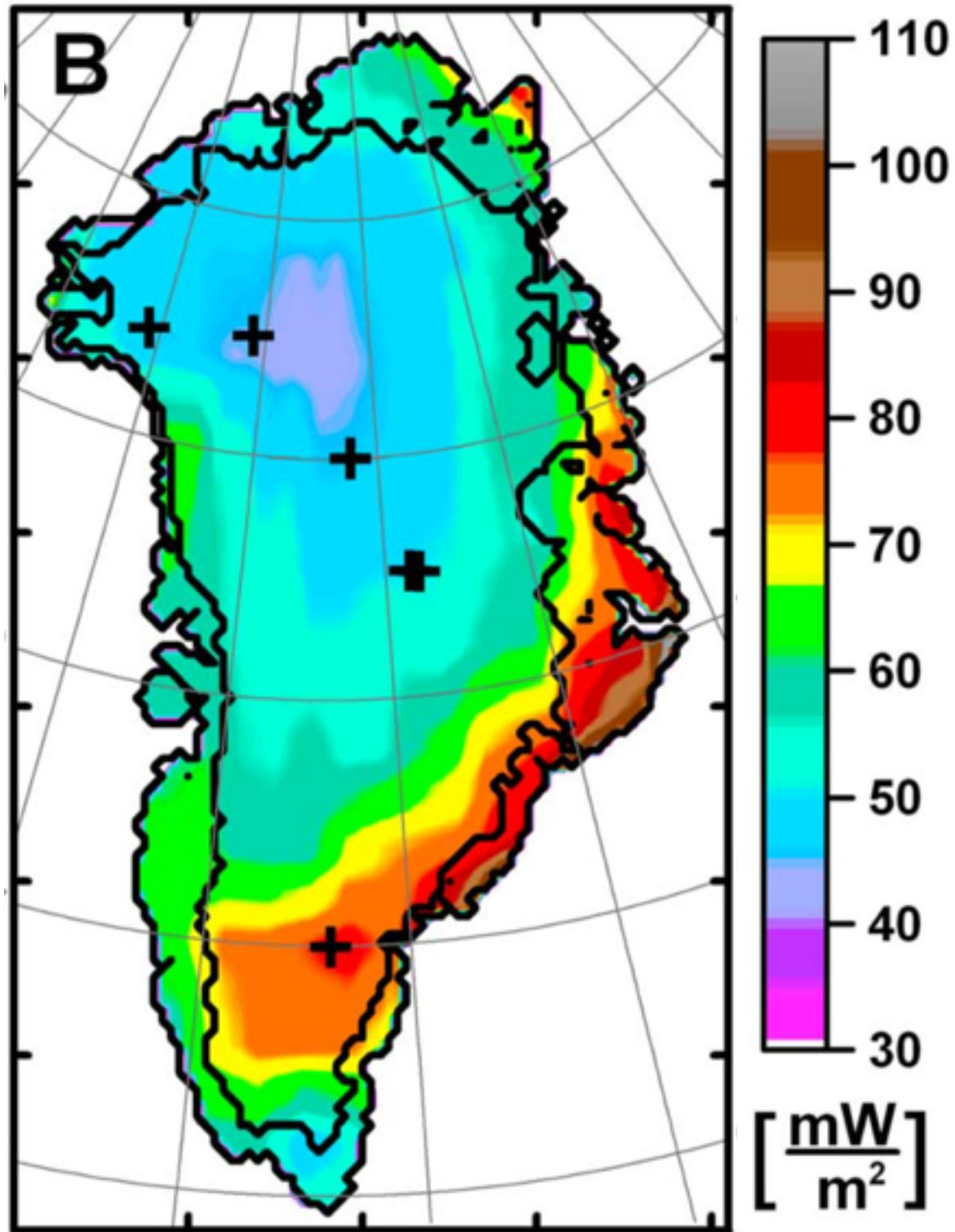
- bedrock topography, geothermal heat flux



bedrock elevation
(Bamber et al., 2013;
Morlighem et al., 2014)

bedrock elevation: BEDMAP2
(Fretwell et al., 2013)





geothermal heat flux
(Shapiro et al., 2004;
Rogozhina et al., 2012)

What does an ice sheet model need?

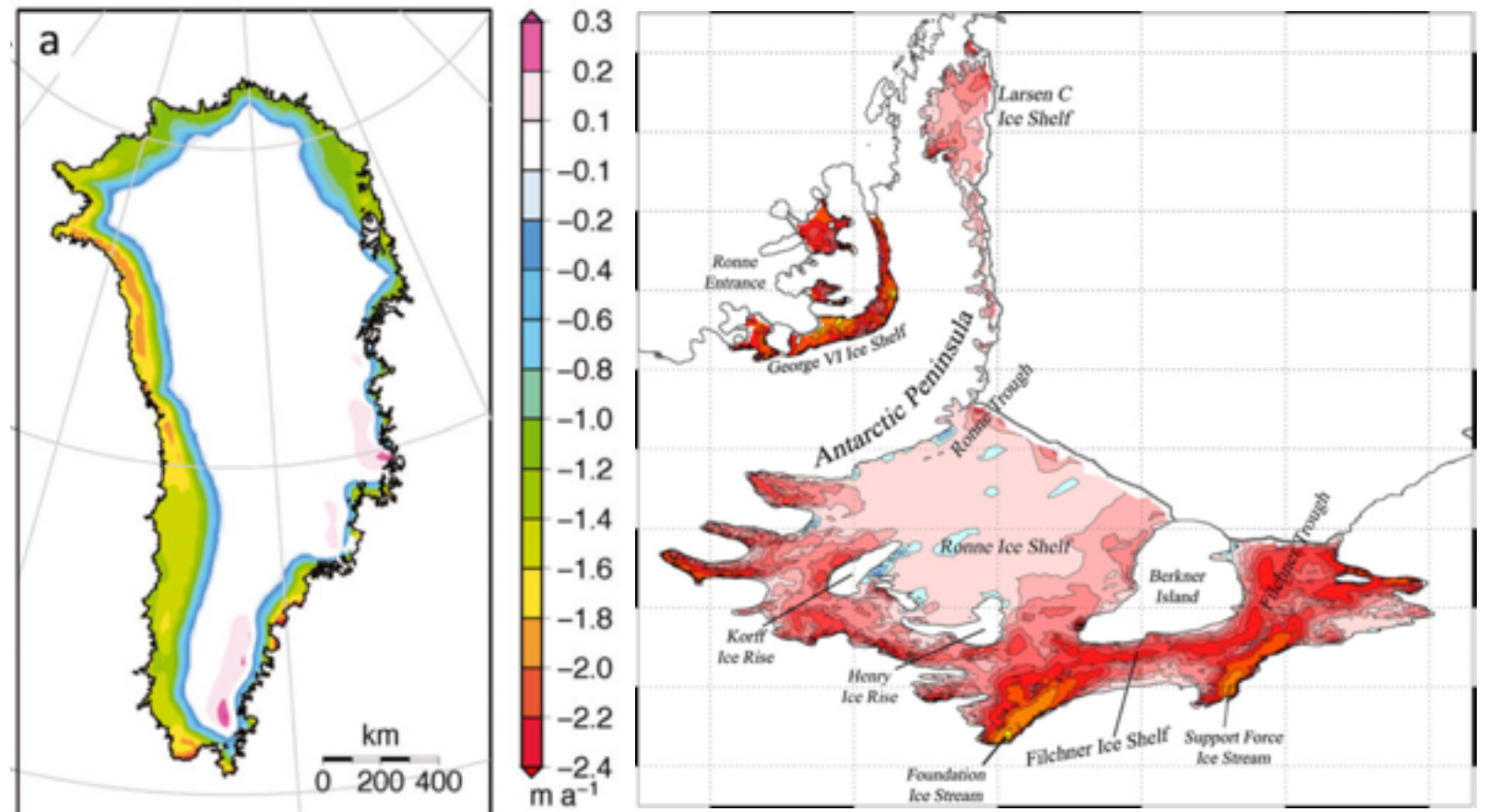
3. Climate forcing or mass balance

- atmosphere:
 - ▶ temperature & precipitation, or
 - ▶ surface mass balance (SMB)
- ocean
 - ▶ temperature, or
 - ▶ basal melting

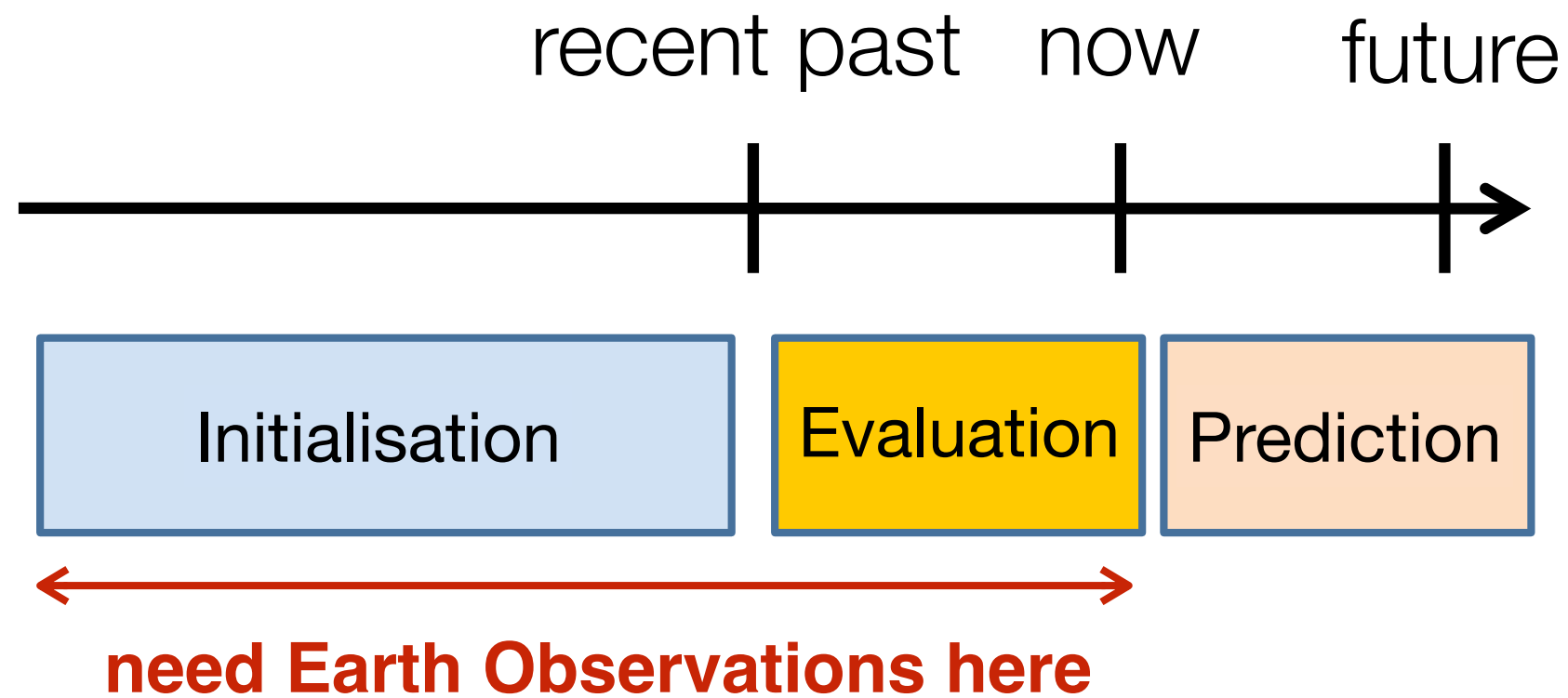
- observations
- climate models
- ice cores
- schematic

Basal melt rates in 2140-2149 under SRES scenario A1B (Timmerman & Hellmer, 2013)

Regional climate model SMB anomalies for 2091-2100 relative to 1989-2008 under SRES scenario A1B (Goelzer et al., 2013)



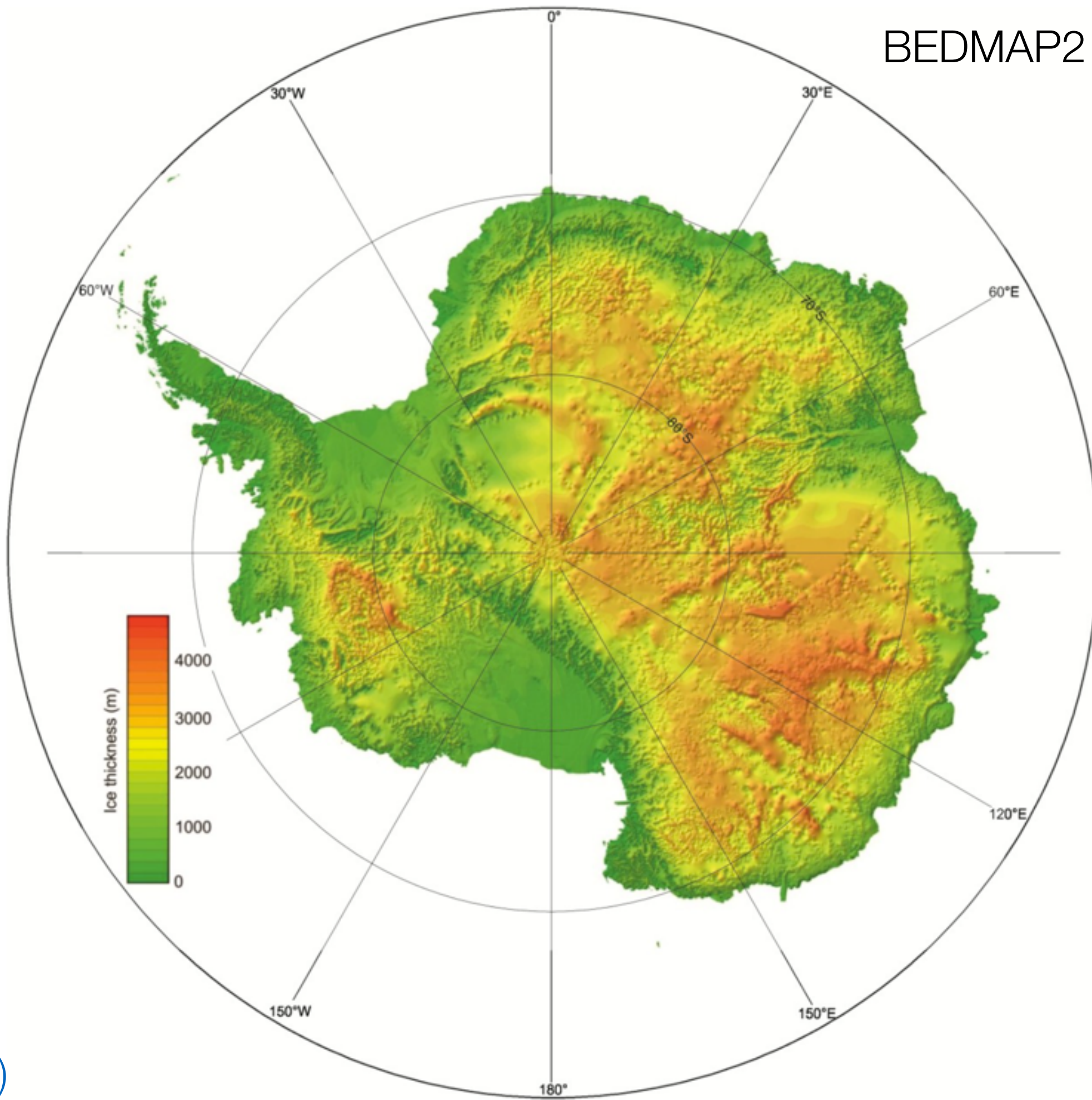
Earth Observations



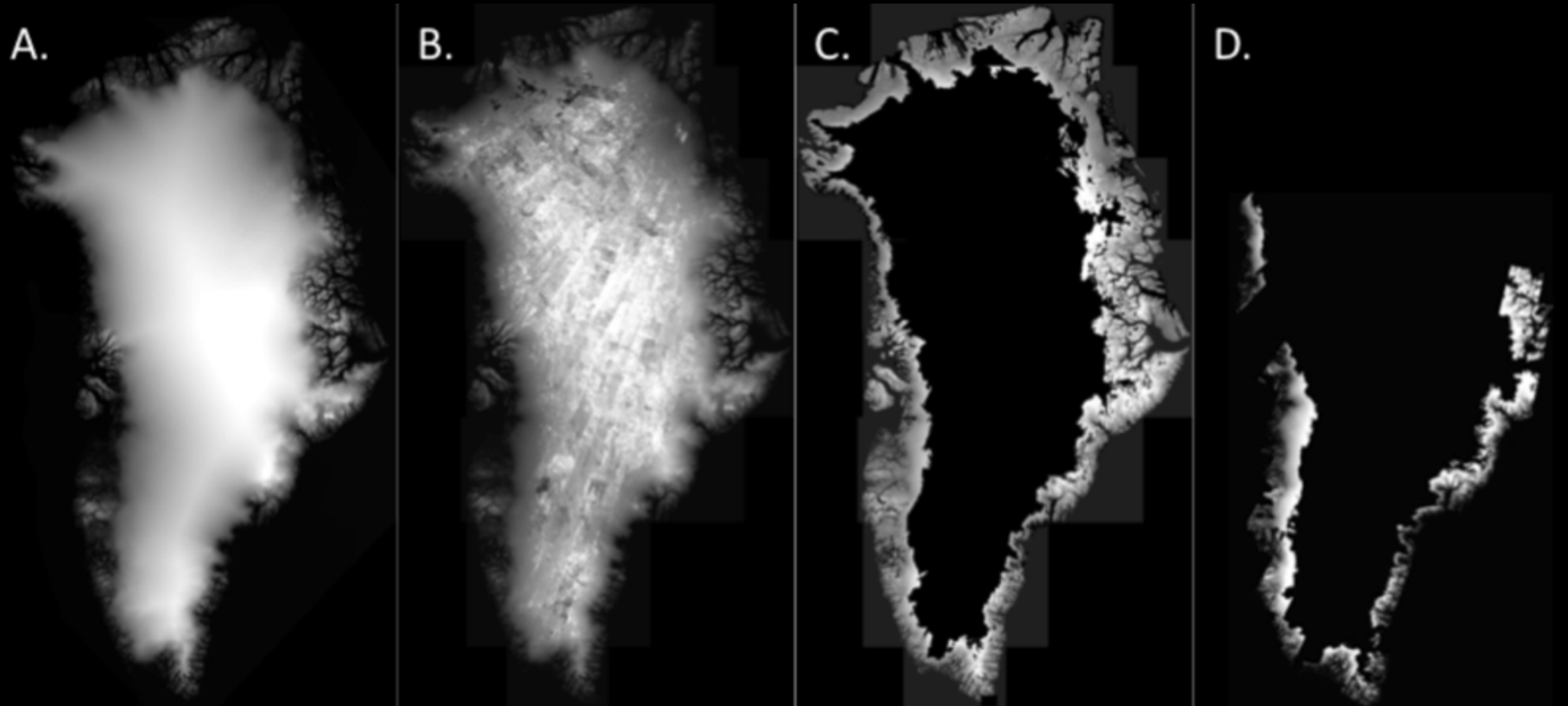
Initialising an ice sheet model

- Need to find self-consistent values for ice sheet state
 - geometry, flow, ice temperature, basal traction coefficient, ...
- Consistent with observations and reconstructions
 - EO: geometry, elevation changes, velocities
 - recent climate, reconstructed palaeoclimate
- Even though both are imperfect
- Data assimilation of various kinds, e.g.:
 - tuning and inverse methods to estimate basal traction coefficient from surface or balance velocities
 - setting geometry equal to observed, then allowing model to 'relax' to quasi-equilibrium state

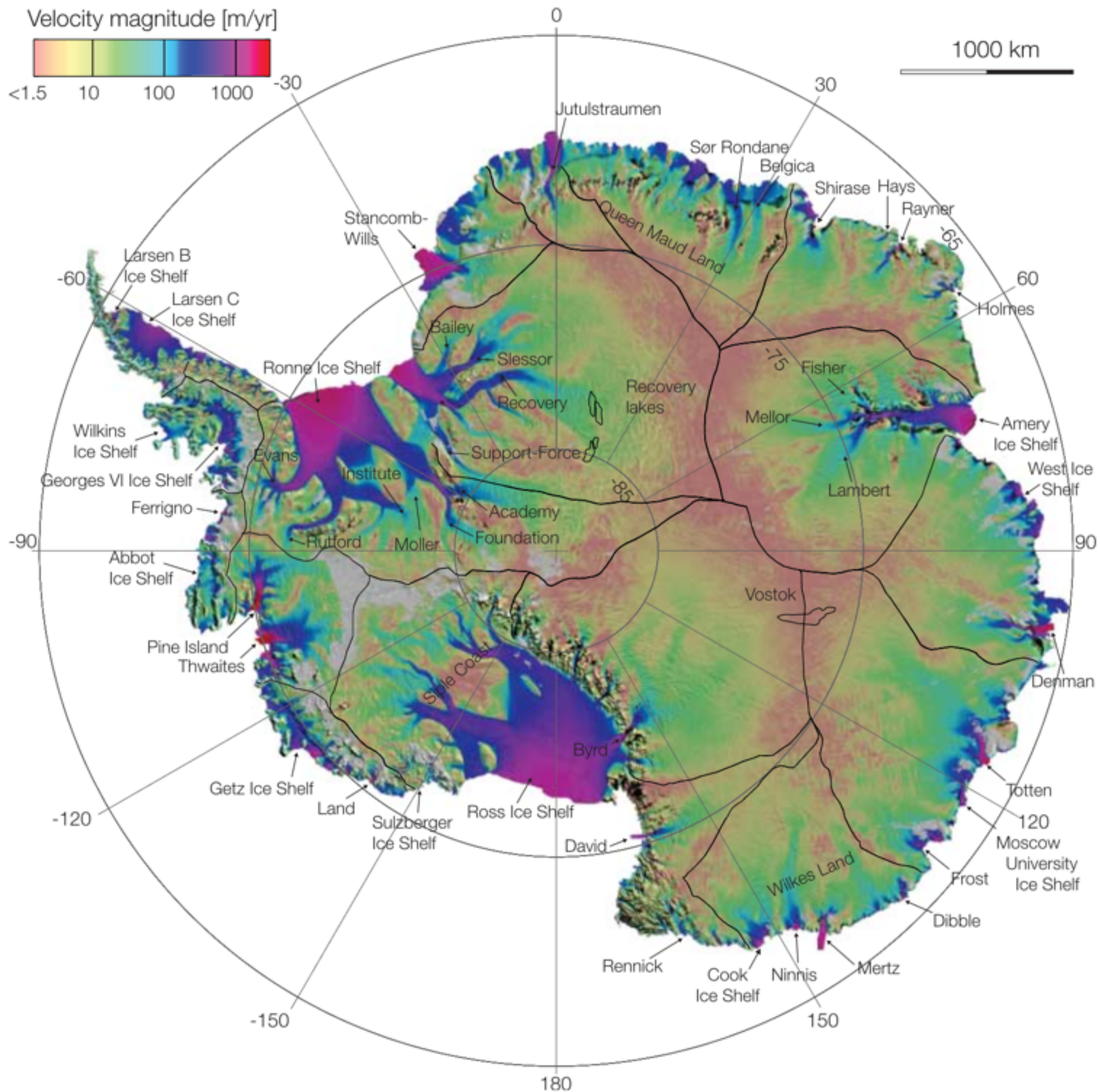
BEDMAP2



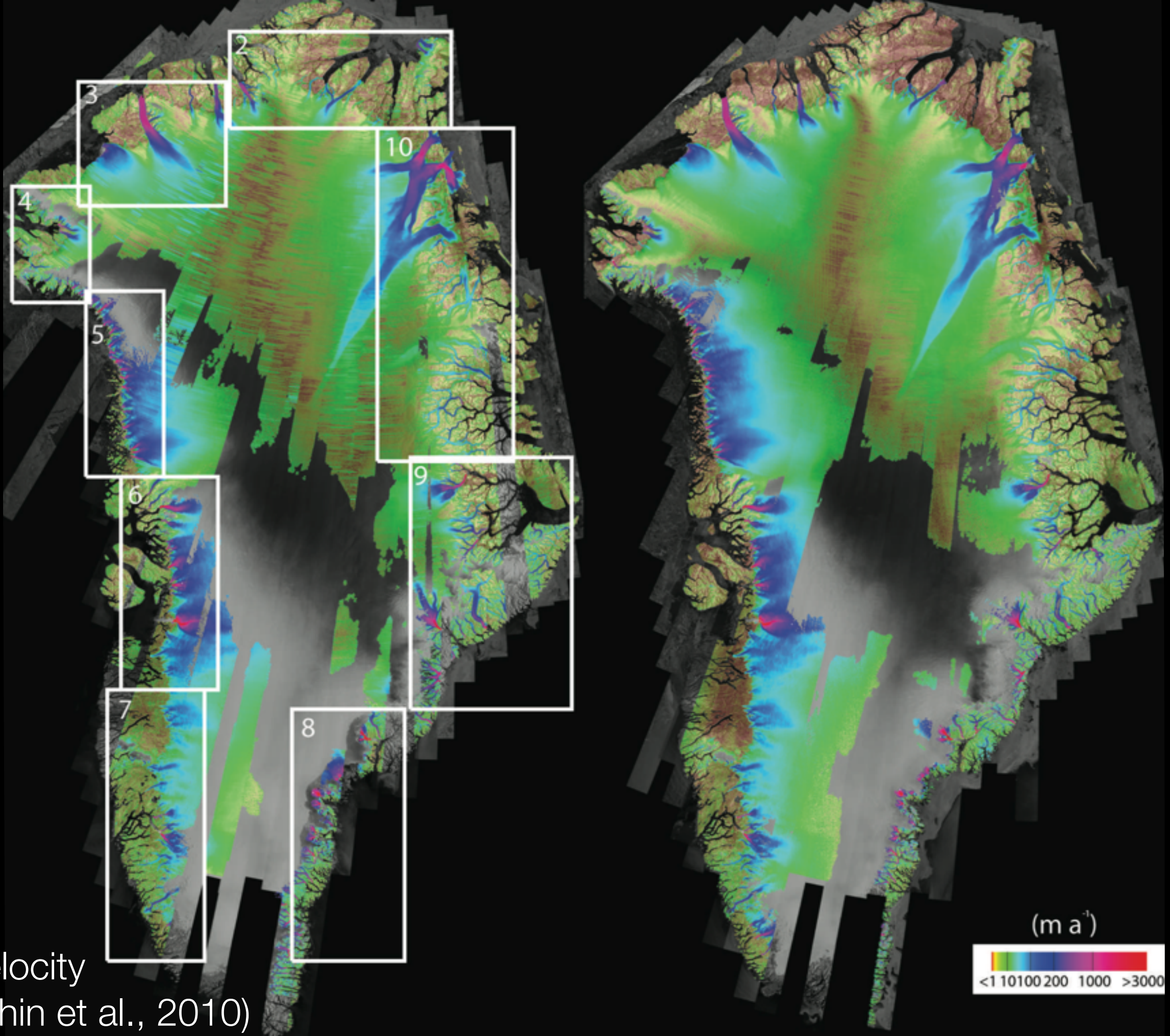
ice thickness
(Fretwell et al., 2013)



surface elevation
(Howat et al., 2014)



ice velocity
(Rignot et al., 2011)



ice velocity
(Joughin et al., 2010)

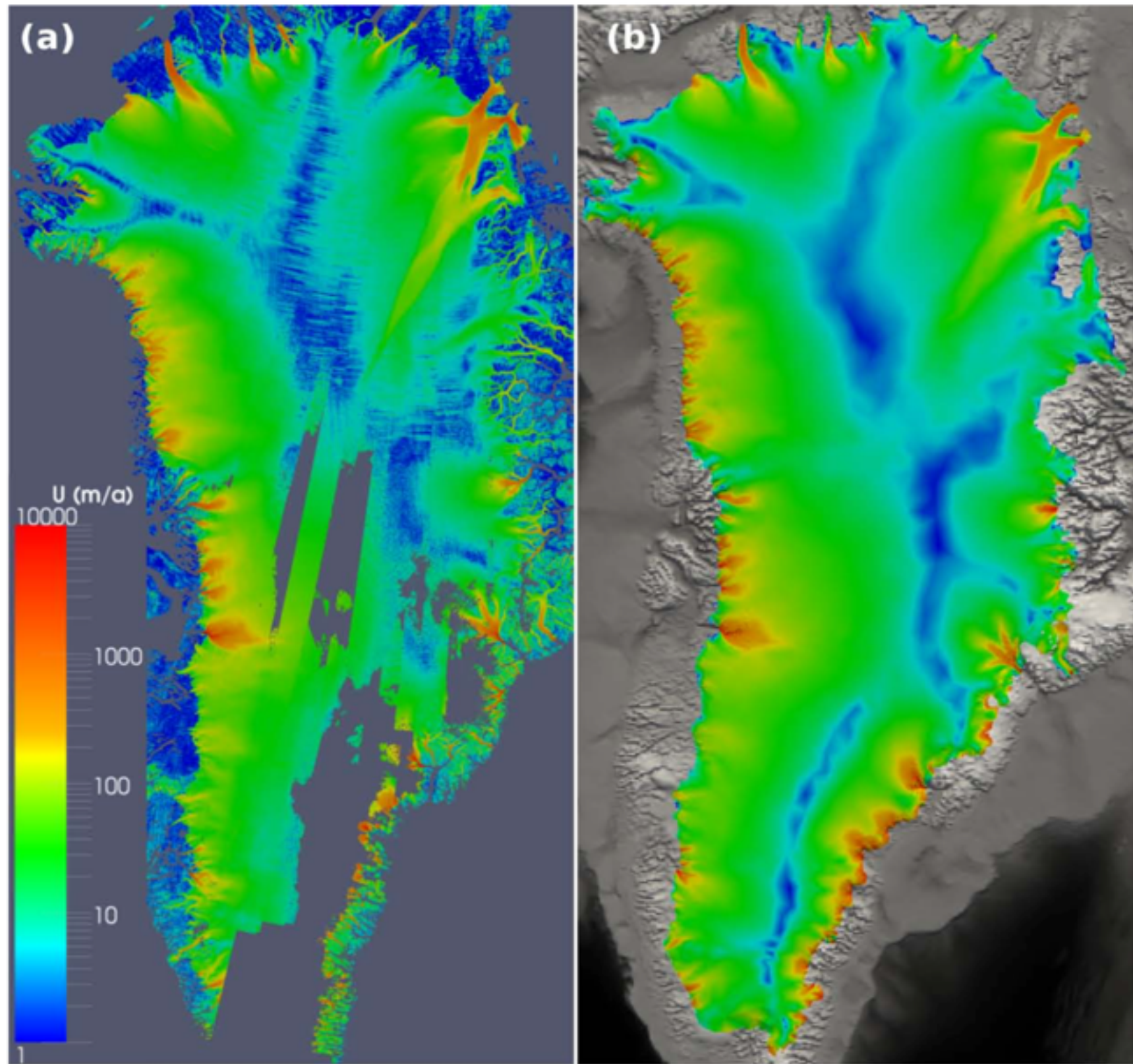
Estimating basal traction coefficient

minimise mismatch
between modelled and
observed velocities

e.g. cost function

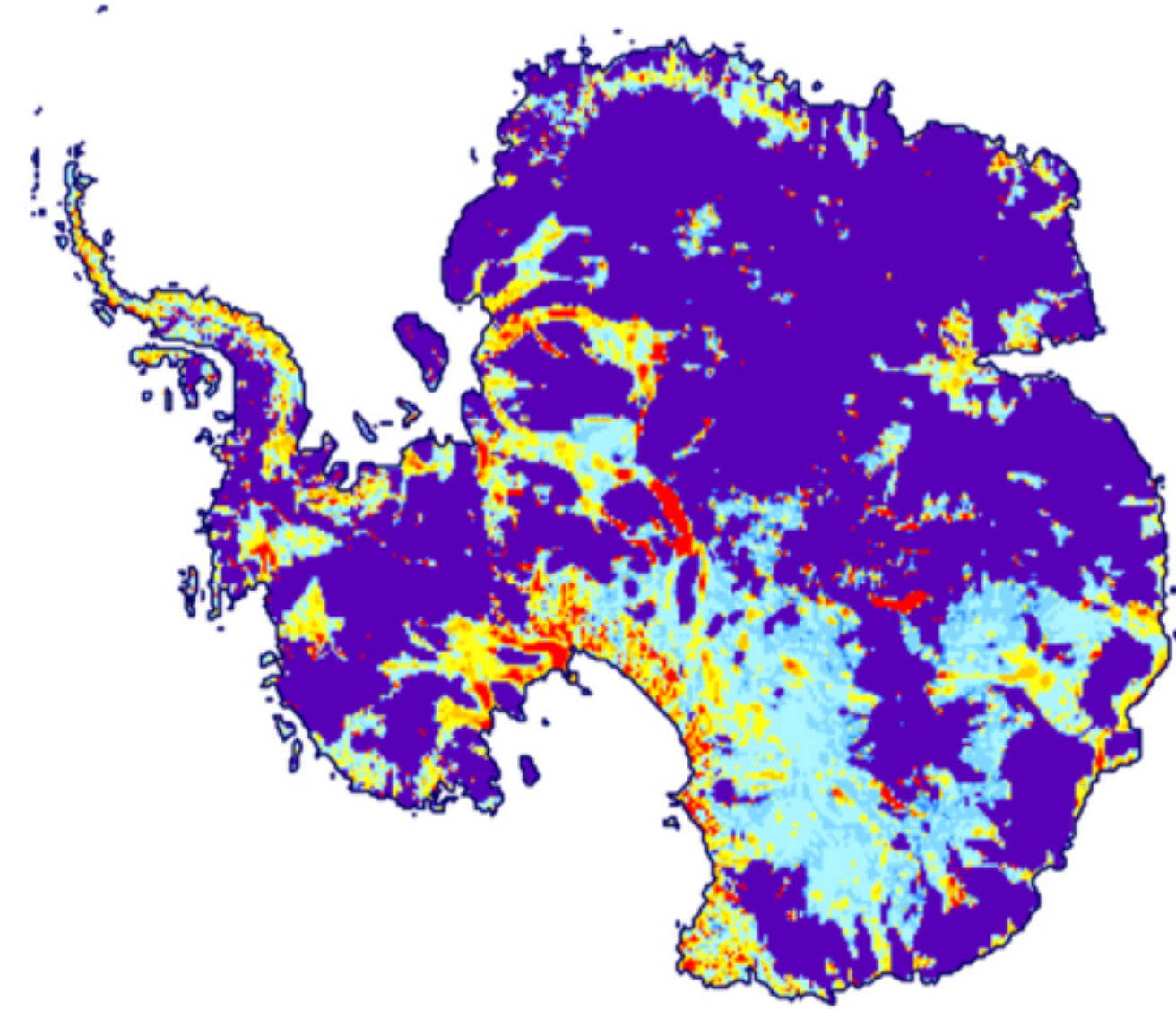
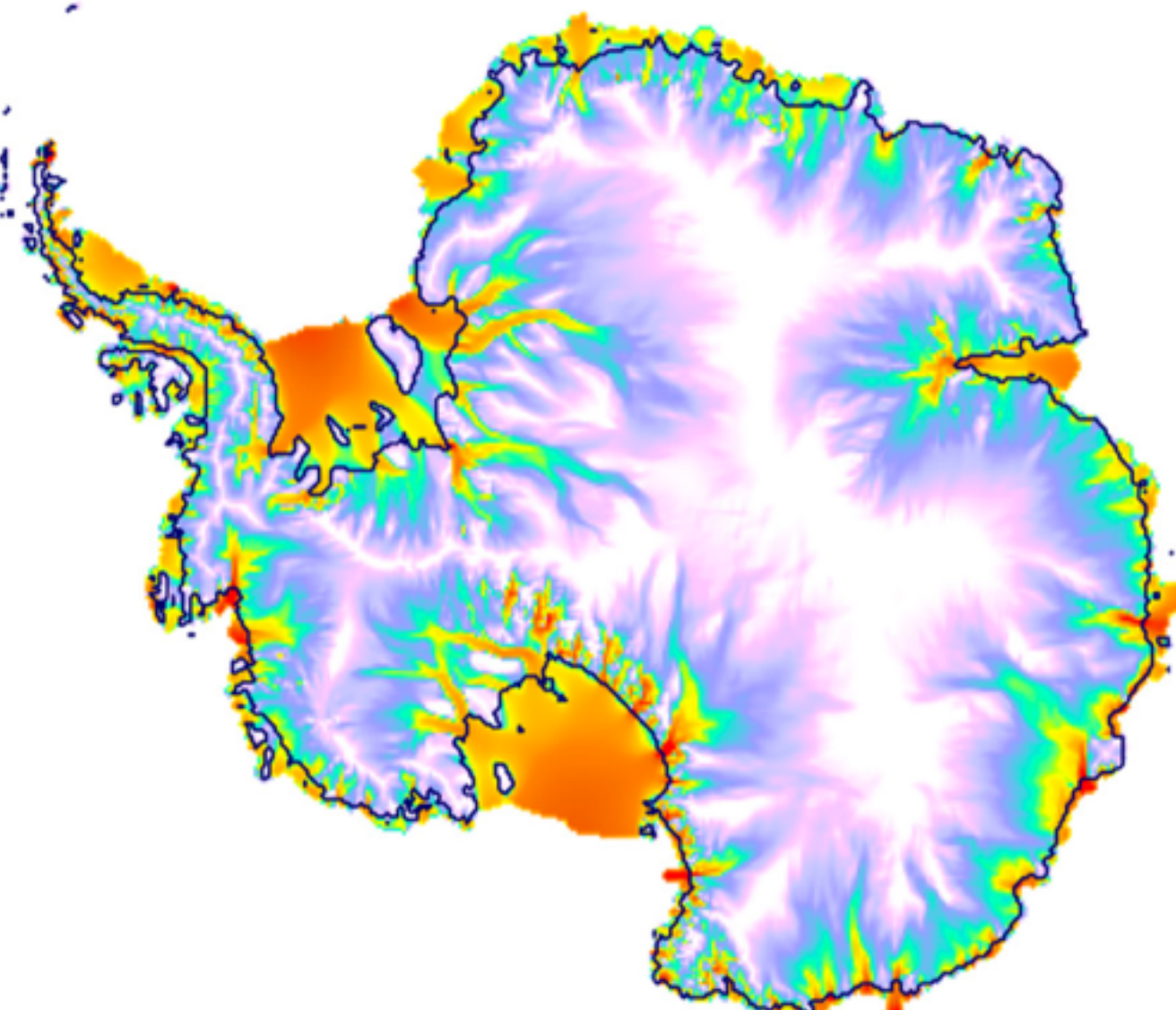
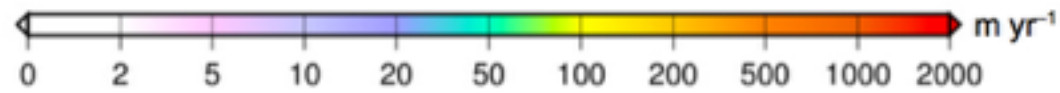
$$J_o = \int_{\Gamma_s} \frac{1}{2} \left(|\mathbf{u}_H| - |\mathbf{u}_H^{\text{obs}}| \right)^2 d\Gamma,$$

observed and initial model
surface velocities
(Gillet-Chaulet et al., 2012)



Estimating basal traction coefficient

inversion gives estimate of basal traction coefficient



initial average velocities
(Ritz et al., 2015)

initial log(effective basal traction coeff.)

Large initialisation uncertainties

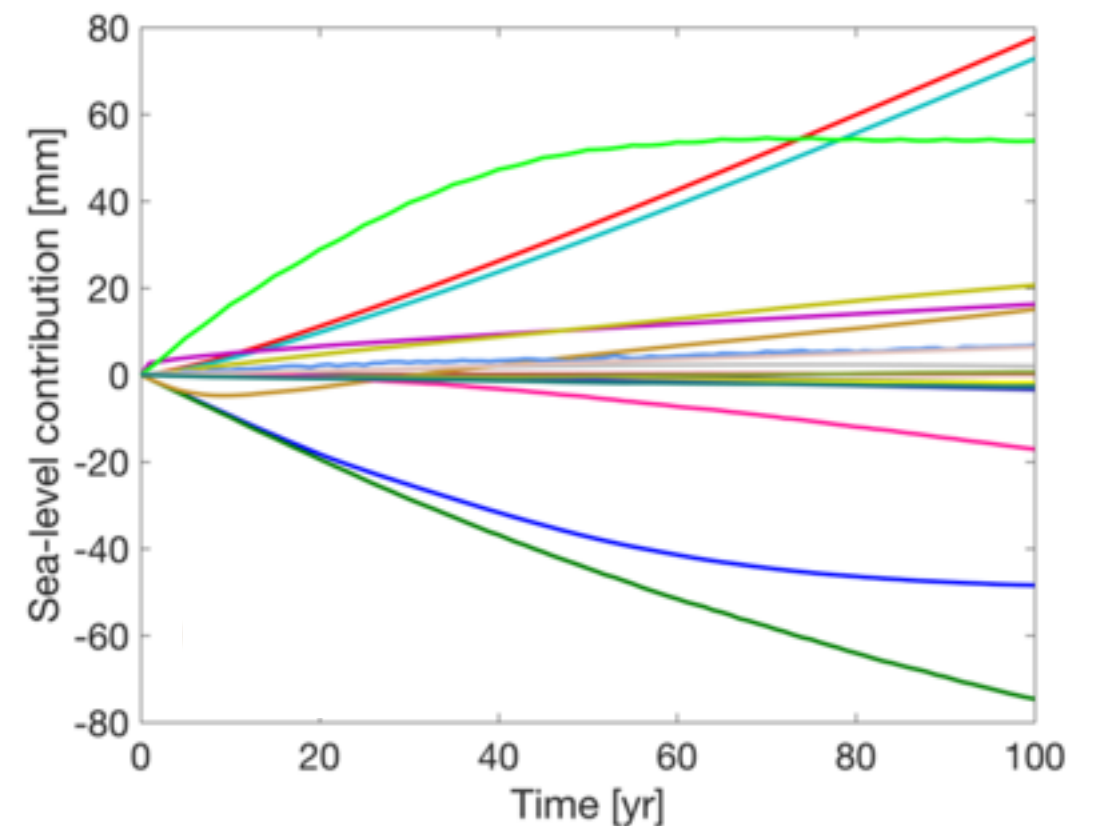
- Different methods
 - formal vs ad-hoc
 - free vs fixed geometry spin-up
 - glacial-interglacial cycle(s) vs recent climatology
 - mass balance corrections vs subtracting drift from predictions
- Different datasets and time periods
 - sometimes multiple variants
 - mismatches in time coverage
 - definition of “recent” climatology
- Different model structures
 - derive different initial states even if same method and data

	CISM	Elmer/Ice	GISM-HO	GISM-SIA	GRISLI	MPAS
DERIVATION OF ICE TEMPERATURE						
Spin-up simulation	Quasi-SS; fixed geom. (B13)	One g-ig cycle with SICOPOLIS (G97)	Two g-ig cycles; IT rescaled to obs. thick.	Two g-ig cycles; IT rescaled to obs. thick.	Quasi-SS; fixed geom. (B13)	Quasi-SS; fixed geom. (B13)
Spin-up SAT	E09, constant	E09, constant	Two g-ig cycles, evolving	Two g-ig cycles, evolving	E09, constant	E09, constant
BASAL DRAG CALIBRATION						
Method	Tuning	Control	n/a	n/a	Iterative inverse	Tuning
Target velocities	Balance	Surface	n/a	n/a	Surface	Balance
INITIALISATION						
Relaxation	n/a	55 years	1000 years (restrictions)	1000 years (restrictions)	200 years	n/a
Drift	Synthetic	Control	Synthetic	Synthetic	Control	Synthetic
Climate SMB Dates	ERA-I MAR 1989–2008	ERA-I MAR 1989–2008	ERA-40 PDD 1960–1990	ERA-40 PDD 1960–1990	ERA-I MAR 1989–2008	ERA-I MAR 1989–2008

initialisation methods in ice2sea project (Edwards et al., 2014)

How much does it matter?

- Short-term ice sheet prediction like weather not climate
 - ice sheet responds on centennial timescales
 - decadal-century scale response depends strongly on initial state
- **Drift** if no mass balance corrections
 - subtract from predictions
- More important than ever
 - robust decadal-century scale predictions for adaptation

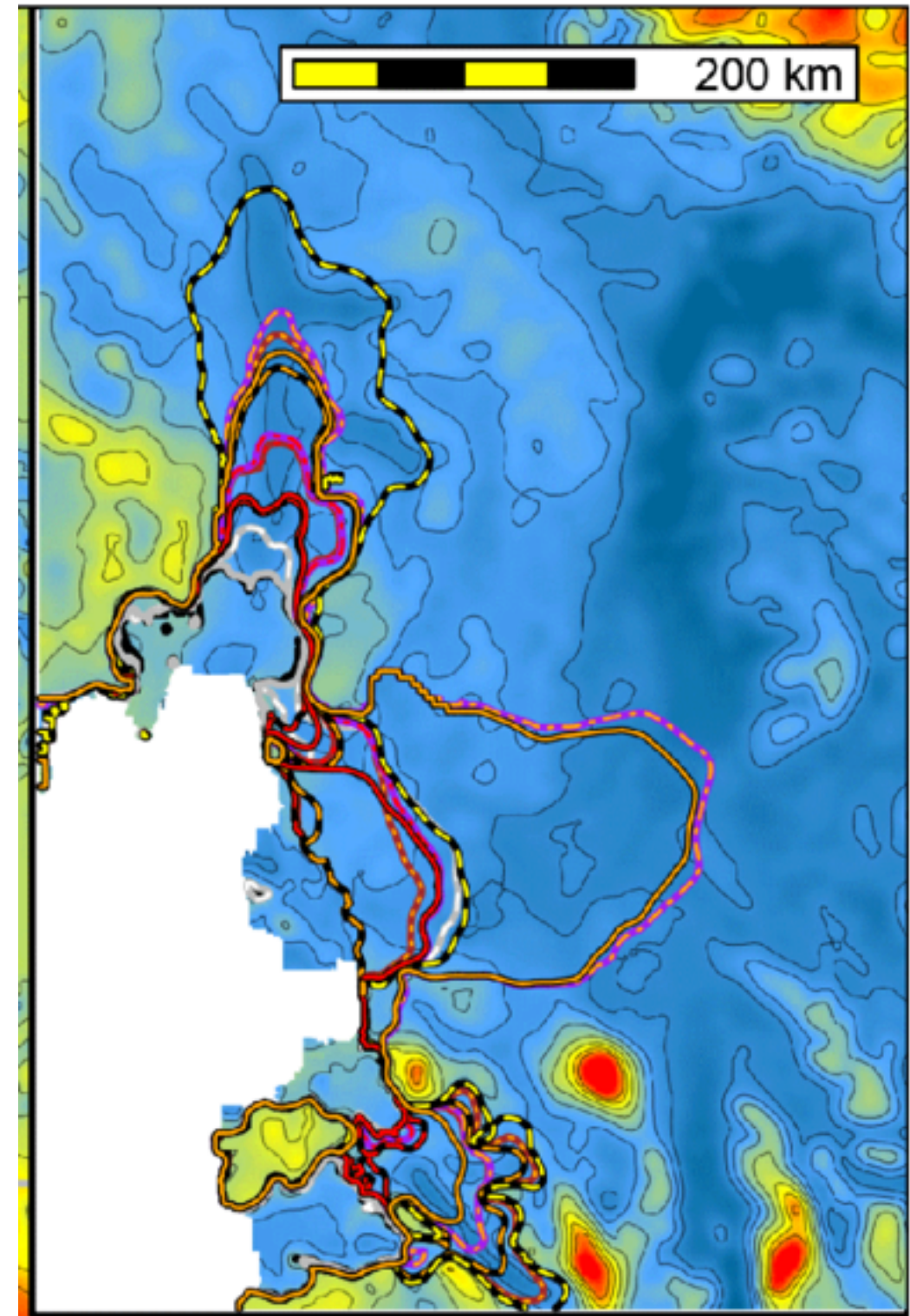


Greenland model drift
(Goelzer et al., 2016, EGU abstract)

How much does it matter?

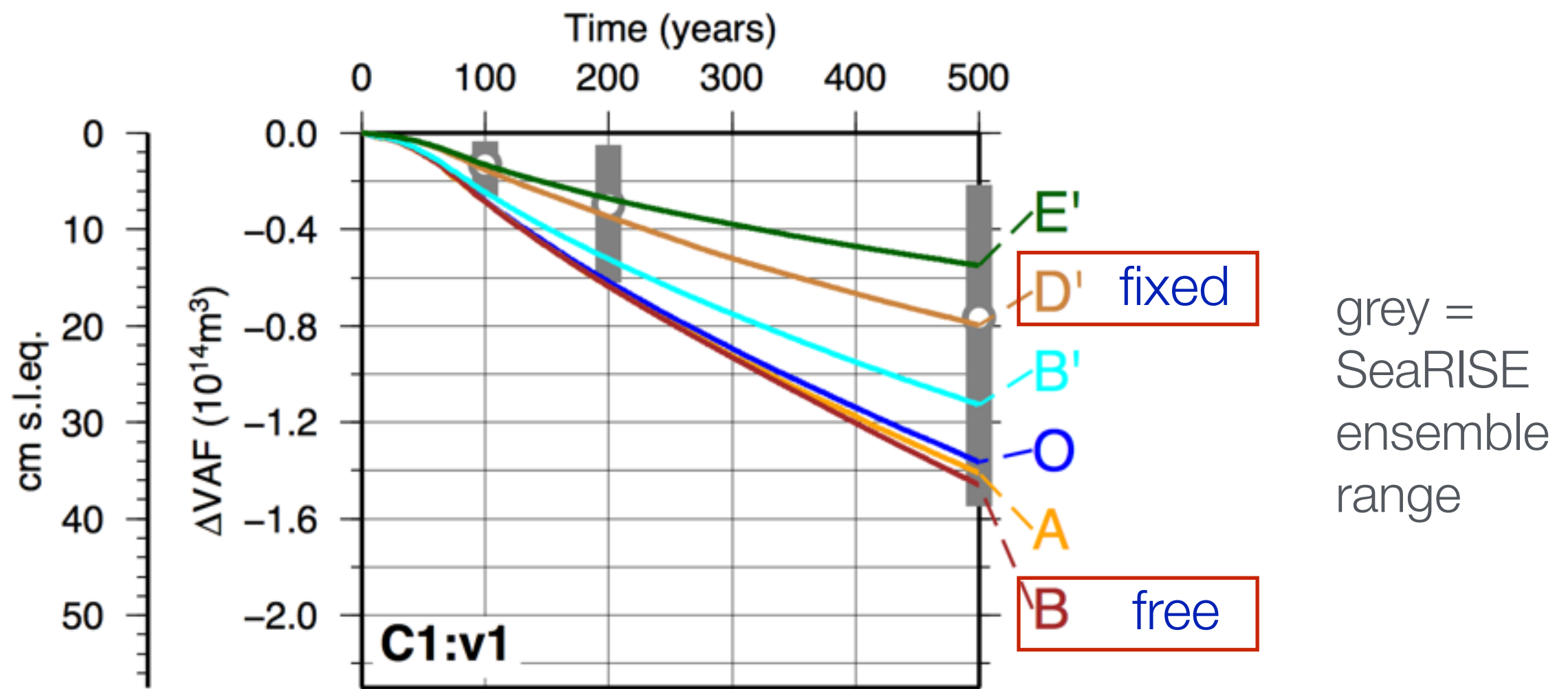
- Initial accumulation from:
 - regional climate model
 - initialisation: mass balance corrections inferred for this climate

“Within the Amundsen Sea Embayment the largest single source of variability is the onset of sustained retreat in Thwaites Glacier, which can triple the rate of eustatic sea level rise....depends strongly on its initial state”



How much does it matter?

- Greenland ice sheet
 - 500 years of A1B scenario
- glacial-interglacial cycle spin-up
 - fix geometry to observed or allow to evolve freely?



grey =
SeaRISE
ensemble
range

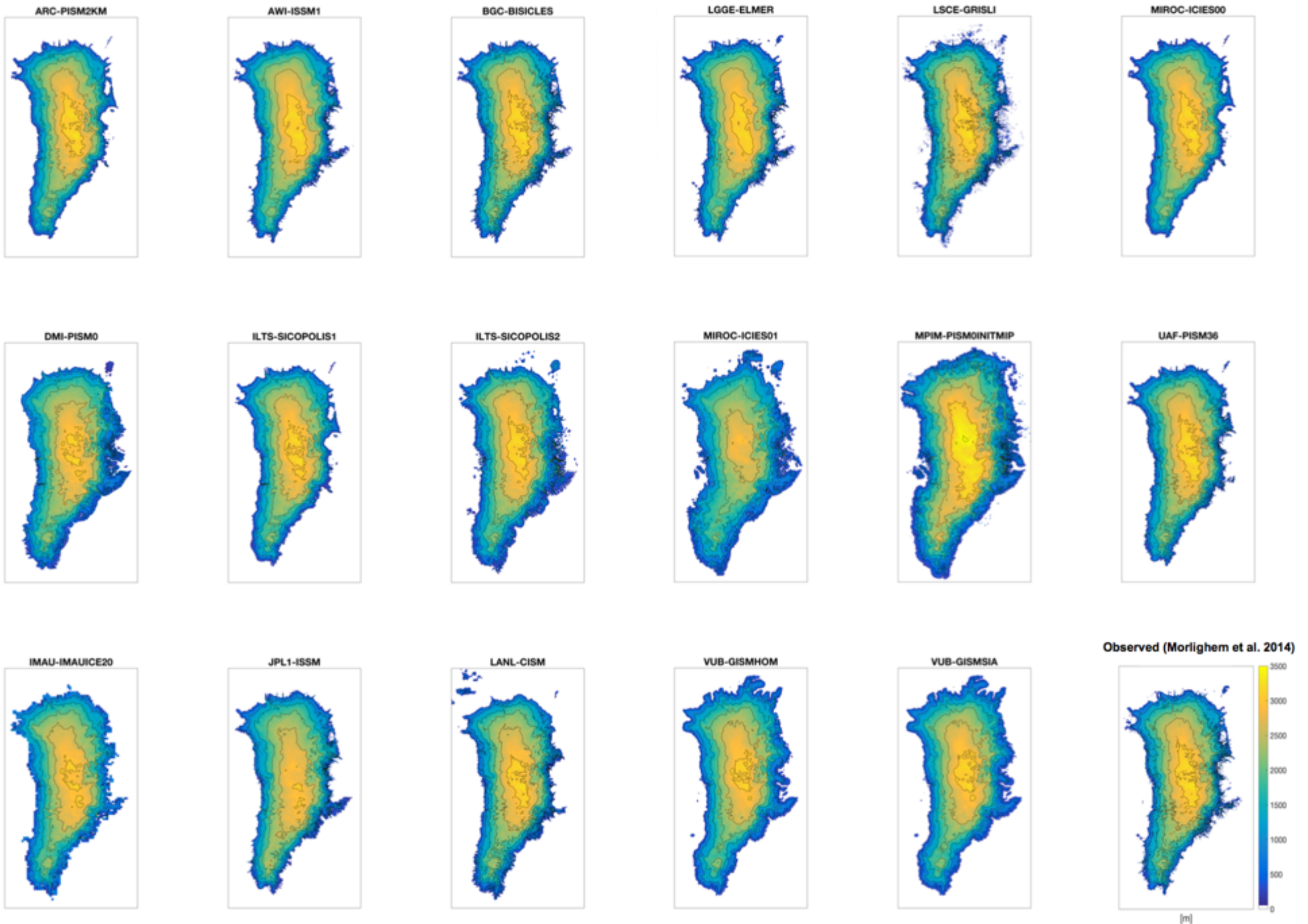
initMIP Goals

- Compare and evaluate the initialisation methods used in the ice sheet modelling community
- Estimate uncertainty associated with initialisation
- Get the ice sheet modelling community started with ISMIP6 activities



- “Requirements”

- Participants can and are encouraged to contribute with different... initialisation methods
- The choice of model input data is unconstrained to allow participants the use of their preferred model setup without modification.
- The specific year of initialization (between 1950 and 2014) is equally unconstrained



initial ice thickness
 (Goelzer et al, 2016, EGU abstract)

evaluating an ice sheet model

Evaluating an ice sheet model

- Important!
 - tests model adequacy
 - can quantify model uncertainty
- Not much formal statistical inference out there
 - only arbitrary comparisons e.g. RMSE
- **Calibrating** models in statistical framework
 - use ensemble of simulations with different input values and fields
 - compare with observations
 - update knowledge about good/bad parameter values
- Ad-hoc methods also used



Model calibration statistical frameworks

- History matching

- rule out poor versions of model to give confidence intervals

- Bayesian calibration

- highest weights to best versions to give probability distributions

What if:
obs = dog
model = cats?



- Strengths and weaknesses

- HM: “this model can’t simulate dogs”
- BC: “here is the cat that looks most like a dog”
“...but here is my uncertainty about that answer”



History matching: Pine Island Glacier

1. model ensemble

5000 simulations
varying 7 parameters

2. observations

grounding line
thickness
velocity

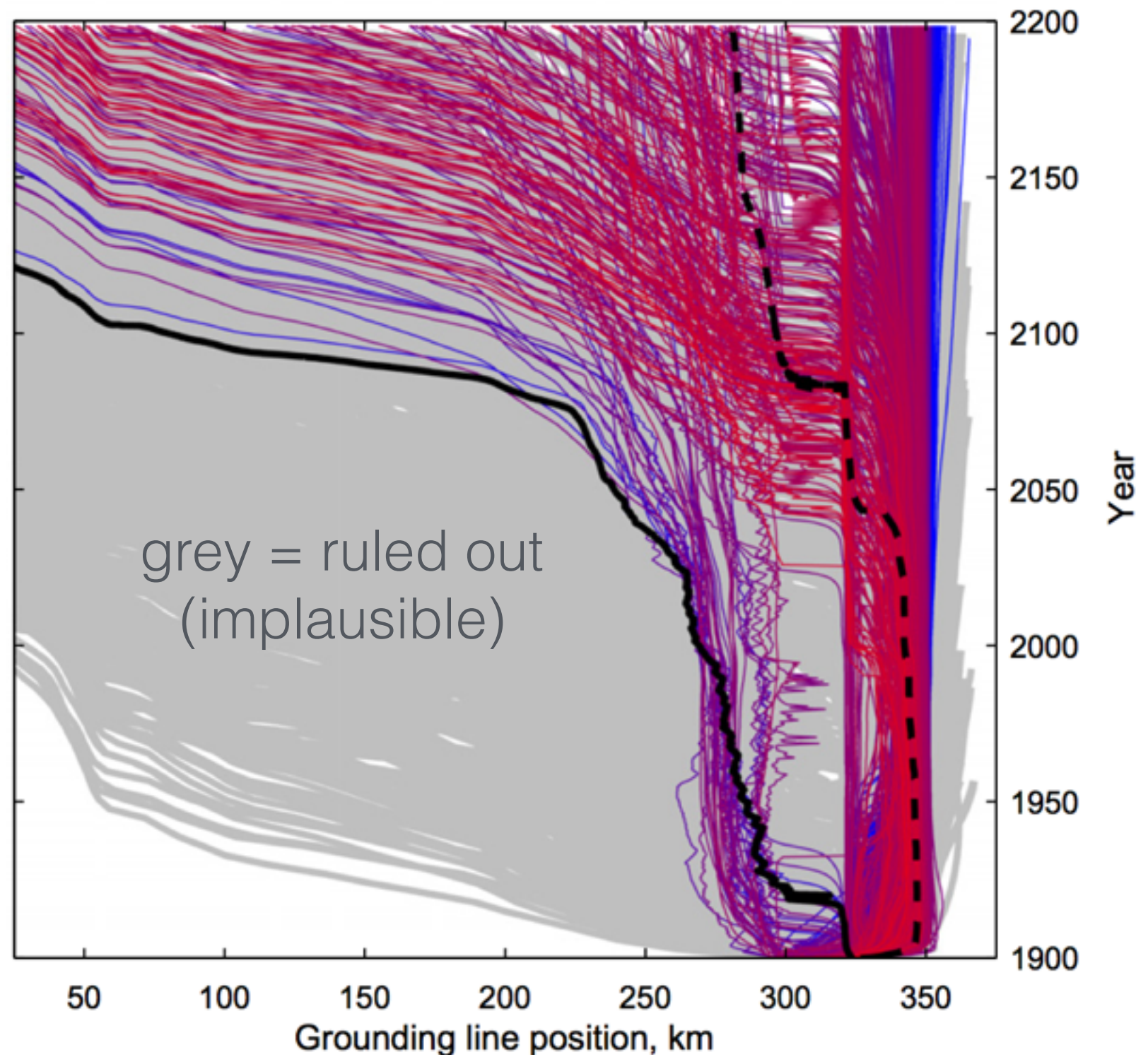
3. choose metric

$$\gamma_t(\phi) = \frac{Z_t - m_t(\phi)}{\sqrt{\text{Var}(\delta_t) + \text{Var}(e_t)}}$$

4. define threshold

$$|\gamma_t(\phi)| \leq 3$$

95% confidence set



Bayesian calibration: Antarctica

1. model ensemble

3000 simulations
varying 16 inputs

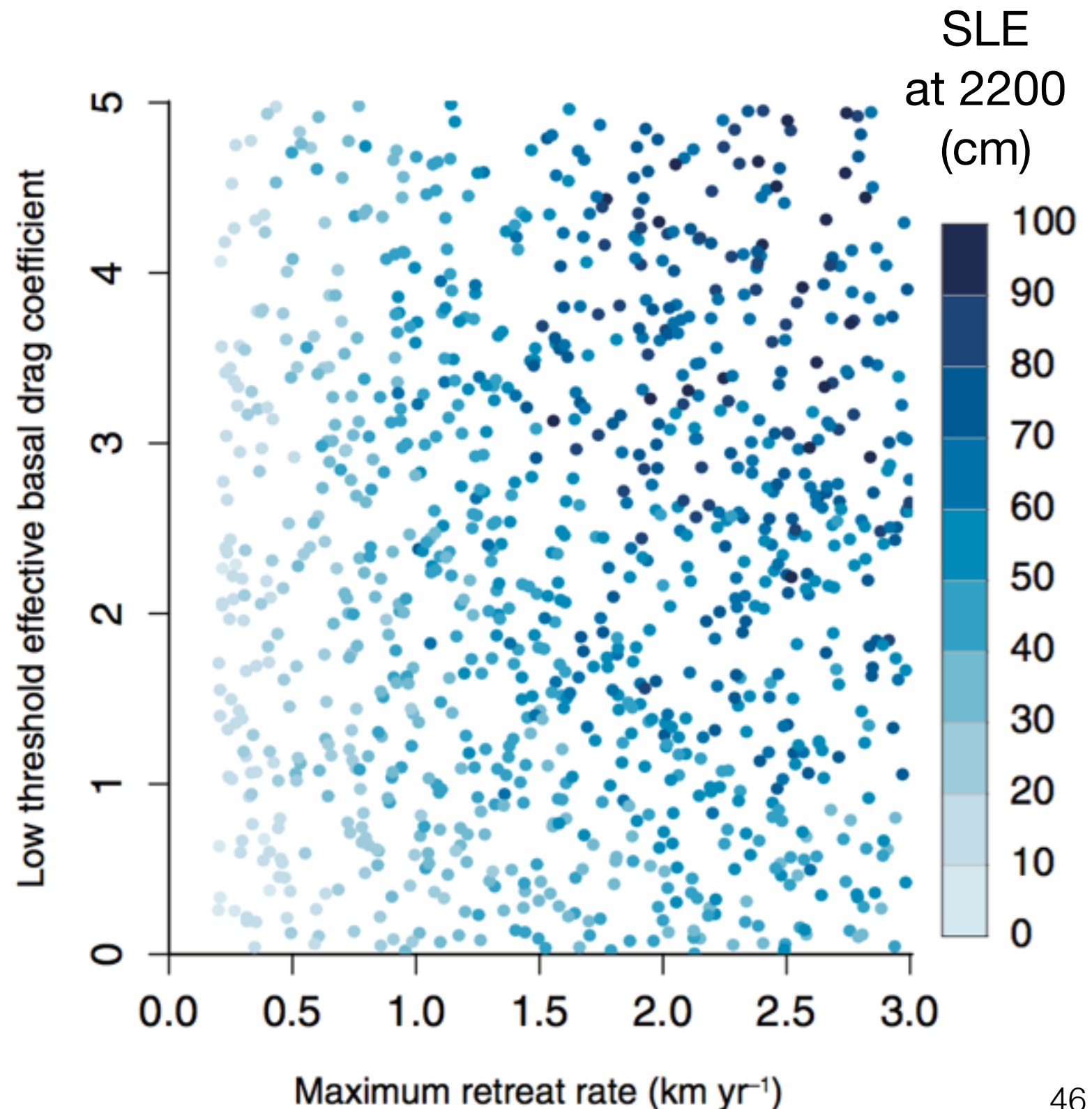
2. observations

Amundsen Sea Embayment
mass trend (IMBIE)

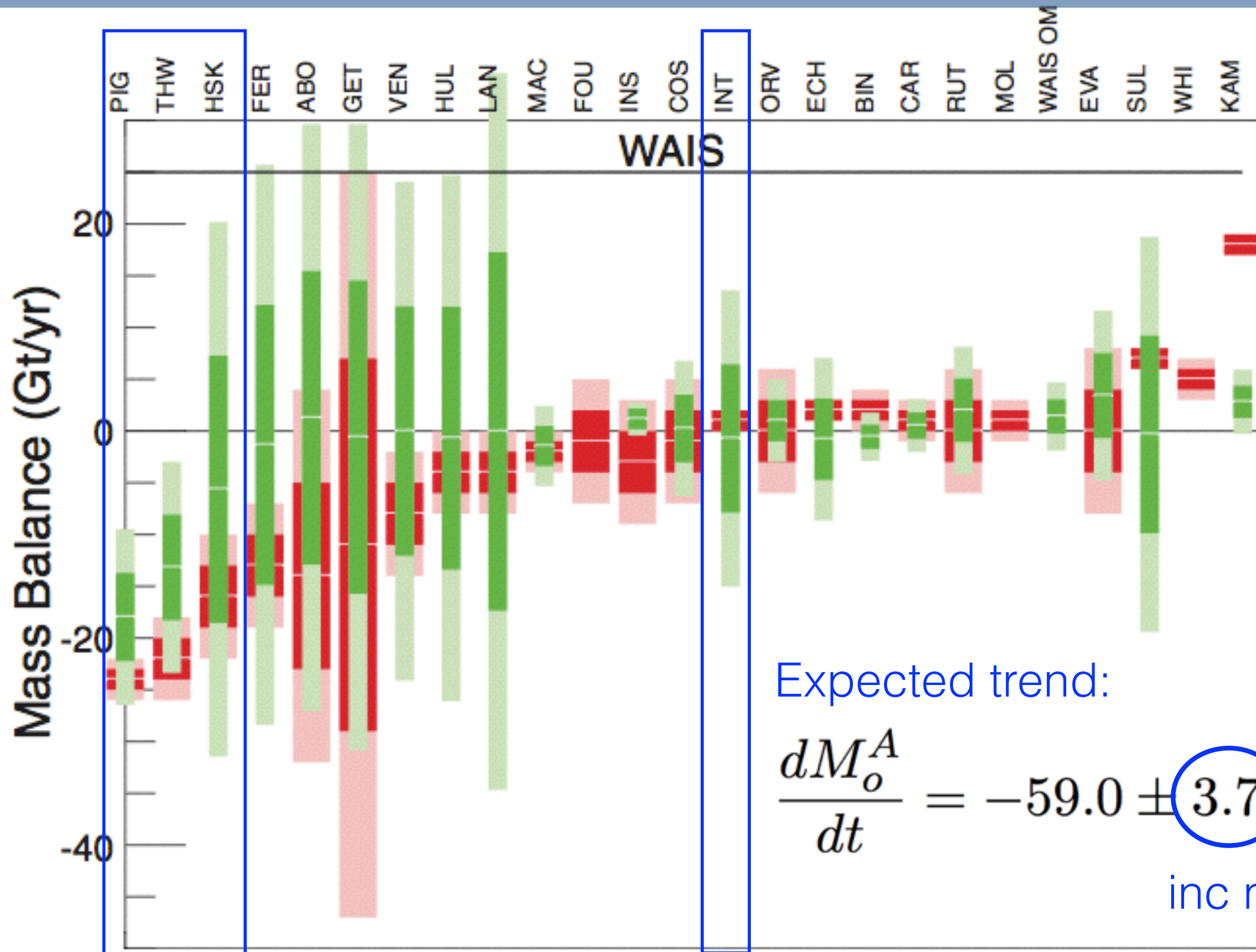
3. choose likelihood

$$w(\phi) \propto \exp \left\{ \frac{-(Z - m(\phi))^2}{2(\text{Var}(\delta) + \text{Var}(e))} \right\}$$

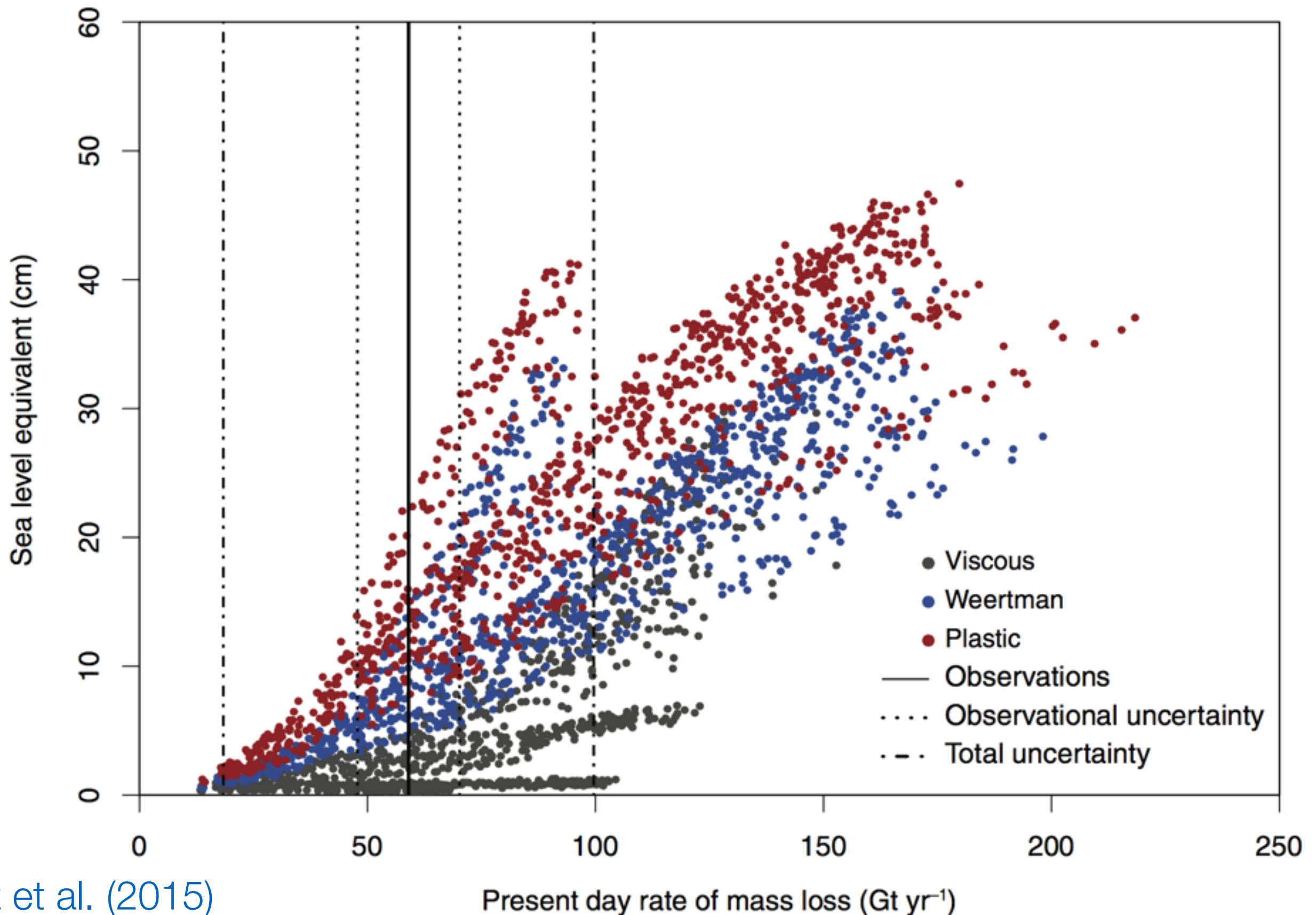
4. normalise and reweight



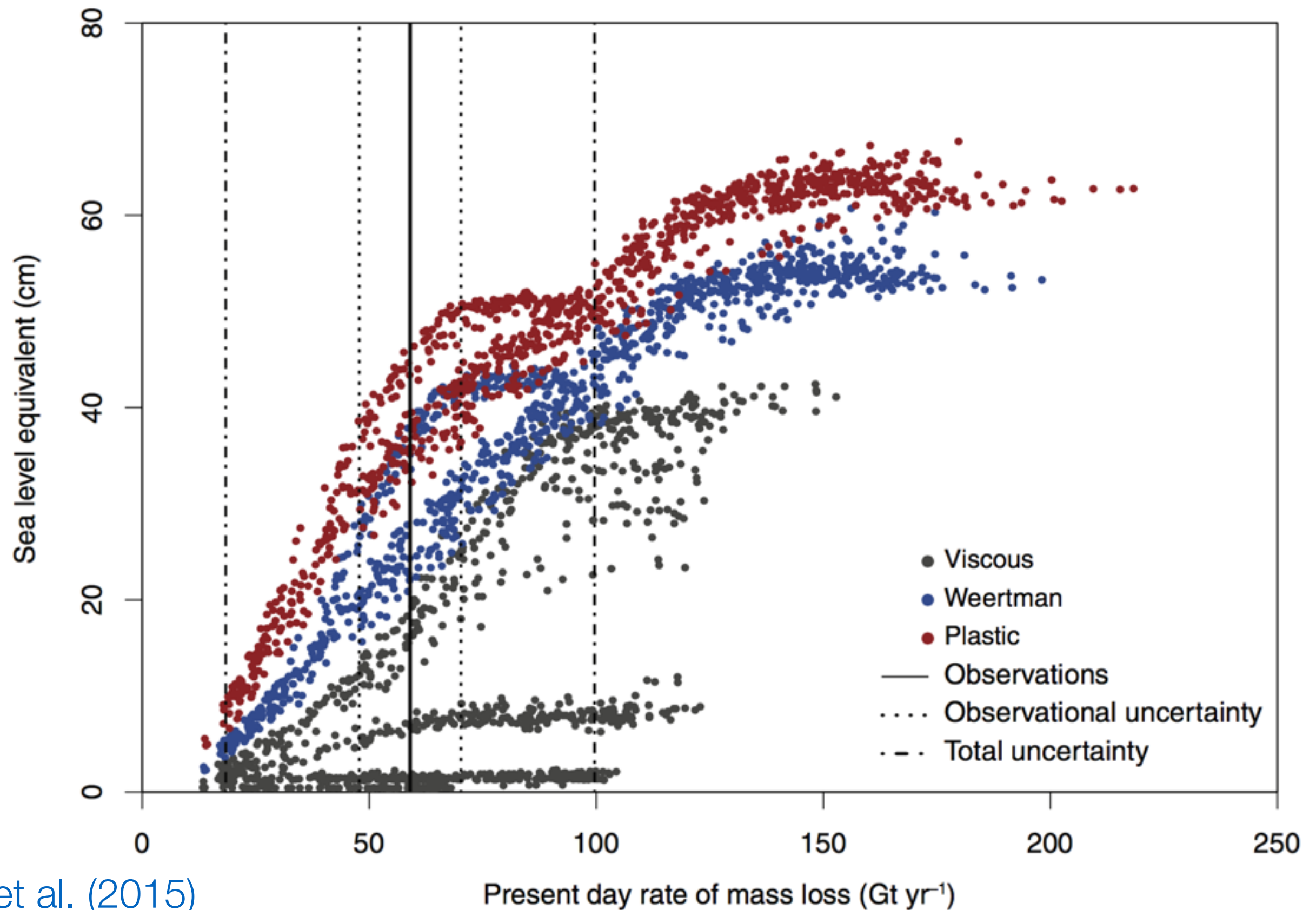
Amundsen Sea Embayment mass trend (1992-2011)



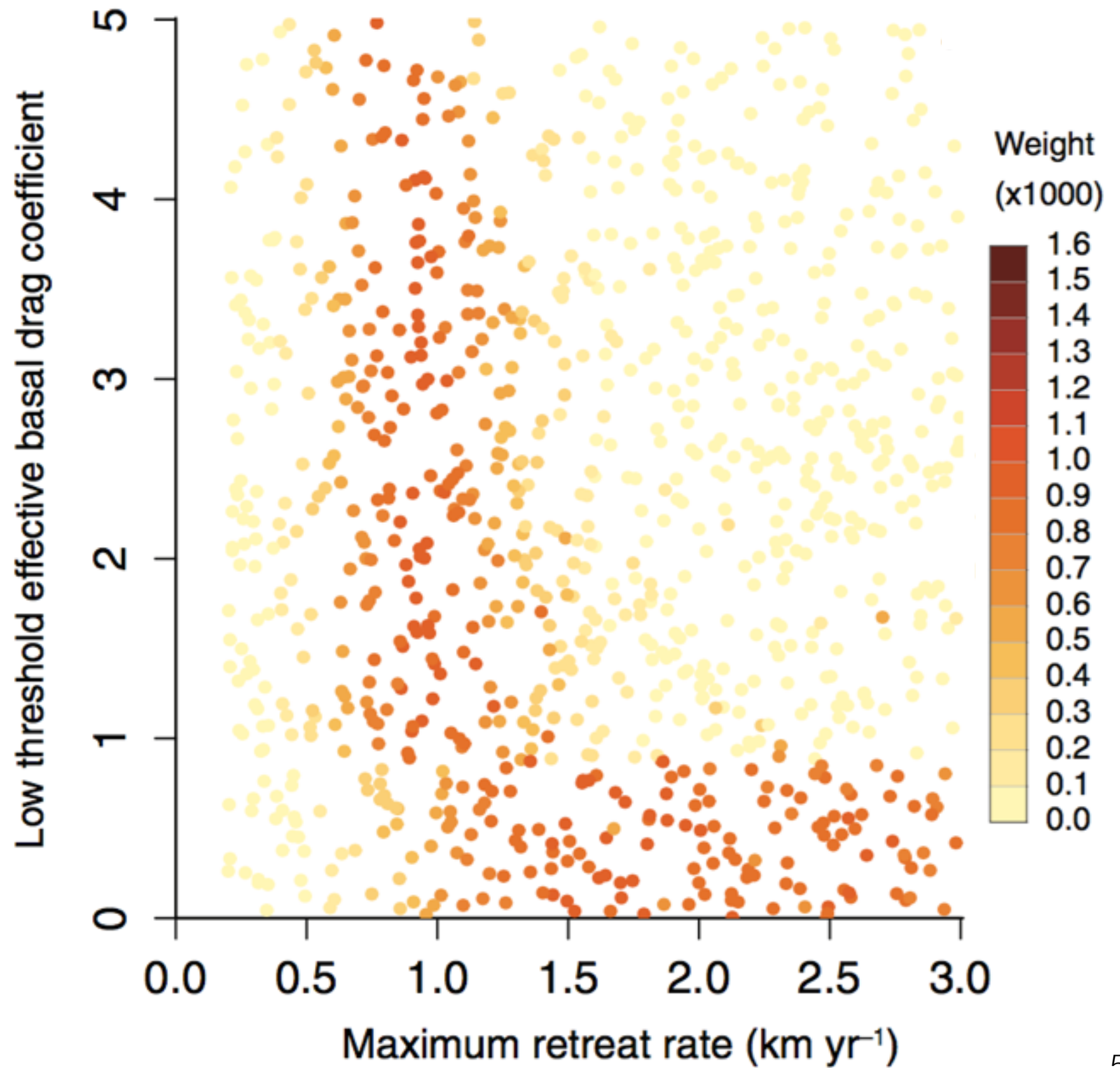
Amundsen Sea Embayment at 2100 vs recent mass trend



Amundsen Sea Embayment at **2200** vs recent mass trend



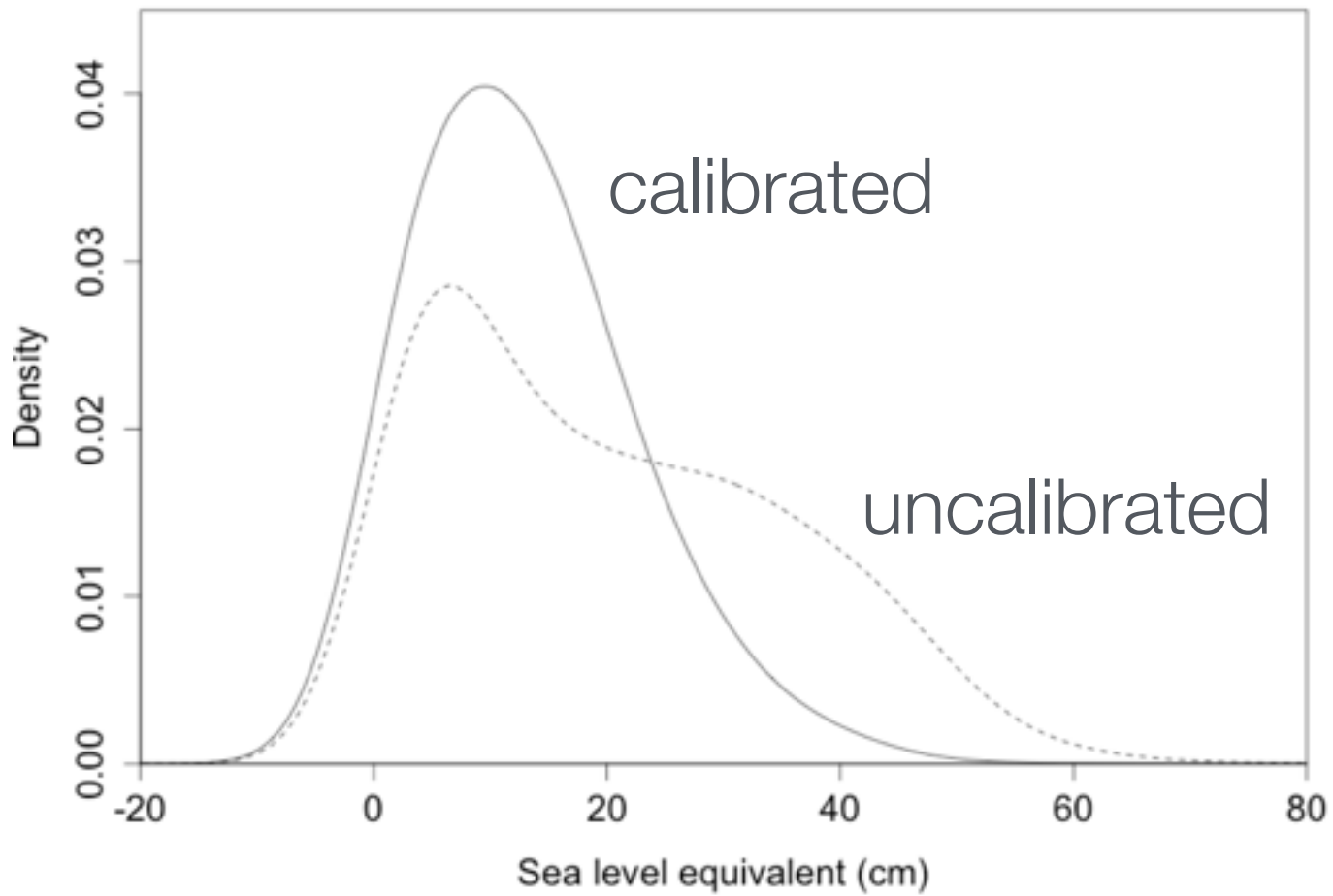
Weights ensemble members



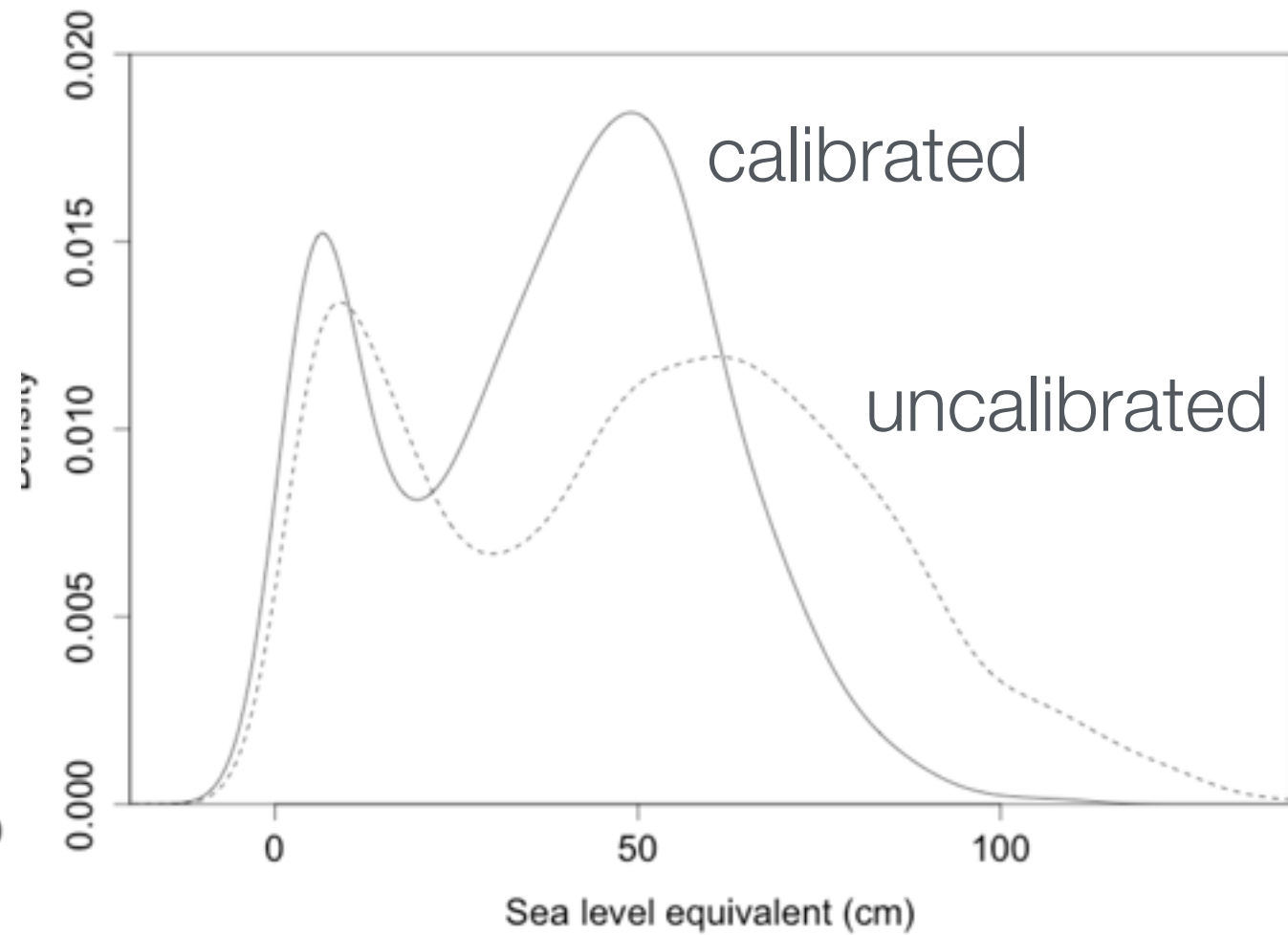
darkest =
highest weight
(highest likelihood)

Effect of calibration on sea level projections

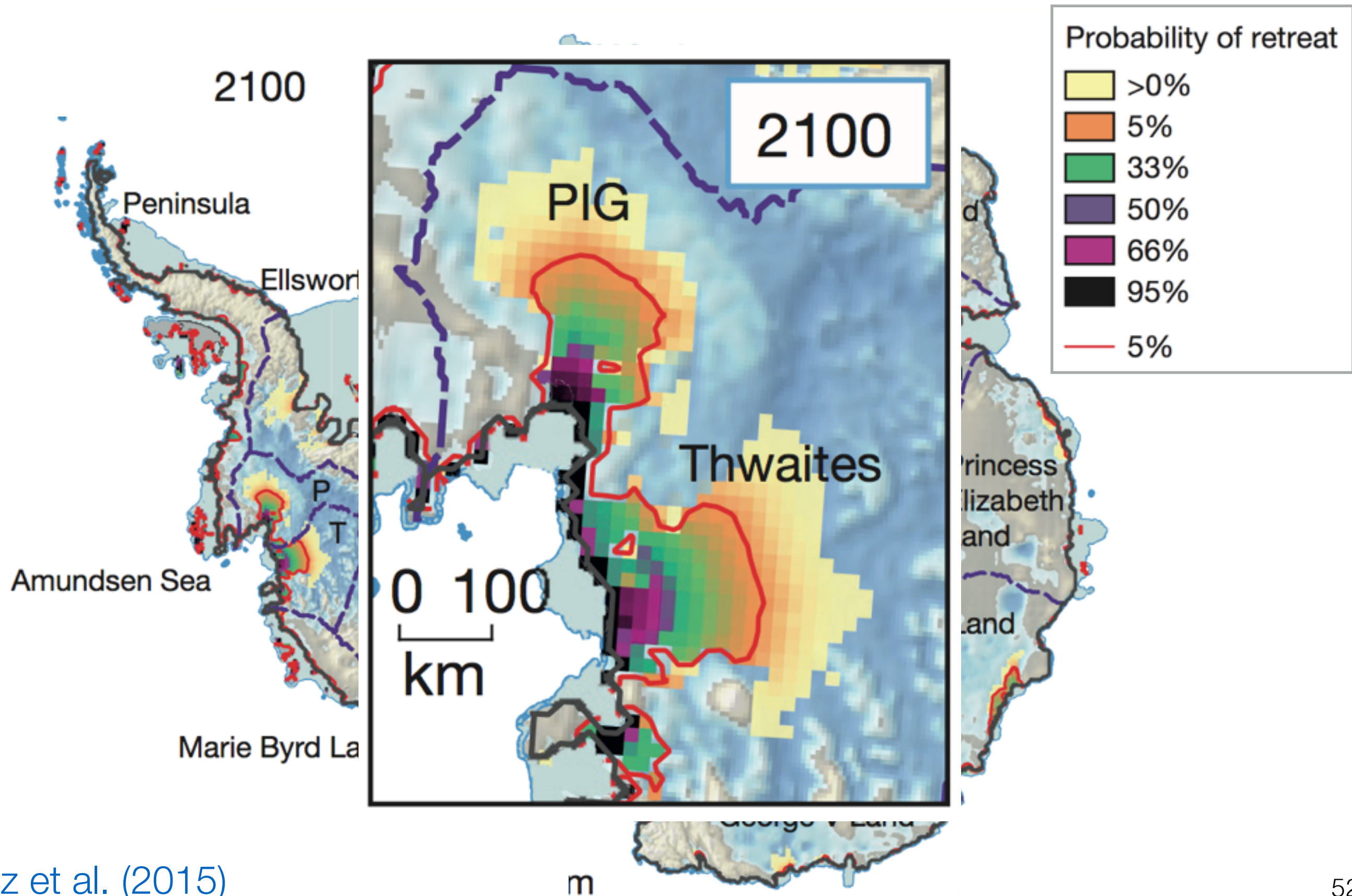
2100



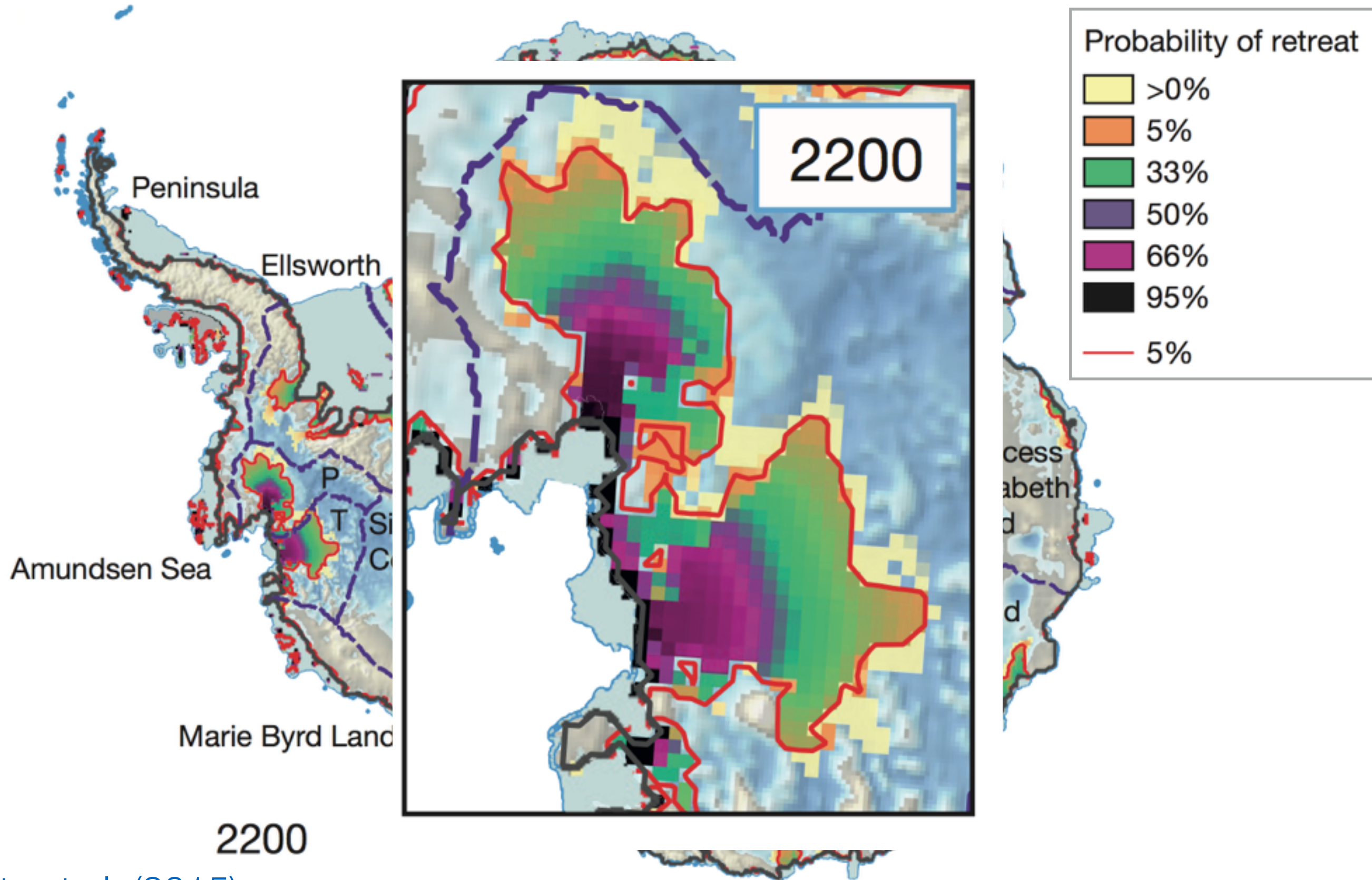
2200

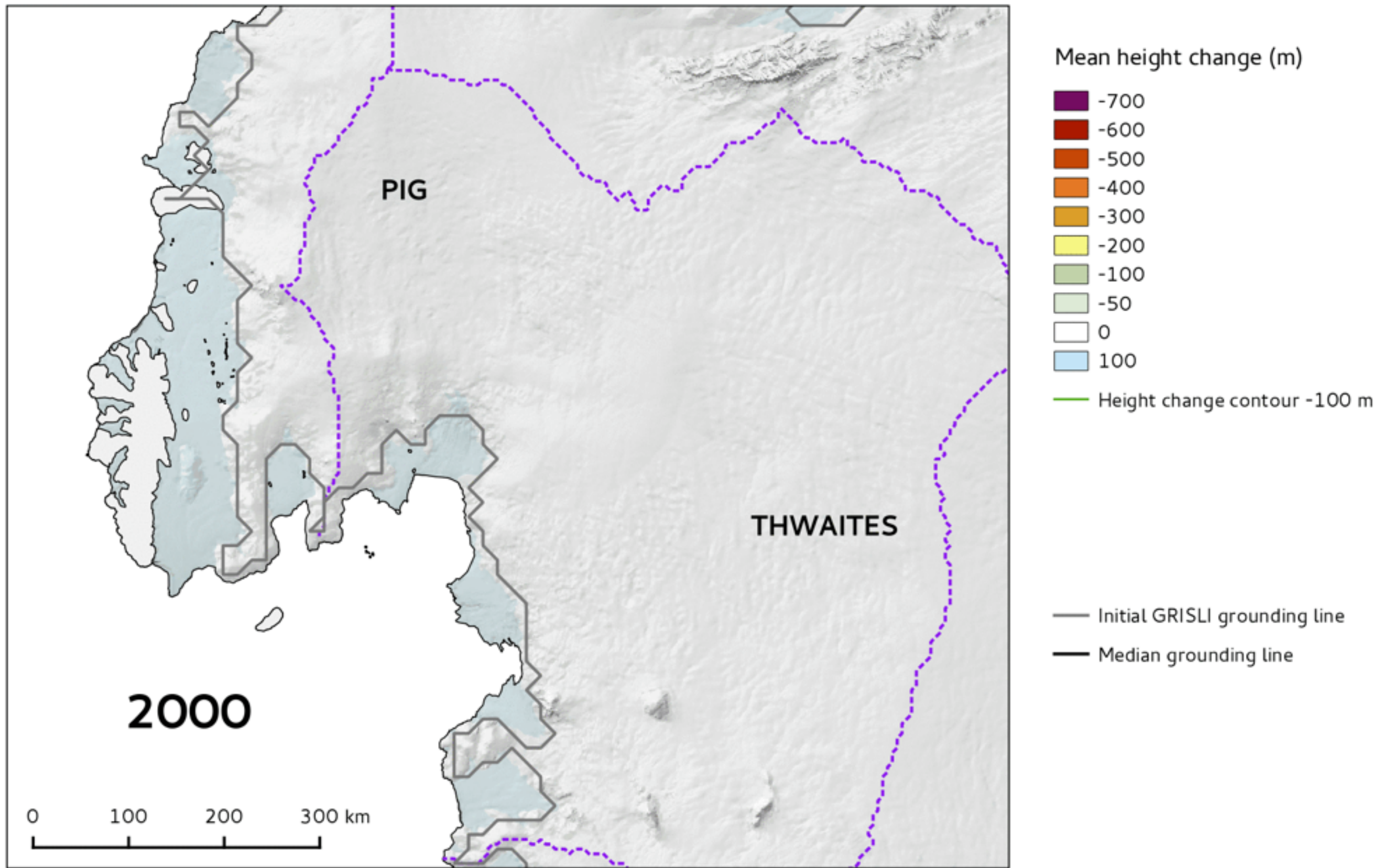


Probability of grounding line retreat at 2100



Probability of grounding line retreat at 2200



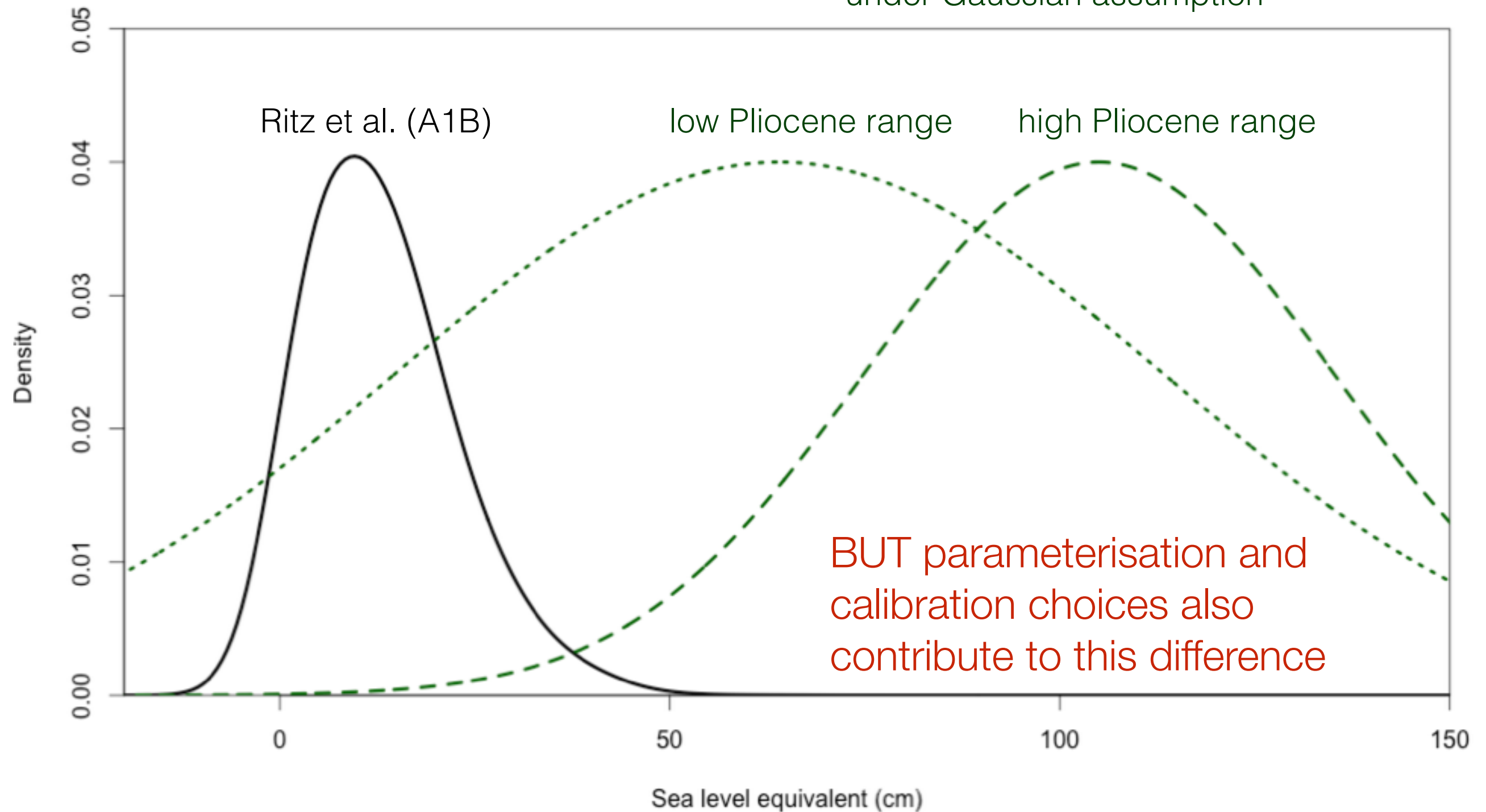


Outlook

- Previous examples: independent model-data comparisons
 - sub-sampled locations (Gladstone et al., 2012)
 - average of region (Ritz et al., 2015)
- Future: use full spatio-temporal information from EO
 - e.g. Won Chang et al.
- Potentially more powerful model calibration
 - But more pitfalls in statistical inference
 - In particular: **correlated uncertainties** in models and observations
- Key question (in my view)
 - maximum rate of Antarctic ice loss
 - does calibration with satellite data bias predictions?

Satellite vs palaeodata bias?

DeConto and Pollard RCP8.5 at 2100
under Gaussian assumption



Summary

- Initialisation of ice sheet models a major uncertainty
 - EO: e.g. geometry, velocity
 - initMIP first semi-systematic step to assessing impact on predictions
 - More to be done here
- Evaluation of ice sheet models is developing
 - EO: e.g. elevation changes, grounding line, mass changes
 - Formal statistical framework gives meaningful inference
 - Moving towards use of EO spatio-temporal patterns
 - Essential to understand correlated uncertainties
 - Antarctica: max rate of ice loss is key uncertainty
- EO will continue to help in reducing & quantifying ice sheet model prediction uncertainties

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