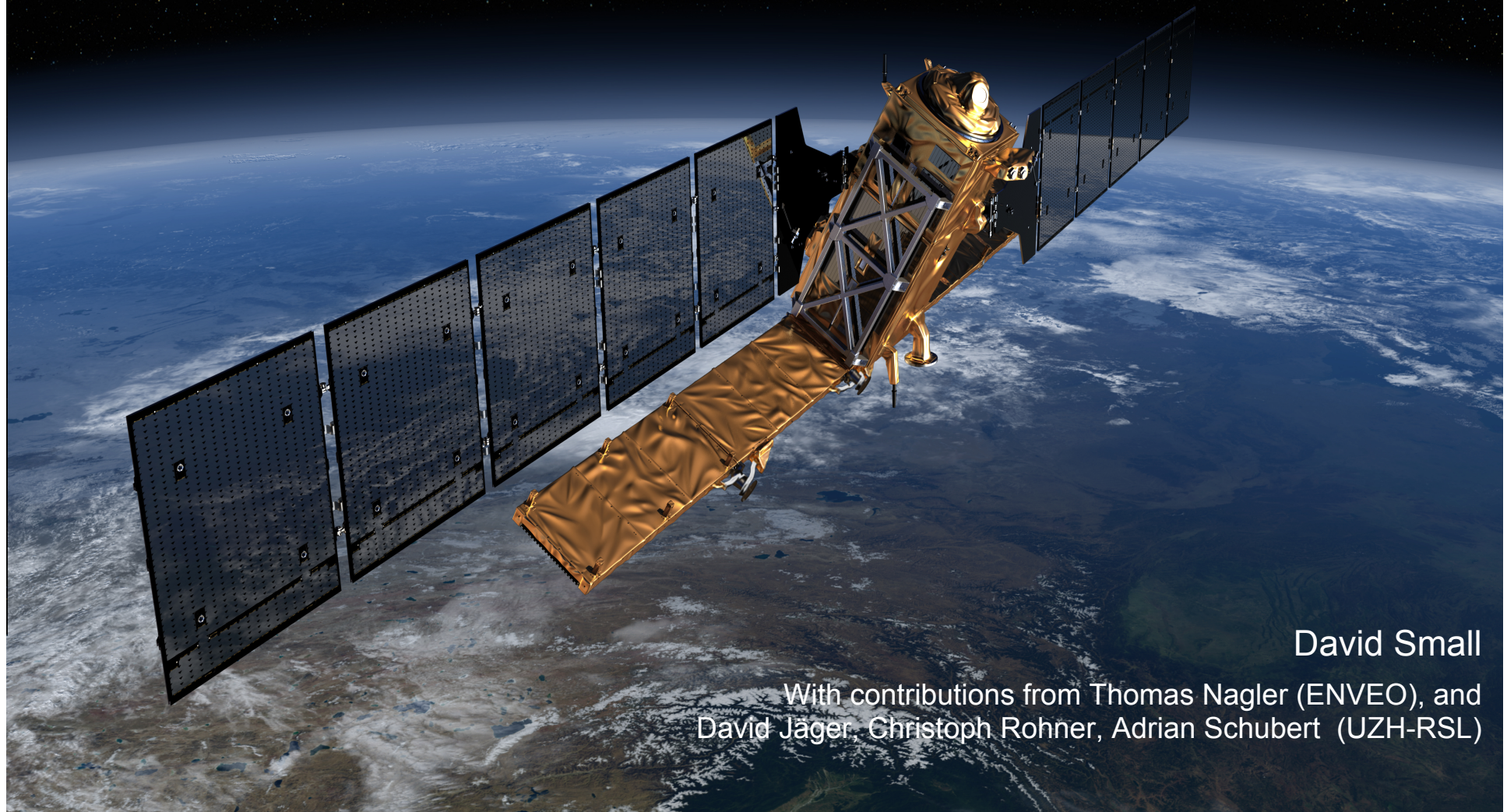




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Zurich <sup>UZH</sup>

ESA Cryosphere Training Course – Sept. 16, 2016  
Leeds, England

## Snow on land from EO



David Small

With contributions from Thomas Nagler (ENVEO), and  
David Jäger, Christoph Rohner, Adrian Schubert (UZH-RSL)



## Outline

### Snow on land from EO

- Motivation
- **Visible & Infrared**
- Microwave radiometry (active): **SAR**
  - Backscatter Normalisation Conventions
  - Radiometric Terrain Corrections
  - Backscatter Compositing
  - Sentinel-1 Examples
- Conclusions





## Remote Sensing of Snow: Motivation

### Knowledge of snow parameterisation important for

- Mitigating large economic impacts of snowfall events
- Snow wetness: run-off modelling: measurements and future prognoses
- Snow wetness: Sudden melt events inducing flooding
- Snow wetness: Hydrology
- Snow distribution and season length – interactions with land cover
- Avalanche modelling
- Climate interactions

J. Dietz, C. Kuenzer, and S. Dech, “**Global SnowPack: a new set of snow cover parameters for studying status and dynamics of the planetary snow cover extent,**” *Remote Sens. Lett.*, 6(11), pp. 844–853, Sep. 2015.

D. R. DeWalle and A. Rango, **Principles of Snow Hydrology**. Cambridge, UK: Cambridge University Press, 2008.



## Multi-sensor approaches

**Multiple sensors** each have own strengths and weaknesses:

- VIS/IR
- Microwave (active & passive)
- Airborne Laser Scanning (ALS)

A. J. Dietz, C. Kuenzer, U. Gessner, and S. Dech, “**Remote sensing of snow – a review of available methods**,” Int. J. Remote Sens., vol. 33, no. 13, pp. 4094–4134, Jul. 2012.

Strengths and weaknesses of respective measurements, e.g.:

- available at night?
- distorted in presence of steep topography?

Difficult in past to integrate e.g. VIS/IR and SAR over large regions due to lack of co-temporal products with similar resolutions / relatively homogenous properties

Future: integrate required ‘harmonised’ measurements, harnessing all strengths?





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# **Visible / Infra-red**



# Composite Products

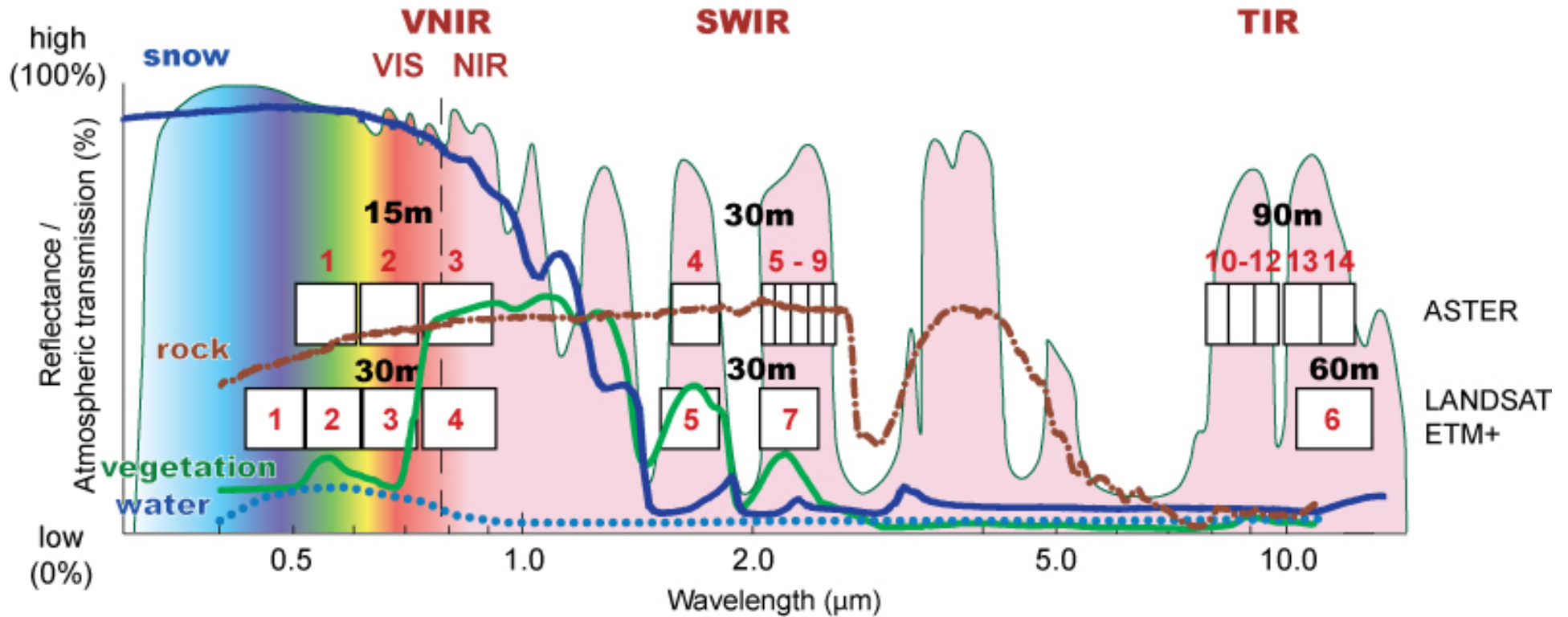
- Assemble wide-area coverage using set window
- Well-established for optical remote sensing
- [https://lpdaac.usgs.gov/dataset\\_discovery](https://lpdaac.usgs.gov/dataset_discovery)
- Composite usage a recognised remedy to cloud coverage issue
  - Clouds are a **spatio-temporal** phenomenon that introduce difficulties to interpreting single optical image sets
- For SAR data, topography introduces difficulties to interpretation of single image acquisition due to:
  - gradients in terrain (**spatial**)
  - variations in imaging geometry depending on track (**temporal**)
- Composite products could help resolve issues in interpreting SAR backscatter
- Yet no widely-established standard SAR-based composite products to date

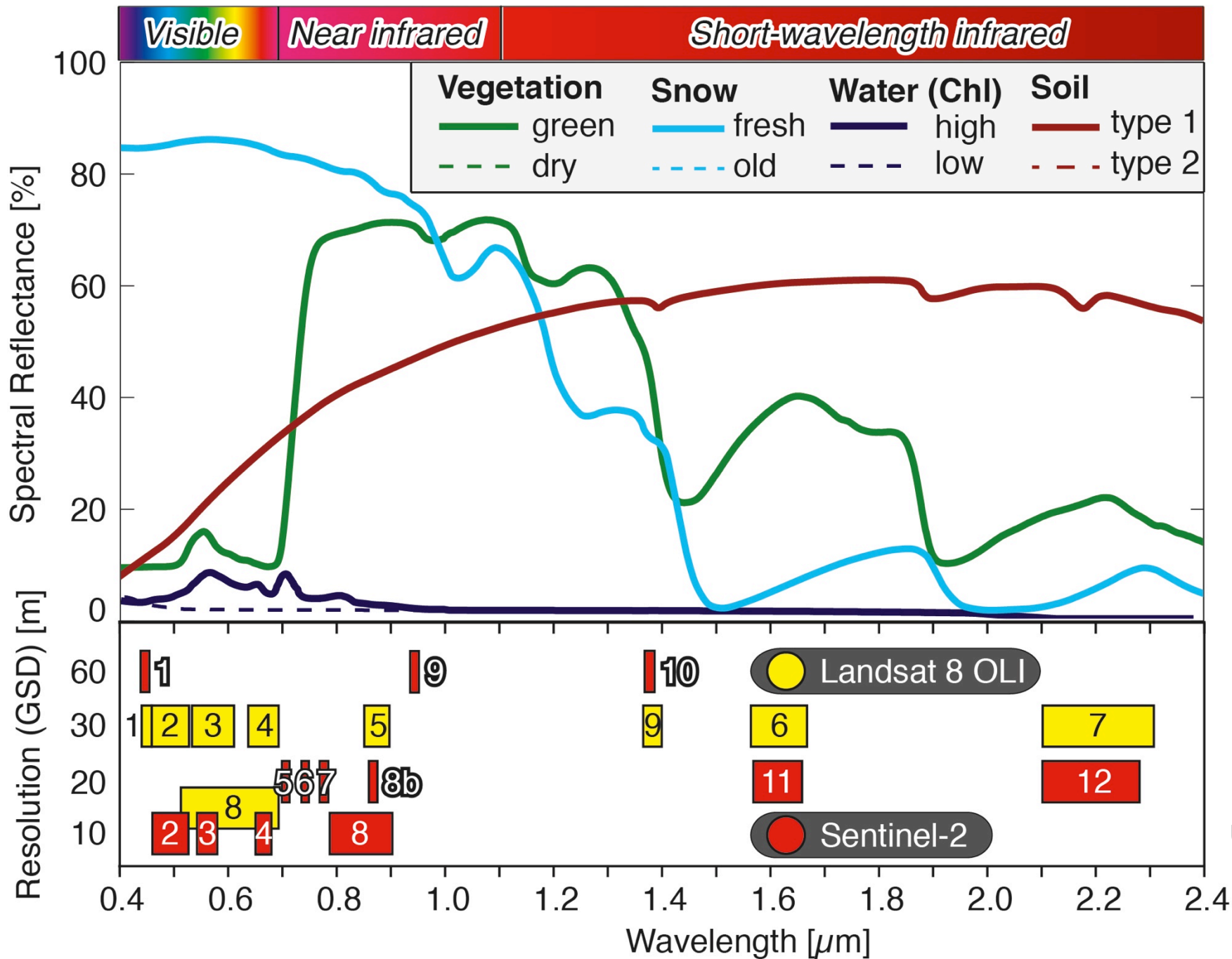
Name	Dataset	Product	Pixel Size	Temporal Granularity
<a href="#">MCD15A2</a>	Combined MODIS	Leaf Area Index and Fractional Photosynthetically Active Radiation	1000	Composites
<a href="#">MCD15A3</a>	Combined MODIS	Leaf Area Index and Fractional Photosynthetically Active Radiation	1000	Composites
<a href="#">MCD43A1</a>	Combined MODIS	Bidirectional Reflectance Distribution Function and Albedo	500	Composites
<a href="#">MCD43A2</a>	Combined MODIS	Bidirectional Reflectance Distribution Function and Albedo	500	Composites
<a href="#">MCD43A3</a>	Combined MODIS	Bidirectional Reflectance Distribution Function and Albedo	500	Composites
<a href="#">MCD43A4</a>	Combined MODIS	Bidirectional Reflectance Distribution Function and Albedo	500	Composites
<a href="#">MCD43B1</a>	Combined MODIS	Bidirectional Reflectance Distribution Function and Albedo	1000	Composites
<a href="#">MCD43B2</a>	Combined MODIS	Bidirectional Reflectance Distribution Function and Albedo	1000	Composites
<a href="#">MCD43B3</a>	Combined MODIS	Bidirectional Reflectance Distribution Function and Albedo	1000	Composites
<a href="#">MCD43B4</a>	Combined MODIS	Bidirectional Reflectance Distribution Function and Albedo	1000	Composites
<a href="#">MCD43C1</a>	Combined MODIS	Bidirectional Reflectance Distribution Function and Albedo	5600	Composites
<a href="#">MCD43C2</a>	Combined MODIS	Bidirectional Reflectance Distribution Function and Albedo	5600	Composites
<a href="#">MCD43C3</a>	Combined MODIS	Bidirectional Reflectance Distribution Function and Albedo	5600	Composites
<a href="#">MCD43C4</a>	Combined MODIS	Bidirectional Reflectance Distribution Function and Albedo	5600	Composites
<a href="#">MOD09A1</a>	Terra MODIS	Reflectance	500	Composites
<a href="#">MOD09Q1</a>	Terra MODIS	Reflectance	250	Composites
<a href="#">MOD11A2</a>	Terra MODIS	Temperature and Emissivity	1000	Composites
<a href="#">MOD11C1</a>	Terra MODIS	Temperature and Emissivity	5600	Composites
<a href="#">MOD13A1</a>	Terra MODIS	Vegetation Indices	500	Composites
<a href="#">MOD13A2</a>	Terra MODIS	Vegetation Indices	1000	Composites
<a href="#">MOD13C1</a>	Terra MODIS	Vegetation Indices	5600	Composites
<a href="#">MOD13Q1</a>	Terra MODIS	Vegetation Indices	250	Composites
<a href="#">MOD14A2</a>	Terra MODIS	Thermal Anomalies and Fire	1000	Composites
<a href="#">MOD15A2</a>	Terra MODIS	Leaf Area Index and Fractional Photosynthetically Active Radiation	1000	Composites
<a href="#">MOD17A2</a>	Terra MODIS	Gross Primary Productivity	1000	Composites
<a href="#">MOD44A</a>	Terra MODIS	Vegetation Continuous Cover/Fields	250	Composites
<a href="#">MYD09A1</a>	Aqua MODIS	Reflectance	500	Composites
<a href="#">MYD09Q1</a>	Aqua MODIS	Reflectance	250	Composites
<a href="#">MYD11A2</a>	Aqua MODIS	Temperature and Emissivity	1000	Composites
<a href="#">MYD11C1</a>	Aqua MODIS	Temperature and Emissivity	5600	Composites
<a href="#">MYD13A1</a>	Aqua MODIS	Vegetation Indices	500	Composites
<a href="#">MYD13A2</a>	Aqua MODIS	Vegetation Indices	1000	Composites
<a href="#">MYD13C1</a>	Aqua MODIS	Vegetation Indices	5600	Composites
<a href="#">MYD13Q1</a>	Aqua MODIS	Vegetation Indices	250	Composites
<a href="#">MYD17A1</a>	Aqua MODIS	Gross Primary Productivity	1000	Composites
<a href="#">MYD15A2</a>	Aqua MODIS	Leaf Area Index and Fractional Photosynthetically Active Radiation	1000	Composites
<a href="#">MYD17A2</a>	Aqua MODIS	Gross Primary Productivity	1000	Composites





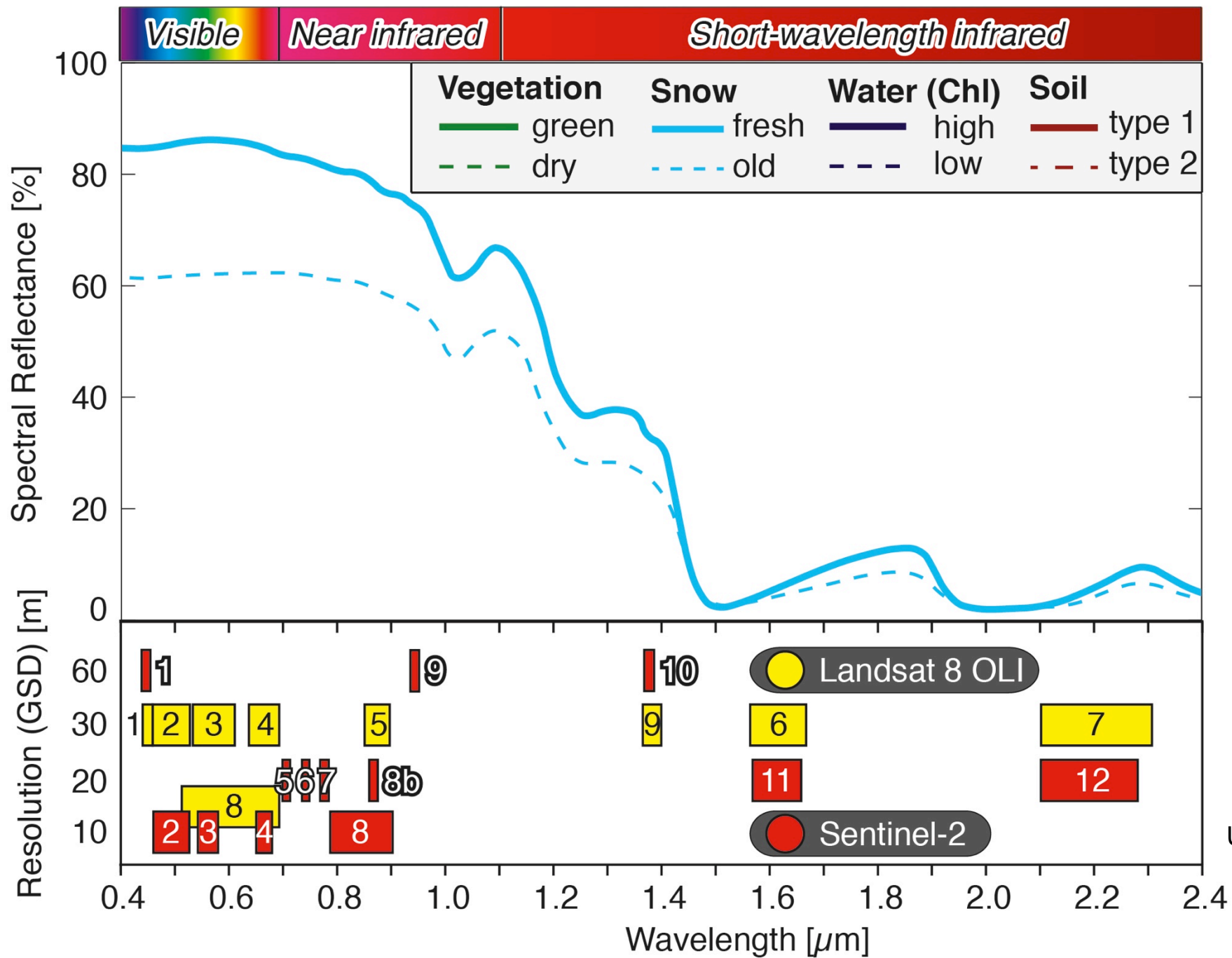
# Remote Sensing of Snow: VIS/IR





Courtesy  
H. Wulf,  
UZH-RSL





Courtesy  
H. Wulf,  
UZH-RSL



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# **Active Microwave: Synthetic Aperture Radar**





## Snow and Dielectric Constant: Attenuation in wet snow

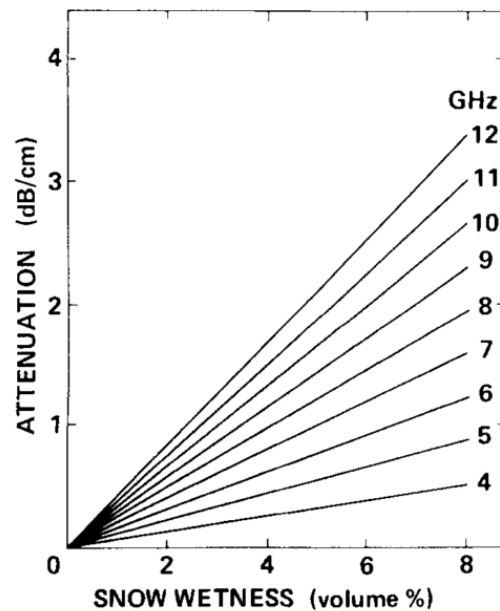
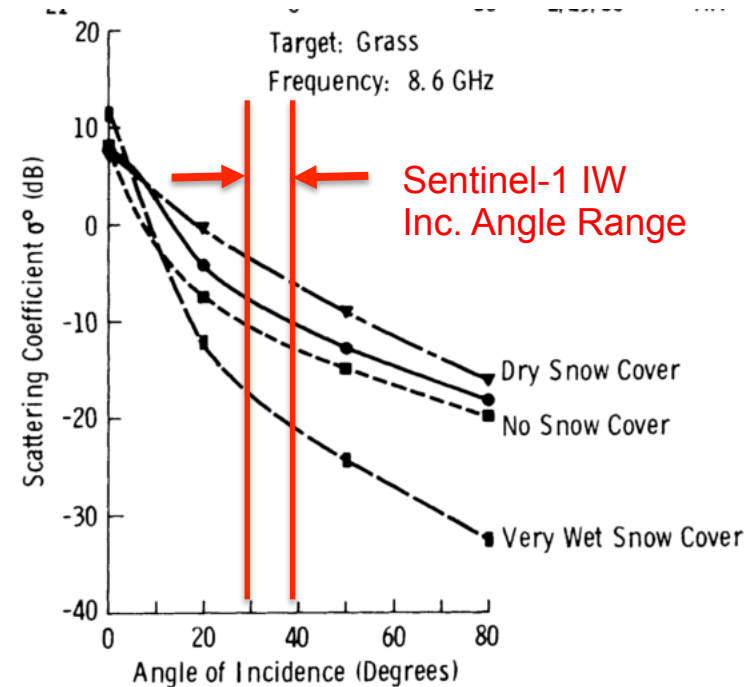


FIG. 5. Variation of attenuation with snow wetness at selected frequencies.



F. T. Ulaby, W. H. Stiles, and M. Abdelrazik, "Snowcover Influence on Backscattering from Terrain," IEEE Trans. Geosci. Remote Sens., vol. GE-22, no. 2, pp. 126–133, Mar. 1984.

W. I. Linlor, "Permittivity and attenuation of wet snow between 4 and 12 GHz," J. Appl. Phys., vol. 51, no. 5, pp. 2811–2816, May 1980.



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# **Active Microwave: Synthetic Aperture Radar**



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# **SAR: Backscatter Normalisation Conventions**





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# Seasonal prioritization of SAR observation windows



Backscatter coefficients [dB] are *ratio of scattered to incident power* over a given area:

<b>RCS</b>	⋮	<b>NRCS</b>	
$\sigma = k \cdot \frac{P_s}{P_i}$		$\beta^0 = \frac{\sigma}{A_\beta}$	$\sigma_E^0 = \frac{\sigma}{A_\sigma} \quad \gamma_E^0 = \frac{\sigma}{A_\gamma}$

Known: transmitted & received power  $P_t$  &  $P_r$

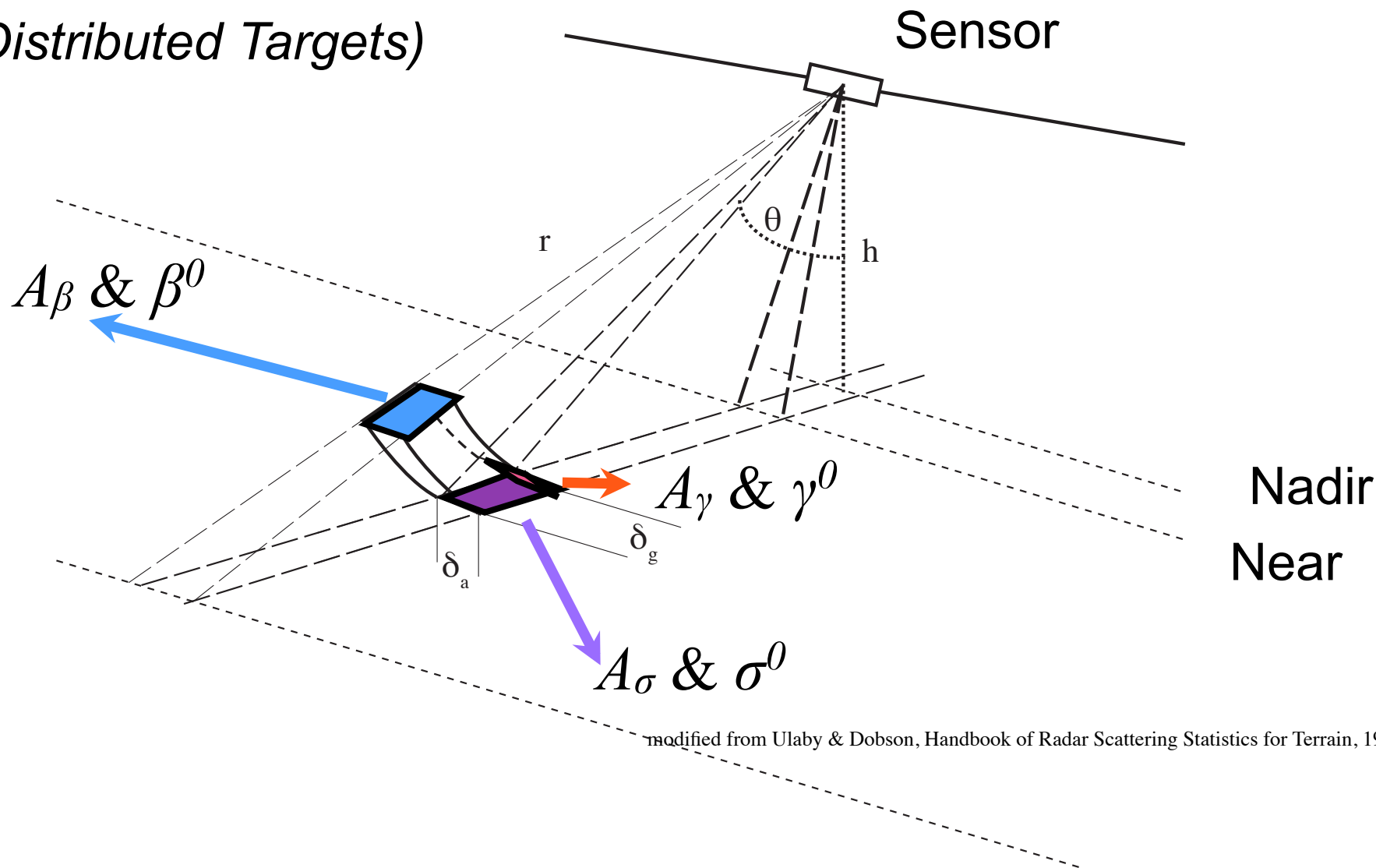
Derive: incident & scattered power  $P_i$  &  $P_s$  from  $P_t$  &  $P_r$

➔  $\beta^0 = k \cdot \frac{f_2(P_r)}{f_1(P_t)} \cdot \frac{1}{A_\beta} \quad \sigma_E^0 = k \cdot \frac{f_2(P_r)}{f_1(P_t)} \cdot \frac{1}{A_\sigma} \quad \gamma_E^0 = k \cdot \frac{f_2(P_r)}{f_1(P_t)} \cdot \frac{1}{A_\gamma}$



# Standard Areas for Normalisation

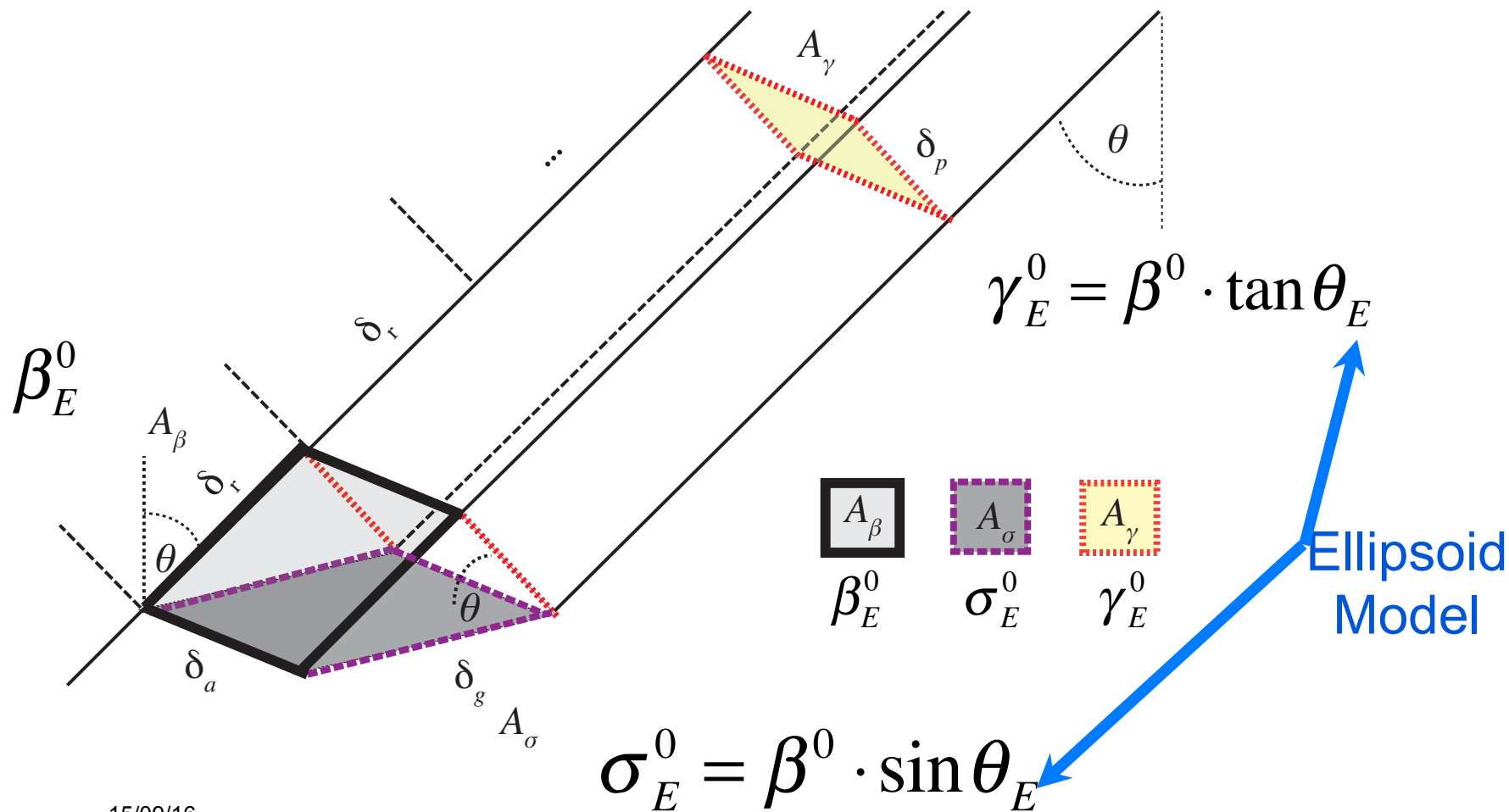
*(Distributed Targets)*



modified from Ulaby & Dobson, Handbook of Radar Scattering Statistics for Terrain, 1989



# Ground Illuminated Area







# Backscatter coefficients are *relative to isotropic scattering*

An idealised **isotropic** scatterer will scatter  
*equally in all directions*

## Real Imaged Objects

- can tend to scatter more **forward** than back to the sensor, focussing energy away from the measurement
  - are darker, generating **negative** *dB* values
- can focus energy **back** towards the sensor (e.g. through corner reflections), generating **positive** *dB* backscatter



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# Wet snow detection with dB thresholding

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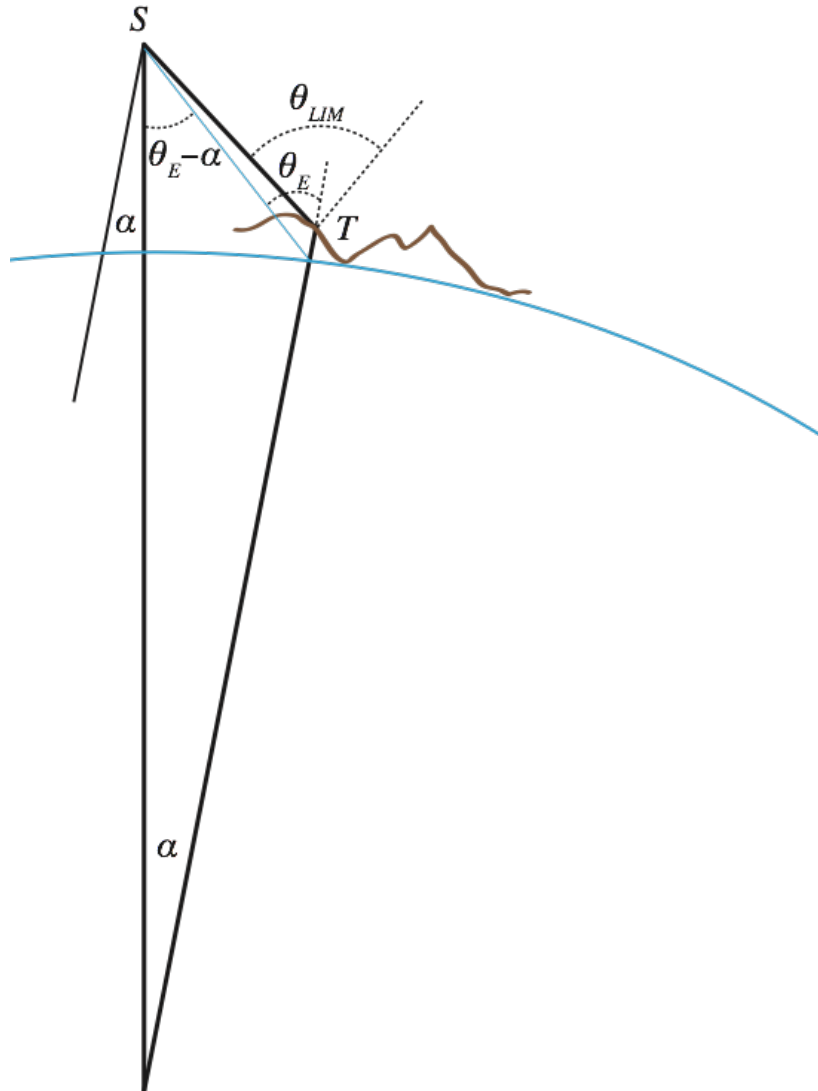
$$(\gamma_{wet}^0 - \gamma_{ref}^0) \text{ [dB]}$$

When difference between candidate image backscatter and dry reference image is lower than -3dB, classify as wet snow

Developed for ERS-1 geometries, VV-pol.

Relies on **exact repeat tracks** (e.g. 35-day ERS repeat) to avoid corruption e.g. by terrain-induced effects

- N. Longépé, S. Allain, L. Ferro-Famil, E. Pottier, and Y. Durand, “**Snowpack Characterization in Mountainous Regions Using C-Band SAR Data and a Meteorological Model**,” IEEE Trans. Geosci. Remote Sens., vol. 47, no. 2, pp. 406–418, Feb. 2009.
- T. Nagler and H. Rott, “**Retrieval of wet snow by means of multitemporal SAR data**,” IEEE Trans. Geosci. Remote Sens., vol. 38, no. 2, pp. 754–765, Mar. 2000.
- N. Baghdadi, Y. Gauthier, and M. Bernier, “**Capability of Multitemporal ERS-1 SAR Data for Wet-Snow Mapping**,” Remote Sens. Environ., vol. 60, no. 2, pp. 174–186, May 1997.



## Incident Angles:

1. Nominal, from **Ellipsoid**:

$$\theta_E$$

2. Local Incident Angle,  
from **height model**:

$$\theta_{LIM}$$



# Normalising $\sigma$ for *terrain*

$$\beta^0, \sigma_E^0, \gamma_E^0$$

are each usable and widely used to normalise the backscatter  $\sigma$ , but one main problem remains:

Each of  $\beta^0$ ,  $\sigma^0$ ,  $\gamma^0$  vary with the local terrain situation (forest on a hill *foreslope* is brighter than forest on *flat* ground, which is brighter than forest on a hill *backslope*)





# Local Incident-angle Mask (LIM)

The most common slope-normalisation methodology found in the literature fails to account for **non-homomorphic** (one to many correspondence) nature of relationship between Earth coordinates (**map geometry**) & **slant range geometry** (native sensor acquisition process)

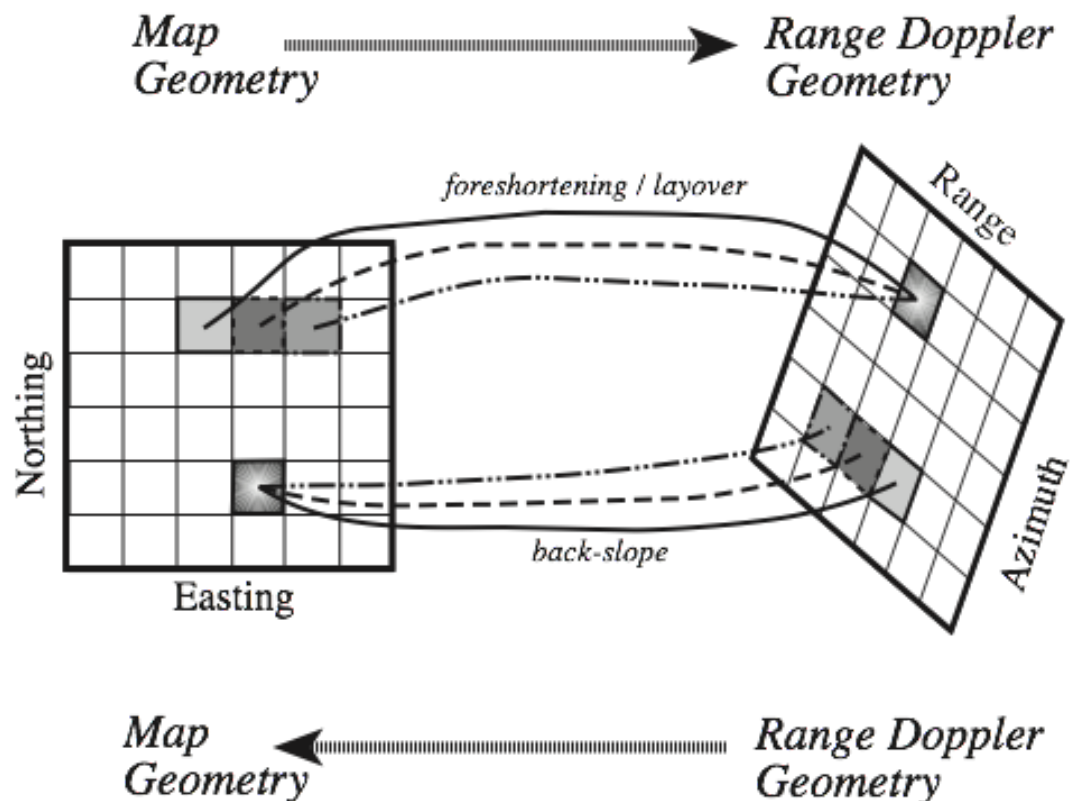
Normalisation for local variation of ground scattering area expressed in **map geometry**:

$$\sigma_T^0 \triangleq \sigma_{NORLIM}^0 = \sigma_E^0 \cdot \frac{\sin \theta_{LIM}}{\sin \theta_E}$$

Kellendorfer et al., TGRS,  
Sept. 1998.



No one-to-one correspondence between slant range and map geometries on **fore-** and **back-**slopes





Relating received to *transmitted* power:

$$\bar{P}_r = \frac{\lambda^2}{(4\pi)^3} \cdot \int_{\text{area illuminated}} \frac{P_t G^2}{R^4} \cdot \sigma^0 dA$$

Ulaby, Moore, Fung,  
1982.

Standard equation of:  $\sigma_E^0 = \beta^0 \cdot \sin \theta_E$

uses an *ellipsoid Earth model* approximation as a standard normalisation area - using ellipsoidal incidence angle  $\theta_E$  as a **proxy for area**

- For radiometric terrain correction, we need to actually *perform the integration* on a DEM



The concept of a **single Local Incident Angle** determining the **terrain's** local normalisation area is **flawed**:

- old concept adapted from **ellipsoidal** incident angle for ocean, sea-ice, & flatlands
- fails to account for:
  - shadow
  - foreshortening
  - layover

Improve sensor model:

- ➔ use local contributing **area**, not angle!
- ➔ and measure that area using the **gamma** convention

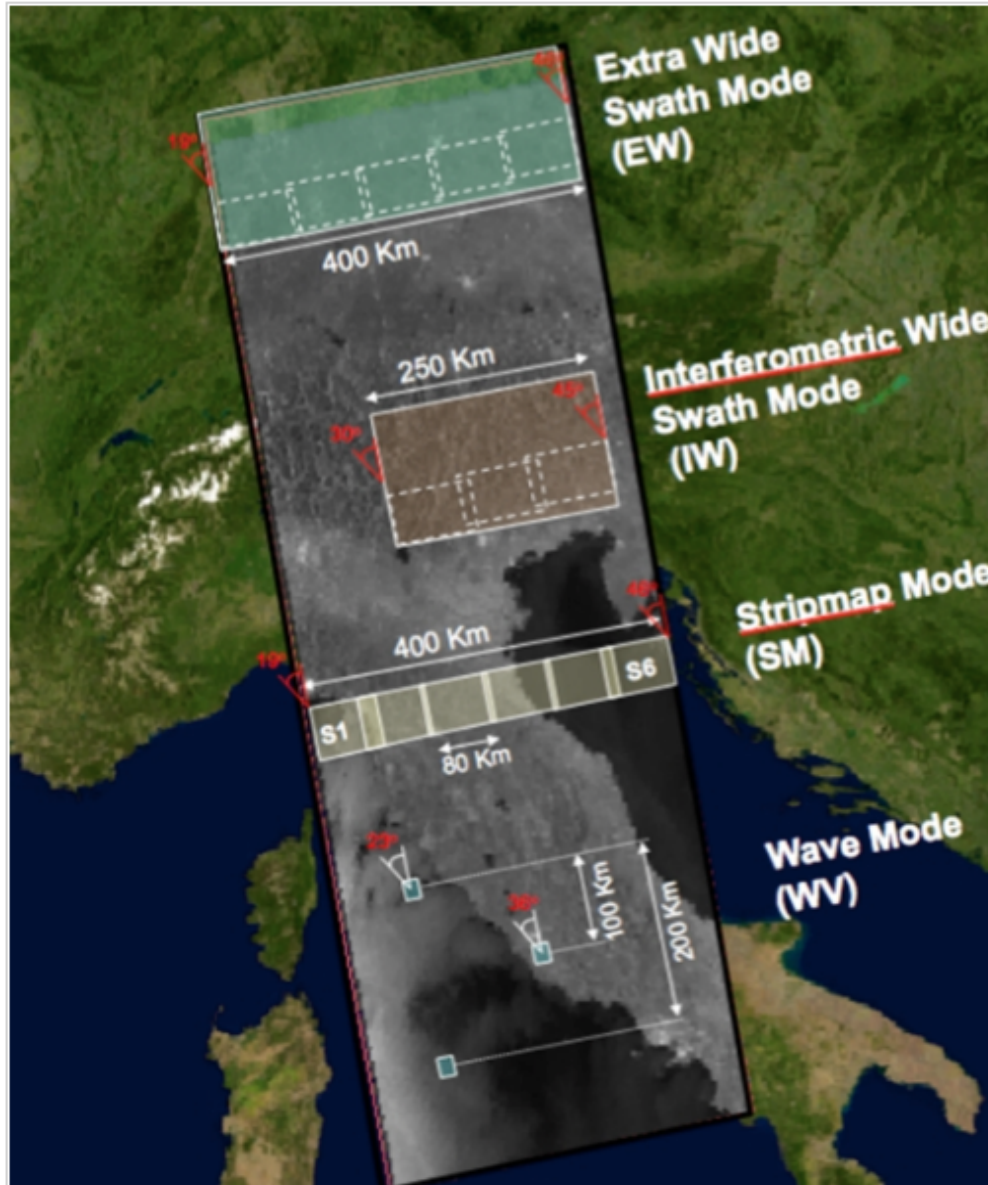
# Radiometric Normalisation Conventions

Convention	1	2	3	4	5
	$\beta^0$	$\sigma_E^0$	$\gamma_E^0$	$\sigma_T^0$	$\gamma_T^0$
Earth Model	<i>None</i>	<i>Ellipsoid</i>		<i>Terrain</i>	
Reference Area	$A_\beta$	$\underline{A}_\sigma$	$\underline{A}_\gamma$	$\hat{A}_\sigma$	$A_\gamma$
Area Derivation	$\delta_r \cdot \delta_a$	$\underline{\delta}_g \cdot \delta_a$	$\underline{\delta}_p \cdot \delta_a$	$\delta_g \cdot \delta_a$	$\int_{DHM} \delta_p \cdot \delta_a$
Normalisation	$\beta^0 = \frac{\sigma}{A_\beta}$	$\beta^0 \cdot \frac{A_\beta}{\underline{A}_\sigma}$ $= \beta^0 \cdot \sin \theta_E$	$\beta^0 \cdot \frac{A_\beta}{\underline{A}_\gamma}$ $= \beta^0 \cdot \tan \theta_E$	$\sigma_E^0 \cdot \frac{\hat{A}_\sigma}{A_\beta}$ $= \sigma_E^0 \cdot \frac{\sin \theta_{LIM}}{\sin \theta_E}$	$\frac{\beta^0 \cdot A_\beta}{A_\gamma}$
Product		<i>GTC</i>		<i>NORLIM</i>	<i>RTC</i>



# Sentinel-1 Acquisition Modes

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← IW is main acquisition mode over land (>80%)

<https://sentinel.esa.int/web/sentinel/user-guides/sentinel-1-sar/revisit-and-coverage>



# Terrain-flattened Gamma Nought

Dept. of Geography / Remote Sensing Laboratories

Interlaken, Switzerland

Sentinel-1A IW GRDH VH-pol.

May 26, 2015

Terrain-flattening: Small D. *Flattening Gamma: Radiometric Terrain Correction for SAR Imagery*, IEEE Trans. on Geoscience & Remote Sensing, 49(8), Aug. 2011, pp. 3081-3093.

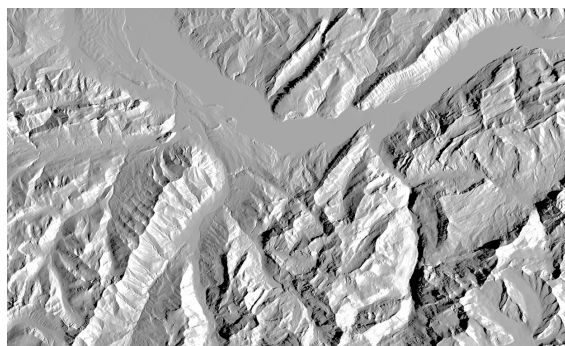
Normalise  $\beta^0$ : divide by simulated image



$\beta^0$

GTC

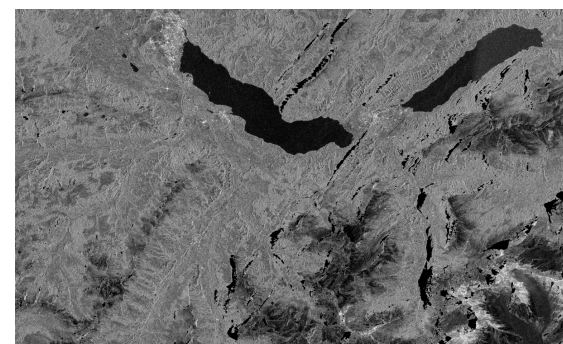
-26dB -1dB



$A_\gamma/A_\beta$

$\gamma_T^0$

=



RTC

$$\gamma_T^0 = \beta^0 \cdot \frac{A_\beta}{A_\gamma}$$






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## Sentinel-1A: **GTC** (Geometrically Terrain Corrected)

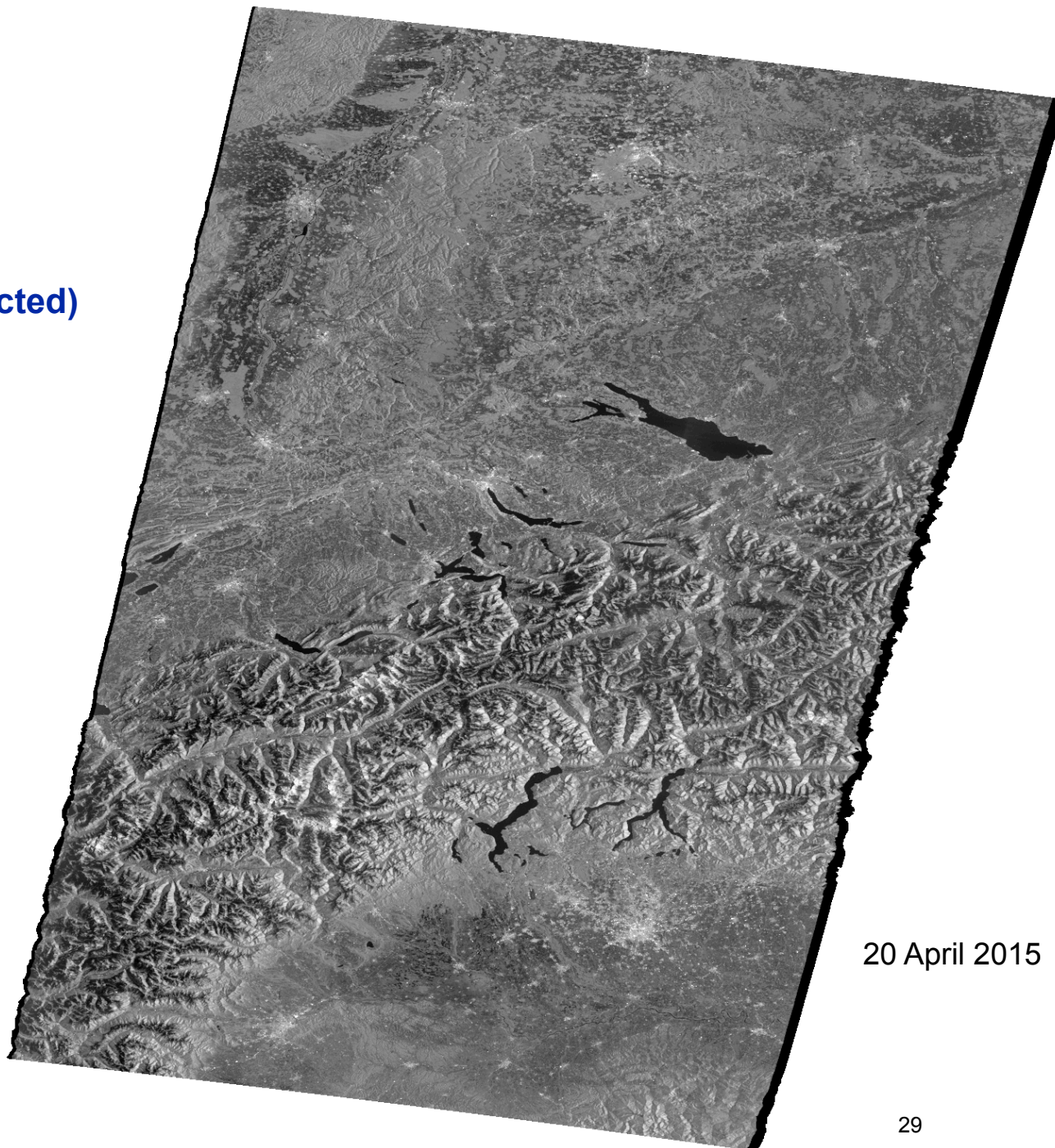
$\gamma_E^0$

-26dB    -1dB



Generated automatically from  
3 IW GRDH products using  
SRTM3

Copernicus Sentinel data (2015)



20 April 2015






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## Sentinel-1A: RTC (Radiometrically Terrain Corrected)

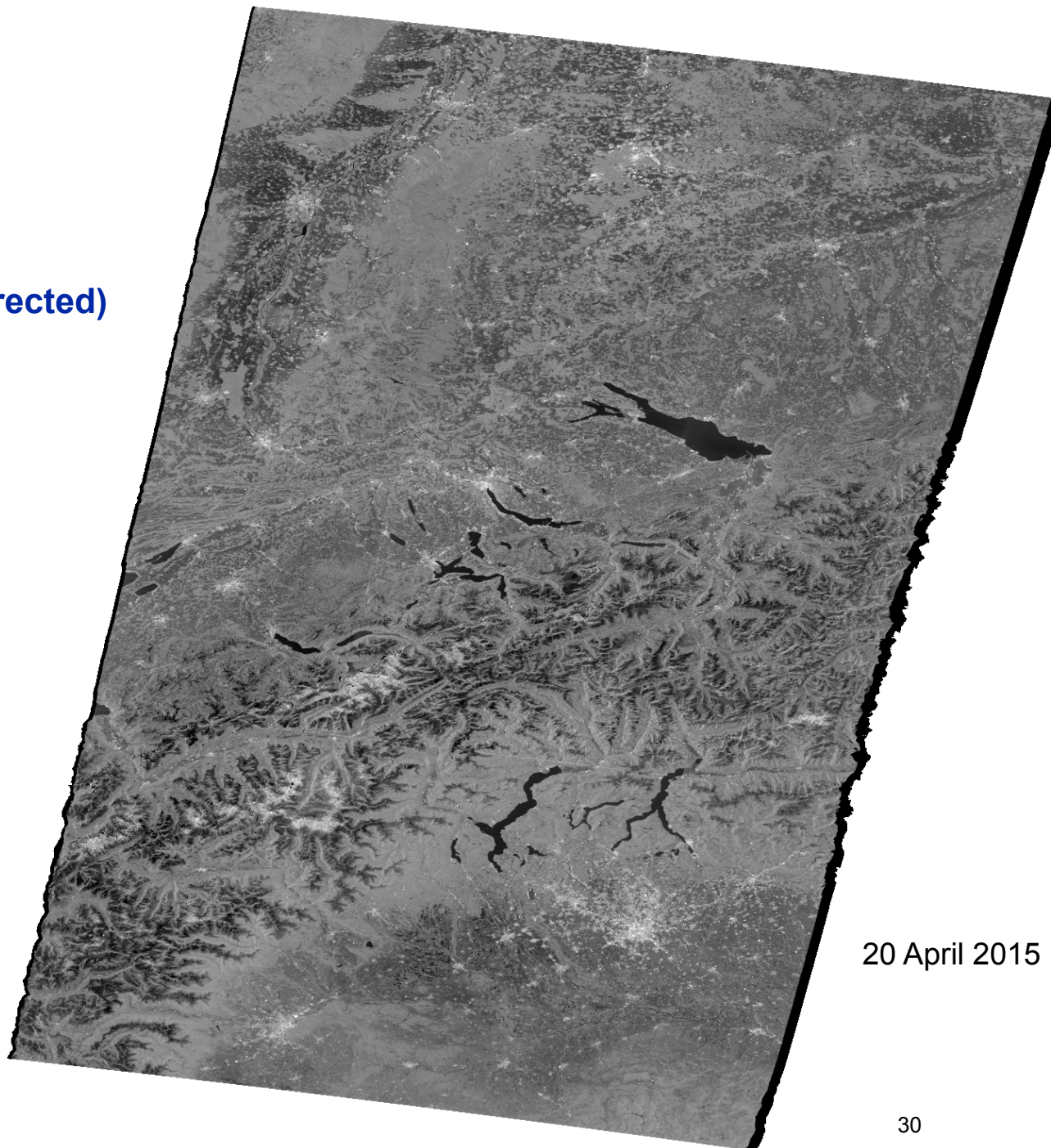
$$\gamma_T^0$$

-26dB      -1dB



Generated automatically from  
3 IW GRDH products using  
SRTM3

Contains modified  
Copernicus Sentinel data (2015)



20 April 2015



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# Sentinel-1A: **GTC** (Geometrically Terrain Corrected)

Dept. of Geography / Remote Sensing Laboratories

Conventional Range-Doppler Orthorectification

Switzerland

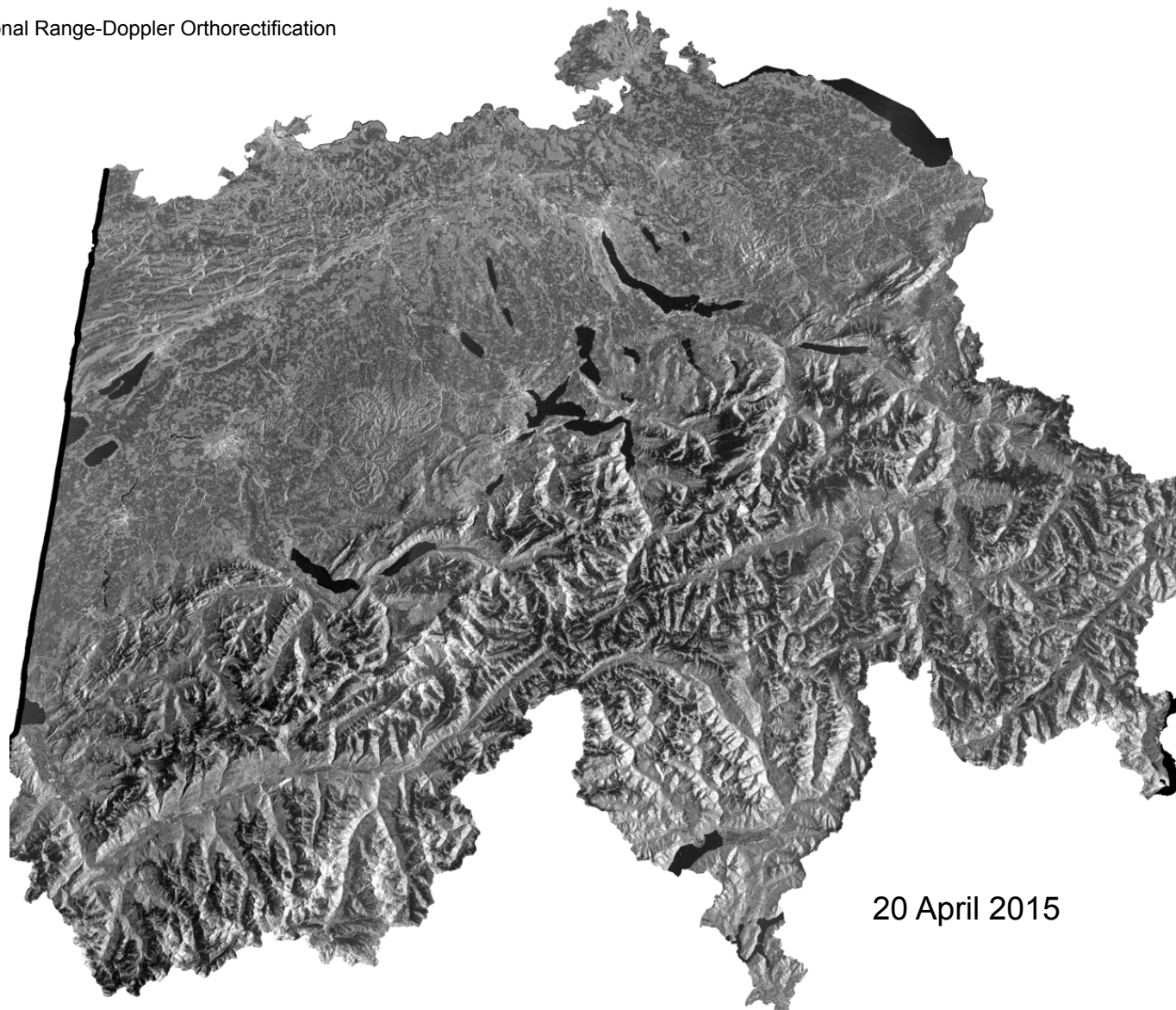
$\gamma_E^0$  VH-pol.

-26dB -1dB



Generated from  
3 IW GRDH products using  
SwissALTI3D DEM (10m)

Copernicus Sentinel data  
(2015)



20 April 2015





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# Sentinel-1A: RTC (Radiometrically Terrain Corrected)

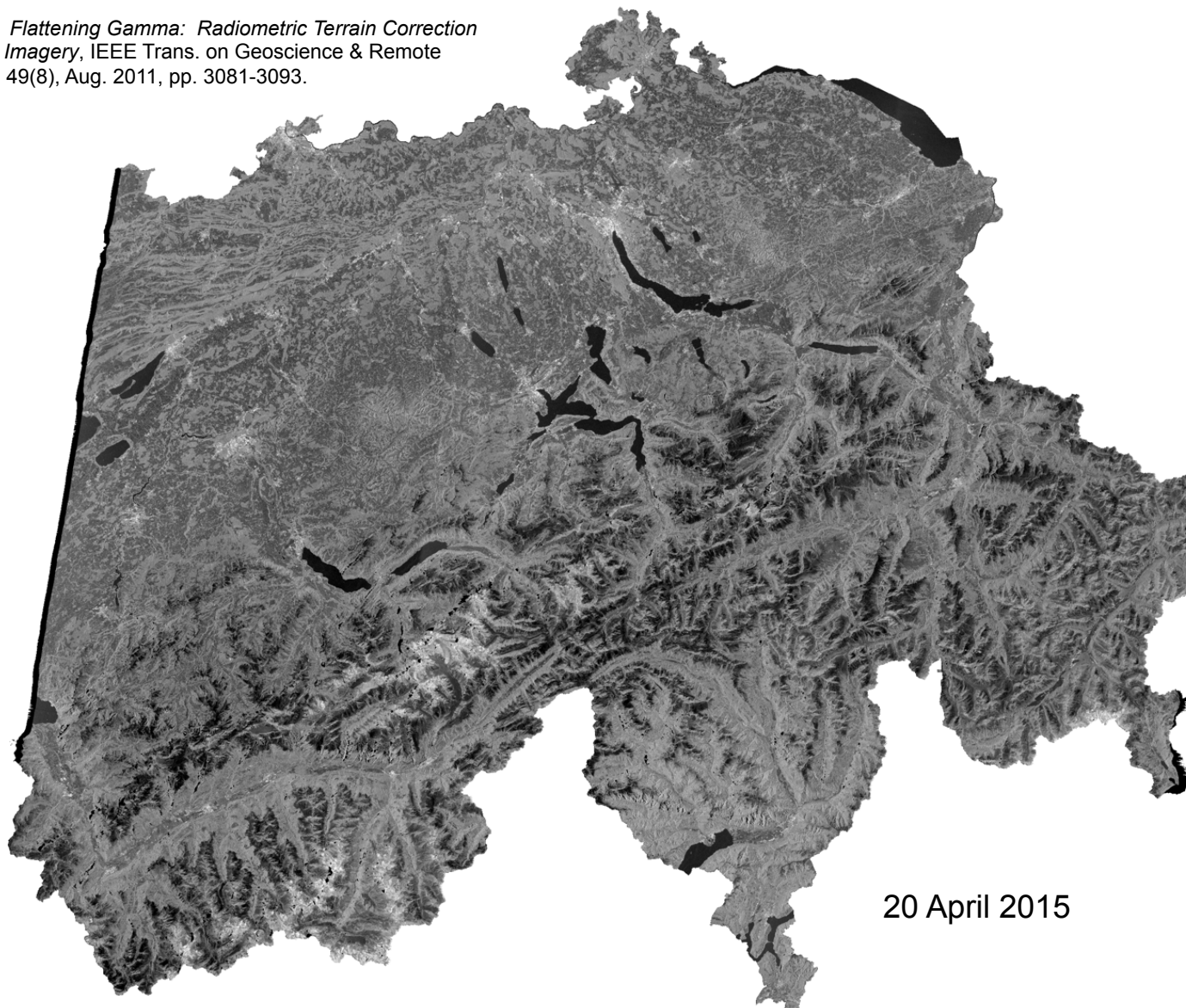
Dept. of Geography / Remote Sensing Laboratories

Small D. *Flattening Gamma: Radiometric Terrain Correction for SAR Imagery*, IEEE Trans. on Geoscience & Remote Sensing, 49(8), Aug. 2011, pp. 3081-3093.

Switzerland

$\gamma_T^0$  VH-pol.

-26dB -1dB



Generated from  
3 IW GRDH products using  
SwissALTI3D DEM (10m)

Contains modified  
Copernicus Sentinel data  
(2015)

20 April 2015

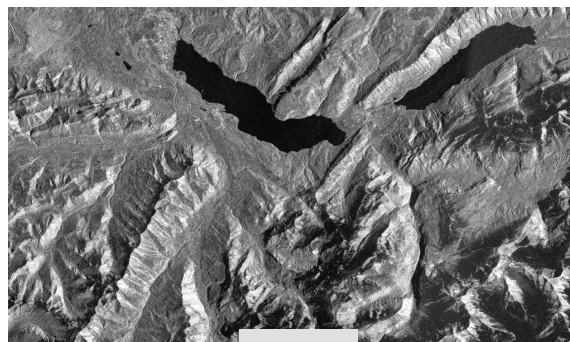




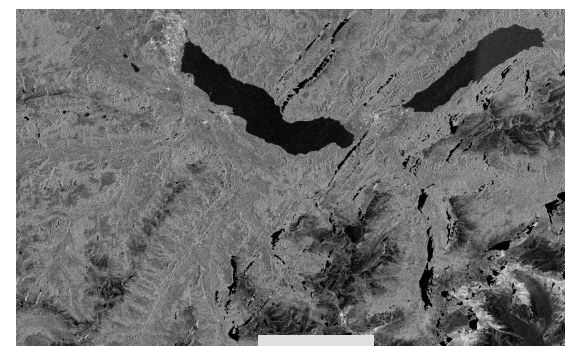
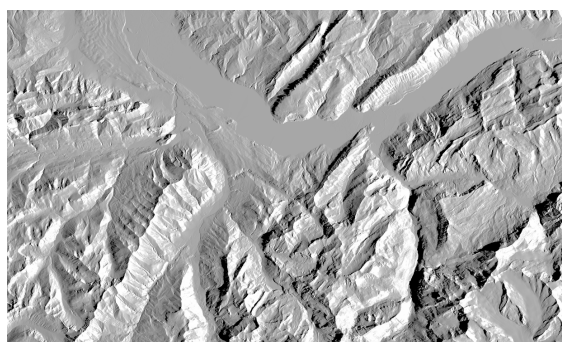
$$\gamma_E^0$$

$$\gamma_T^0$$

2015.05.26 (Desc.)

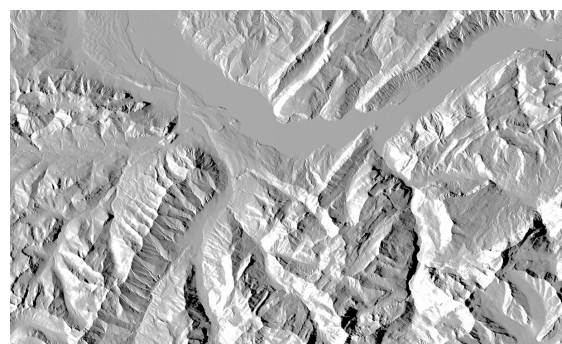


GTC



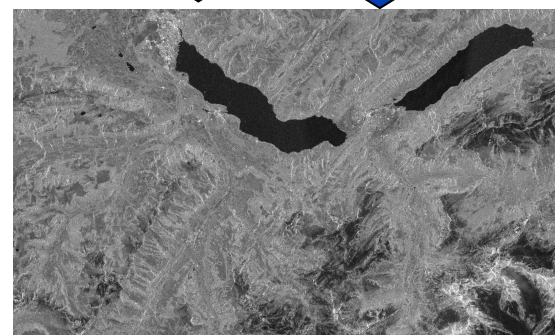
RTC

2015.05.27 (Asc.)



- Combine asc. & desc. observations to generate **composite** with improved local resolution
- Less shadow than single RTC, lower noise

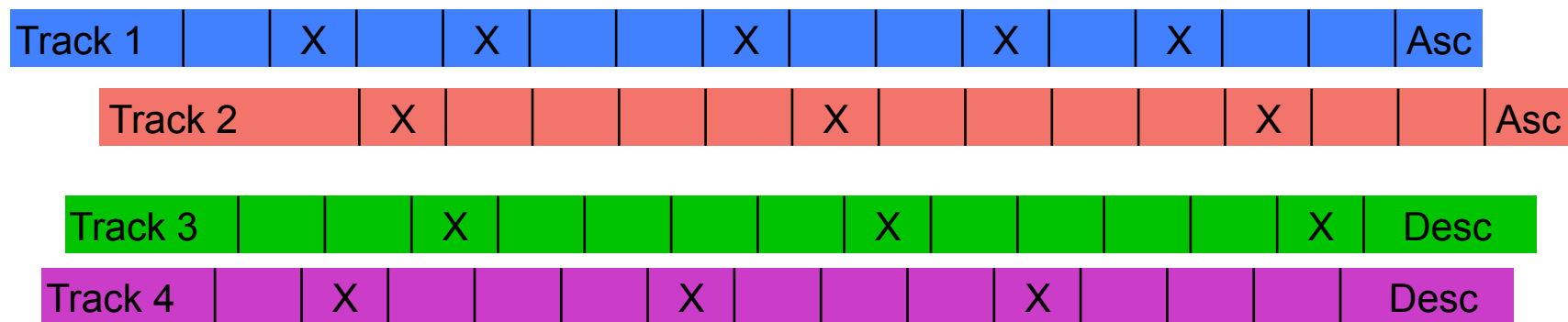
Interlaken, Switzerland



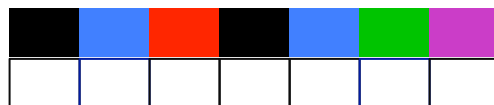
Composite



## Revisit Interval: Breaking the tyranny of exact repeat passes



Backscatter contributions



Weights inverse to local area



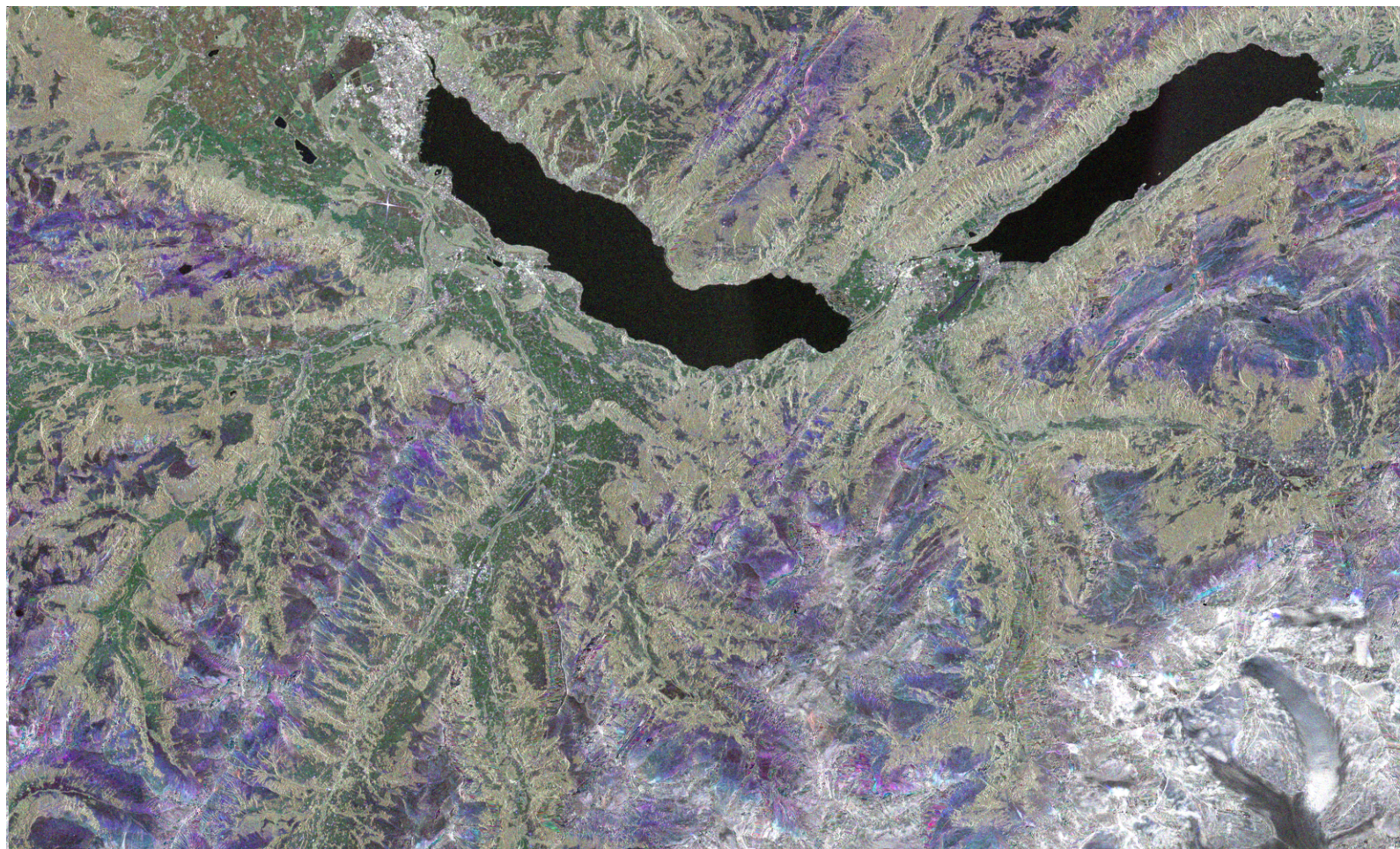
For *Regular Intervals* with temporal resolution better than repeat-pass interval

- Use moving time-window integrating information from all tracks
- The more (diverse!) data (and tracks) the better – esp. combine ascending and descending observations





R=2015.01.02+03 / G=2015.01.14+15 / B=2015.02.07+08 ; (each Asc. + Desc.)

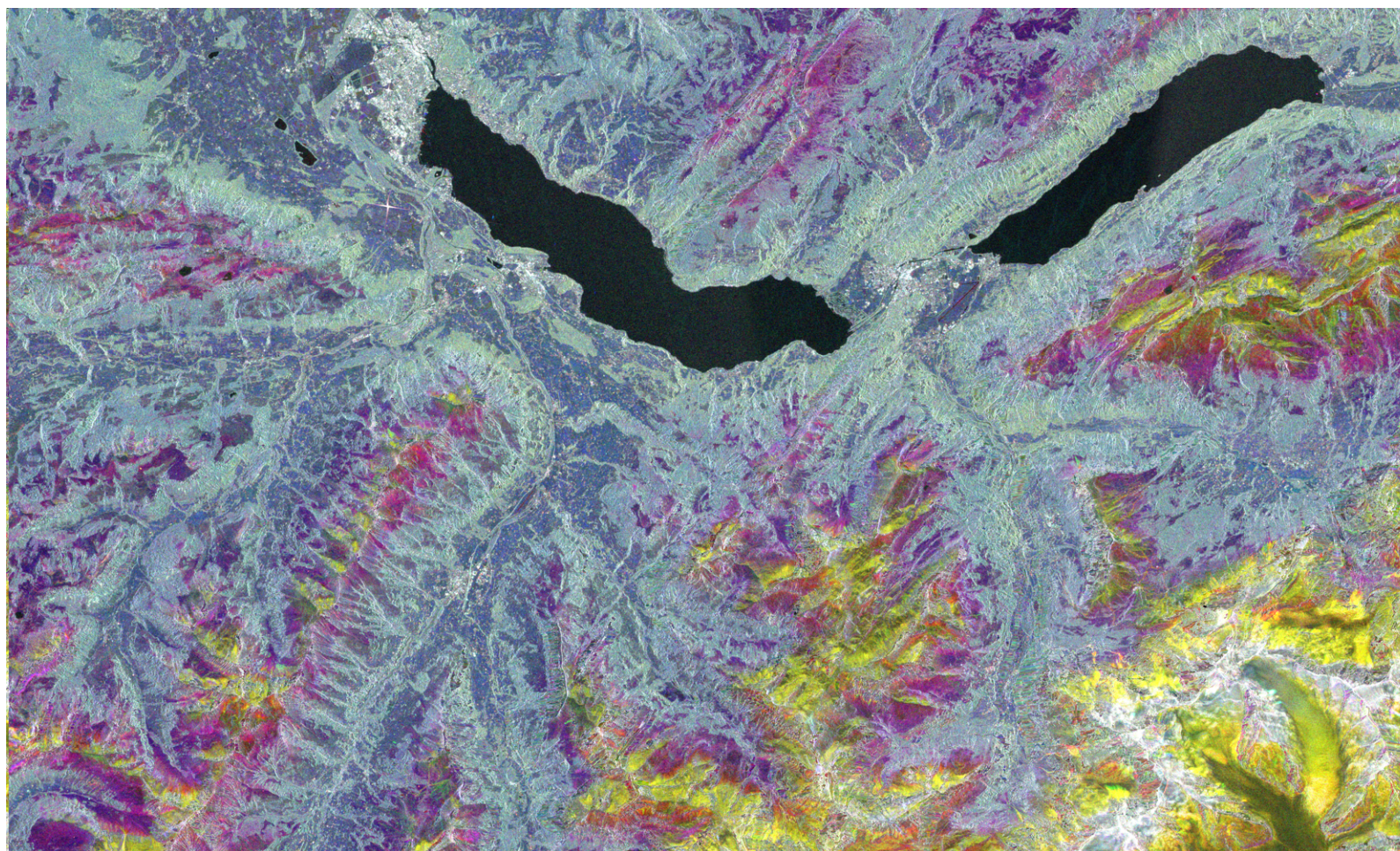


Contains modified  
Copernicus  
Sentinel data (2015)





R=2015.02.07+08 / G=2015.04.08+09 / B=2015.05.26+27 ; (each Asc. + Desc.)



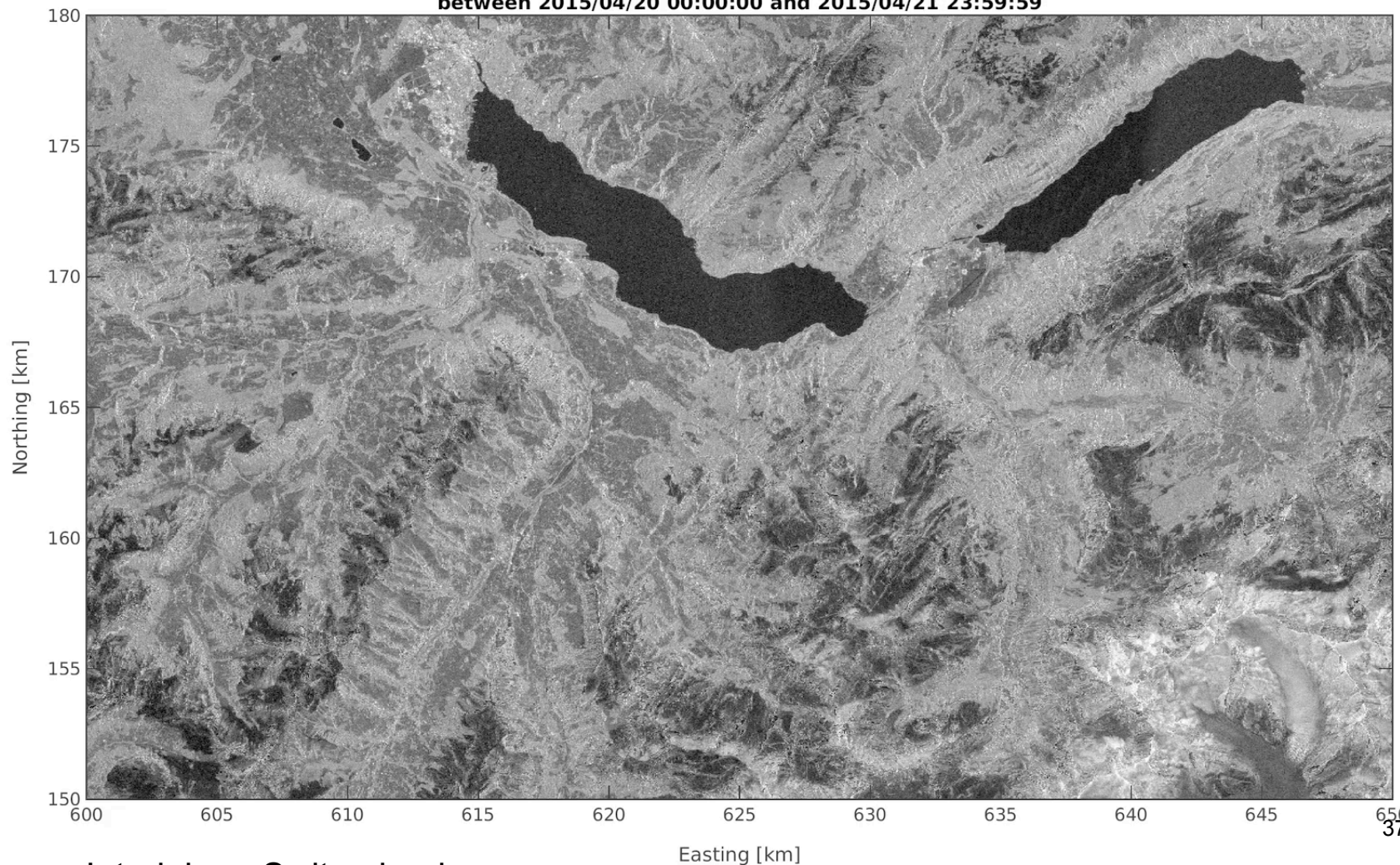
Contains modified  
Copernicus  
Sentinel data (2015)





Jan – May 2015

Composite backscatter from 2 scenes  
between 2015/04/20 00:00:00 and 2015/04/21 23:59:59



Contains modified  
Copernicus  
Sentinel data (2015)

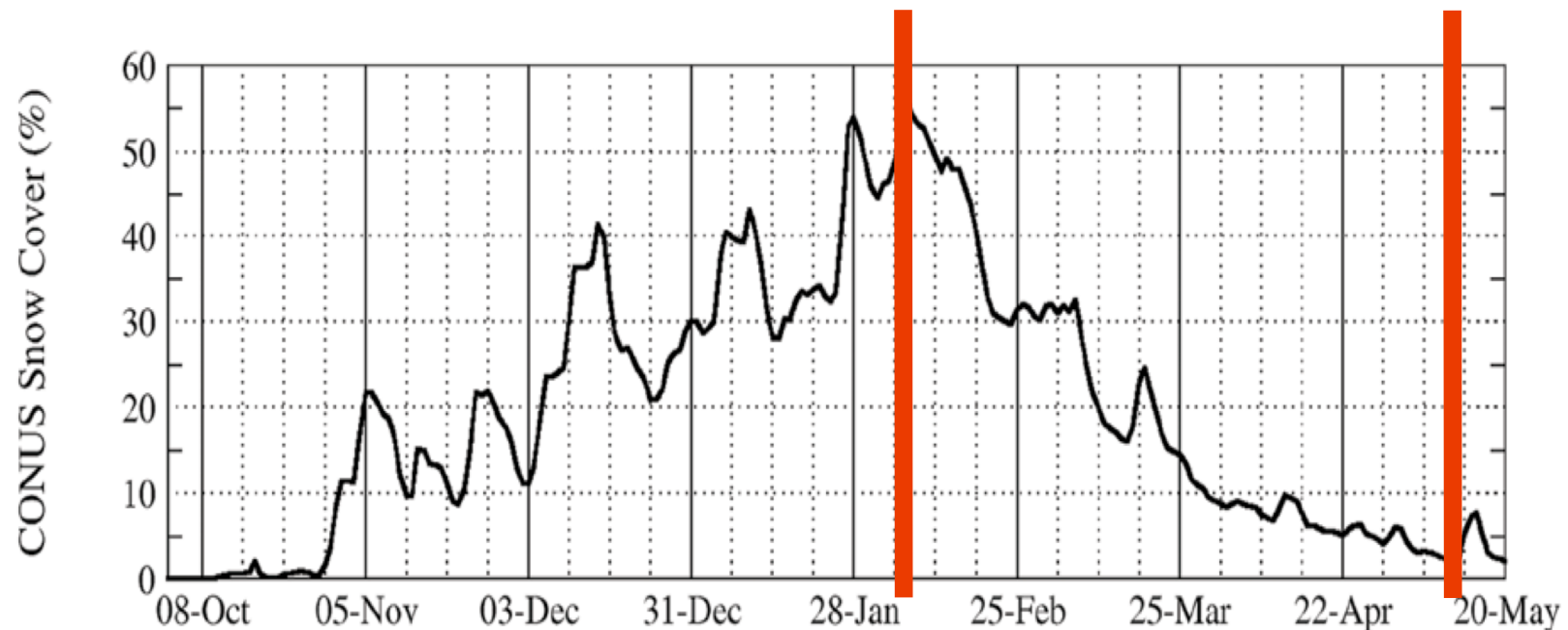
Interlaken, Switzerland



## Evolution of Snow Cover in the US

[IGOS Cryosphere Theme Report, 2007]

**Fig. 3.1.** Percentage of snow-covered area within the conterminous U.S. during the 2003-2004 season with corresponding unique snow depths and SWE.





## Swiss Seasonal Hydrology: 1971-2012 Daily SWE plotted by elevation

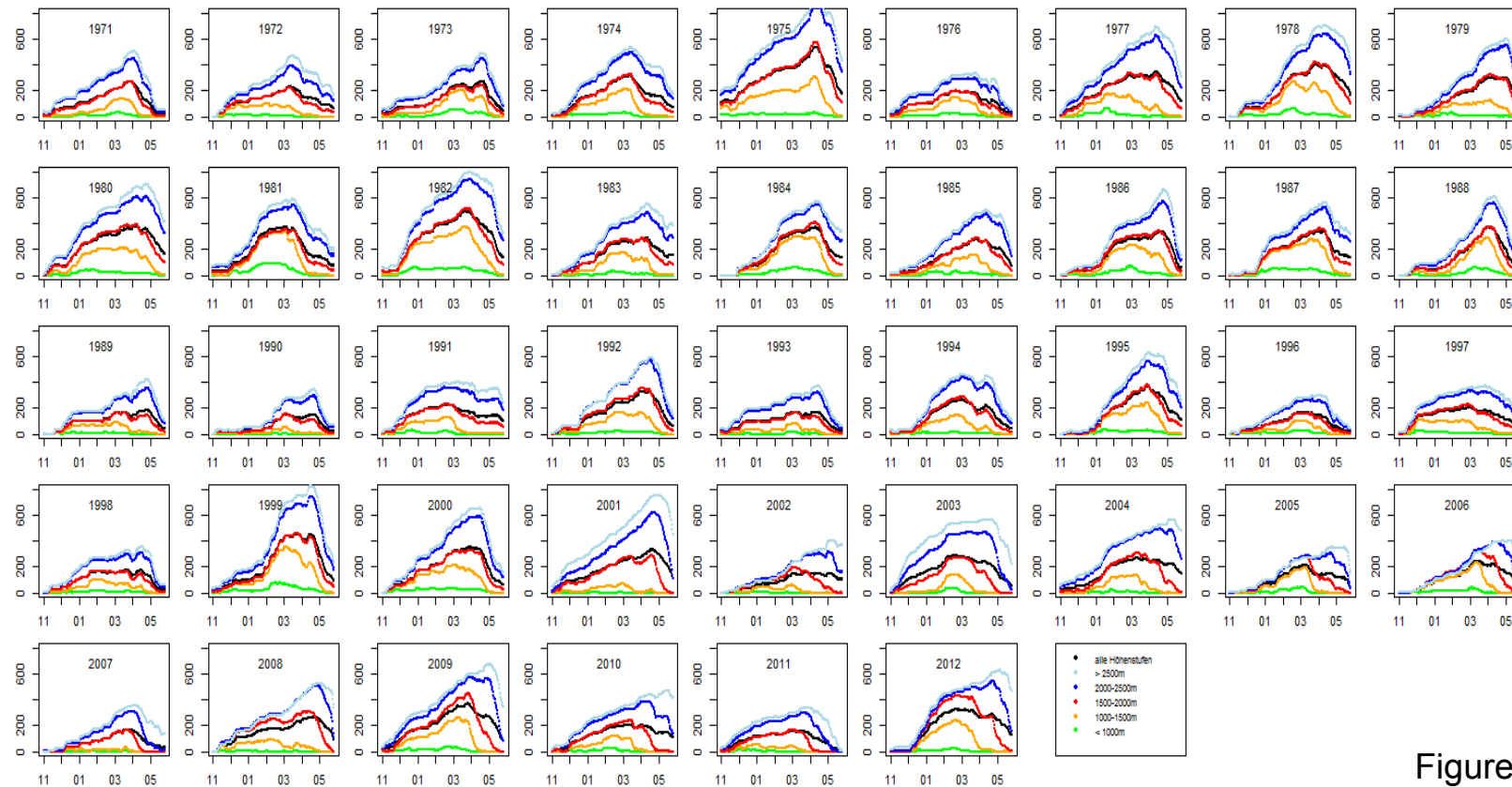


Figure courtesy  
Tobias Jonas, SLF

Melting generally captured at significant elevations between Feb. 15 and May 15



## Key General Recommendations of *WMO White Paper* on SAR Acquisition Planning for Terrestrial Snow Monitoring

R1	Use <b>wide-swath modes</b> to enable wide area monitoring with high temporal resolution (i.e. RSAT2 SCN or SCW, Sentinel-1 IW or EW, TSX “SC Wide” & CSK “Huge Region” ScanSAR modes).
R2	Build combined <b>ascending/descending</b> coverage by default into acquisition plans covering mountainous regions. Favour asc./desc. acquisition sets acquired within a <b>tight time window</b> (1-3 days) to allow a narrow time-attribution to composites generated from these sets.
R3	Concentrate snowmelt acquisitions on the <b>seasonal window</b> when the majority of snow melting occurs (March through May at temperate northern latitudes). The <i>highest temporal resolution possible</i> is requested during this critical melting period. Although some further acquisitions are also requested <i>outside</i> of this seasonal window, lower temporal resolution at these less critical times is acceptable.
R4	Standardise dual-pol. mode acquisitions on <b>VV/VH</b> combination: a cross-platform consistent polarisation simplifies combination of datasets from multiple providers (e.g. S1/RSAT2/RCM or TSX/CSK).
R5	<b>Harmonise acquisition plans</b> of satellites with compatible calibrated backscatter values (e.g. S1/RSAT2/RCM or TSX/CSK). Utilise the available diversity of orbits to achieve the desired diversity of tracks – e.g. to achieve the fullest possible ascending/descending coverage.
R6	Assure <b>full coverage over land also in coastal regions</b> when other modes are by default programmed over ocean (e.g. favour Sentinel-1 IW or EW over WV).
R7	Maintain a <b>regular observation plan also during the winter</b> to assure frequent observations of other important snow parameters, and other phenomena related to the winter period such as avalanches and rain on snow events.





## Science Requirements for wide area snowmelt monitoring

**Spatial resolution:** 100m ✓

Variable	Extent	Spatial resolution	Temporal resolution	Sensor	Auxiliary Data
Snowmelt area	Regional	100m	1 to 5 days	Sentinel-1	Land cover, DEM
Snowmelt liquid water content	Regional	100m	1 to 5 days	Sentinel-1 dual polarisation	Land cover, DEM

[Malenovsky, Z. et al. *Sentinels for science: Potential of Sentinel-1, -2, and -3 missions for scientific observations of ocean, cryosphere, and land*. Remote Sens. Environ. 120, 91–101 (2012)]

**Temporal resolution (target):** 1 day ✗

- “Observation of the **daily geographic extent of snow cover** is essential because it enables inference of several first order effects of snow on many Earth systems.” [IGOS Cryosphere Theme Report, 2007]
- WMO PSTG report “Coordinated SAR Acquisition Planning for Terrestrial Snow Monitoring”, PSTG-SARCWG-SNOW-001, Aug. 2014.



## Data Collections

Region	DEM	Spatial sampling	Temporal resolution [days]	Sensors
Interlaken region, Switzerland	swissALTI3D (2m)	10 m	(selected) 2	S1A IW DV
European Alps	SRTM3 (3s)	3s (~90m)	16	S1A IW DV, RS2 SCW/SCN VV/VH
Coastal British Columbia, Canada	SRTM3 (3s)	3s (~90m)	24	S1A IW SV
Ellesmere Island, Canada	CDEM <sup>1</sup>	400m	4	S1A EW DH RS2 SCWA HH/HV

<sup>1</sup>M. Santoro & T. Strozzi (2012): **Circumpolar digital elevation models > 55° N Canadian Digital Elevation Model Product Specifications**, Edition 1.1, 2013-04-01, GeoGratis



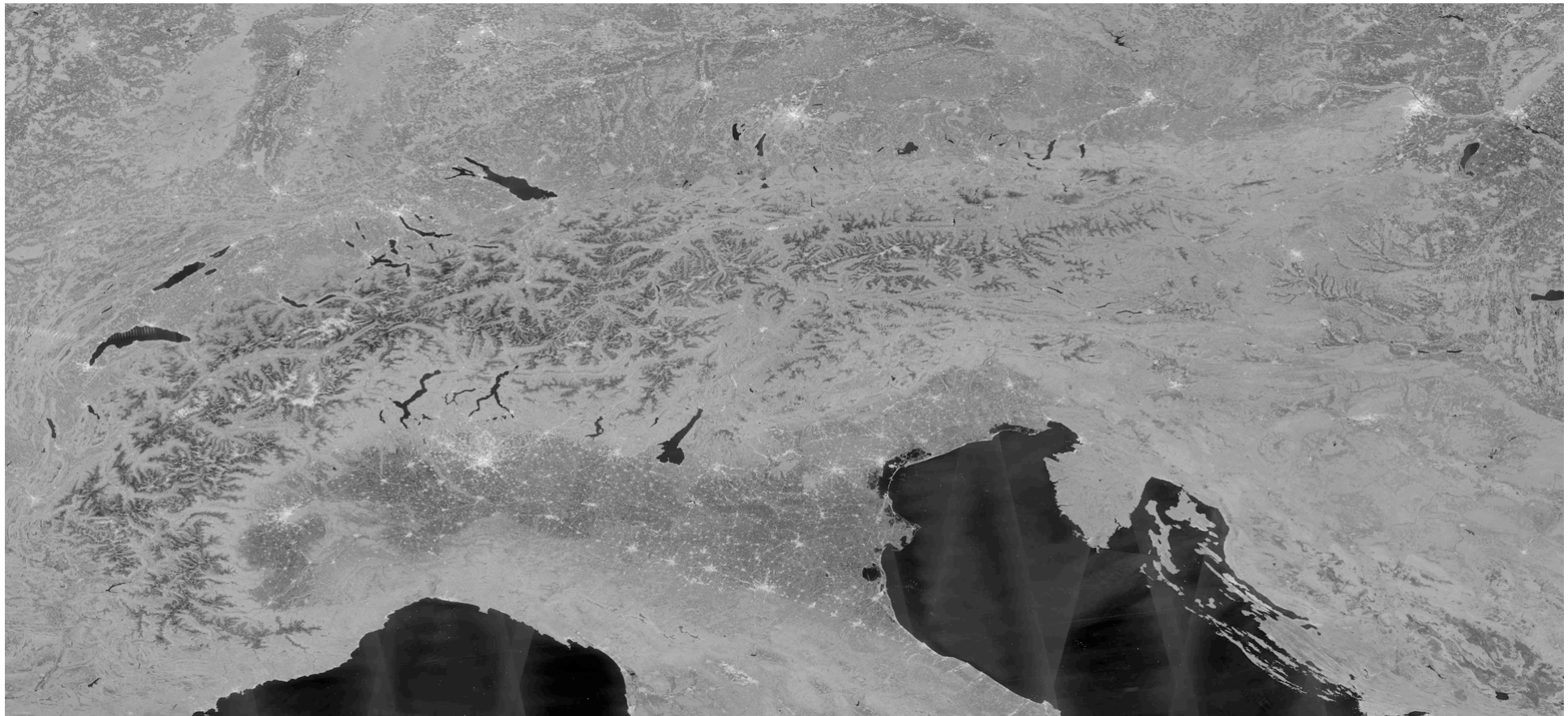
University of  
Zurich <sup>UZH</sup>

# S1A Alps Backscatter Mosaic

Dept. of Geography / Remote Sensing Laboratories

Contains modified Copernicus Sentinel data (2015)

S1A IW VH & VV-pol. Oct. 2014 – Aug. 2016: 12d & 16d windows  
**Jan.-Aug. 2016 VH 16d shown here**



SRTM3 used for geometric and radiometric corrections



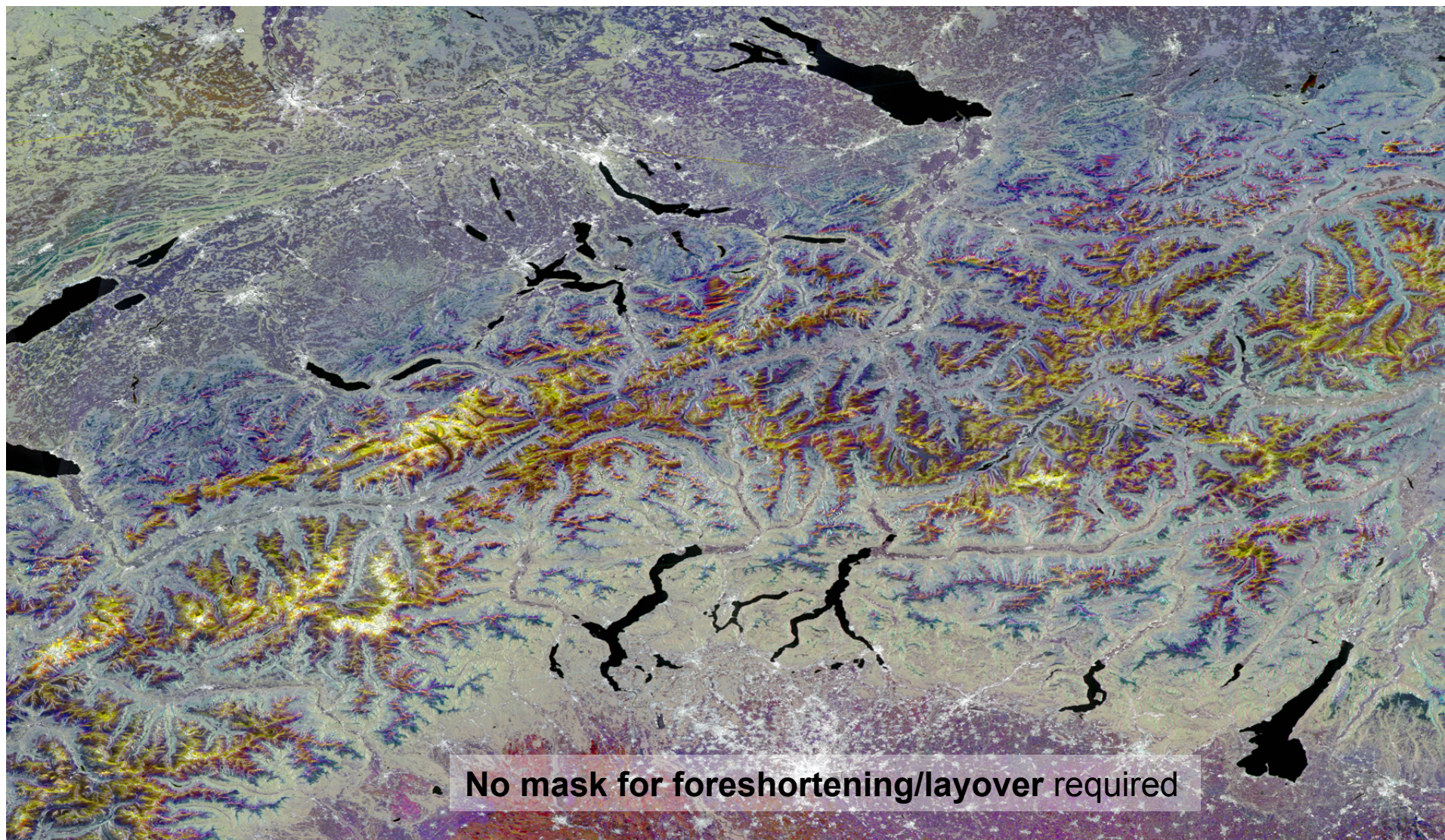


**University of  
Zurich** <sup>UZH</sup>

Contains modified  
Copernicus  
Sentinel data (2015)

## Dept. of Geography / Remote Sensing Laboratories

Sentinel-1 IW 16d Composites 2015 VH: **March 14-29**, **April 7-22**, **May 25-June 9**; -23dB (black) to -6dB (white)



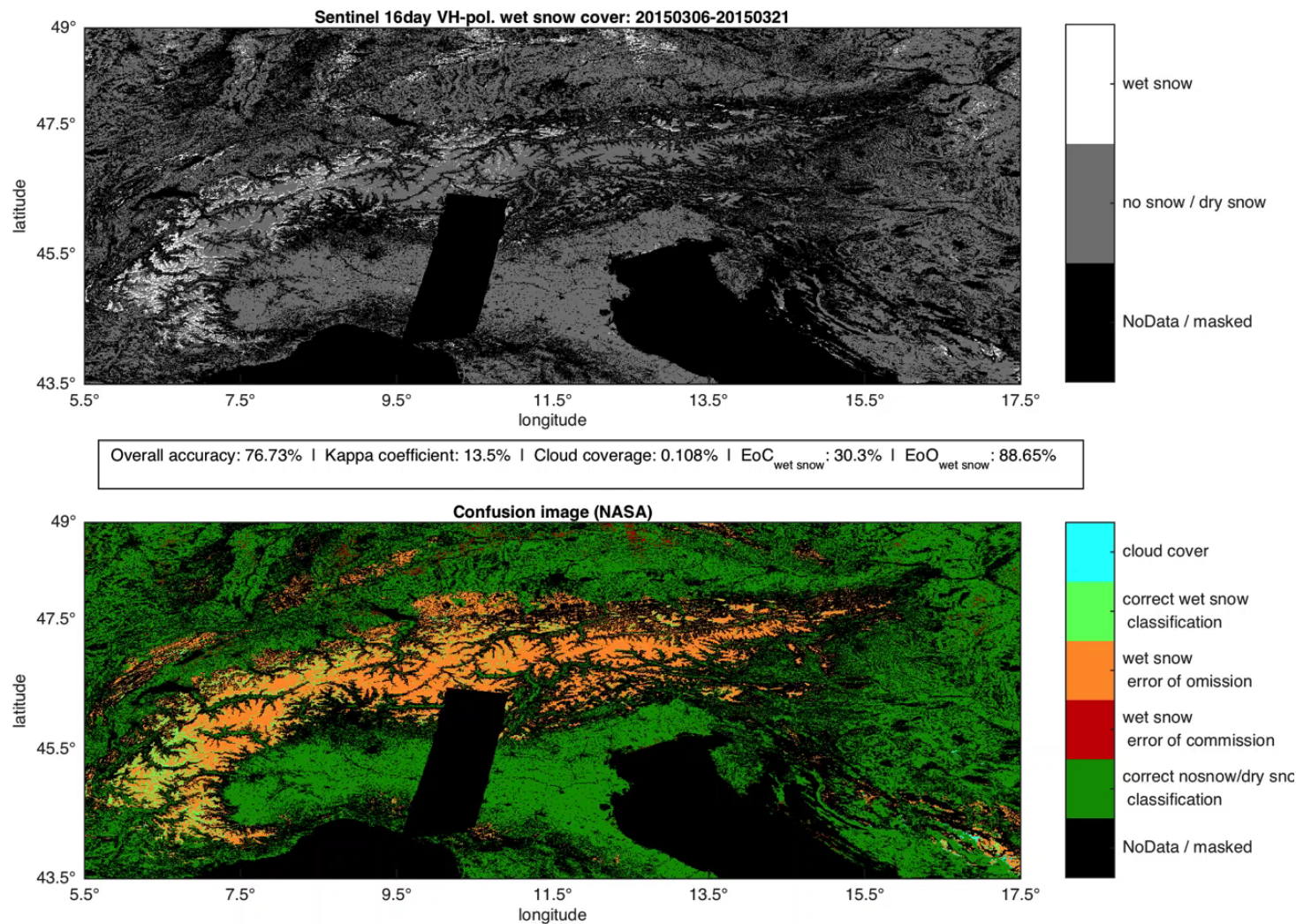
**No mask for foreshortening/layover required**





S1A IW 2015  
VH & VV-pol.

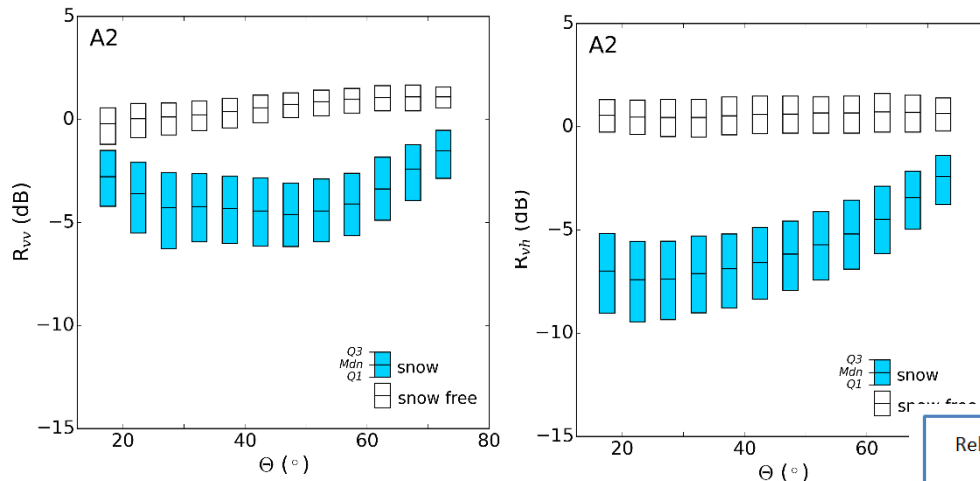
S1-based wet  
snow  
classifications  
compared with  
NASA MODIS  
snow products





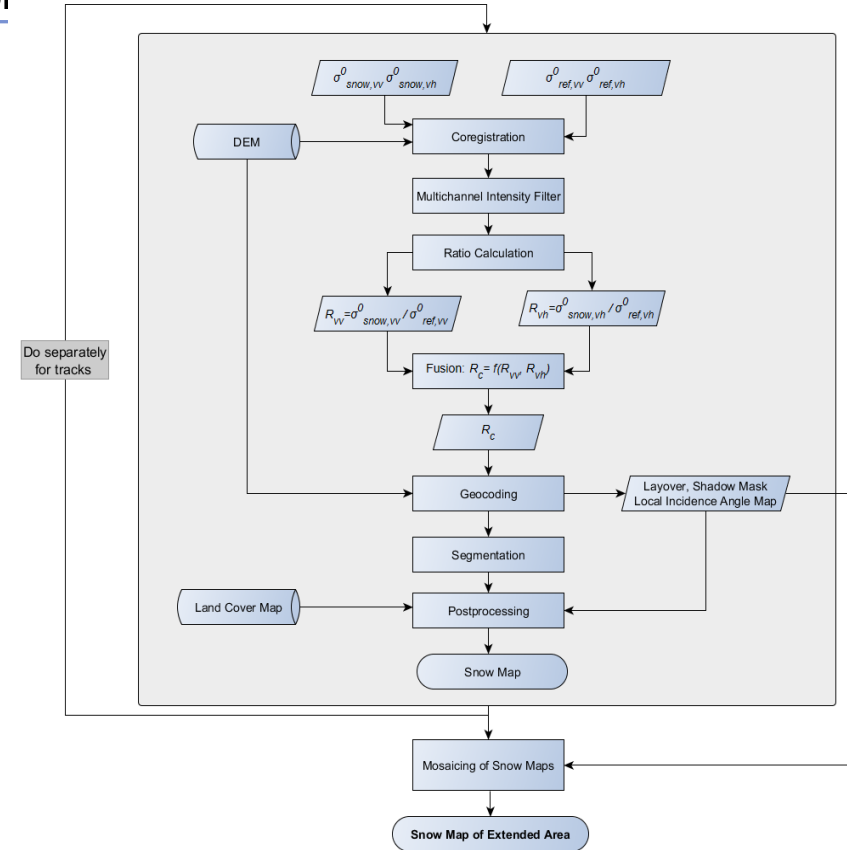
# Sentinel-1 Dual Pol Snow Mapping Method (ENVEO)

Backscatter ratio (median, Mdn, 1<sup>st</sup> and 3<sup>rd</sup> quartile) for Sentinel-1 VV- and VH- polarized channels in dependence on local incident angle. Test area Ötztal.



Nagler et al. 2016; Rem. Sens., 2016, 8(4), 348, doi:10.3390/rs8040348

Figure courtesy Thomas Nagler



Relation for merging  $R_{vv}$  and  $R_{vh}$  ratios in order to create a combined single channel,  $R_c$ :

$$R_c = W R_{vh} + (1 - W) R_{vv}.$$

With:

$$IF (\theta < \theta_1) \rightarrow \{W = 1.0\}$$

$$IF (\theta_1 \leq \theta \leq \theta_2) \rightarrow \left\{W = k \left[1 + \frac{(\theta_2 - \theta)}{(\theta_2 - \theta_1)}\right]\right\},$$

$$IF (\theta > \theta_2) \rightarrow \{W = k\}$$

We use  $k = 0.5$ ,  $\theta_1 = 20^\circ$ ,  $\theta_2 = 45^\circ$ ,  $\theta$  is the local incidence angle.

Wet snow segmentation rule:  $R_c < THR$ , with  $THR = -2$  dB.

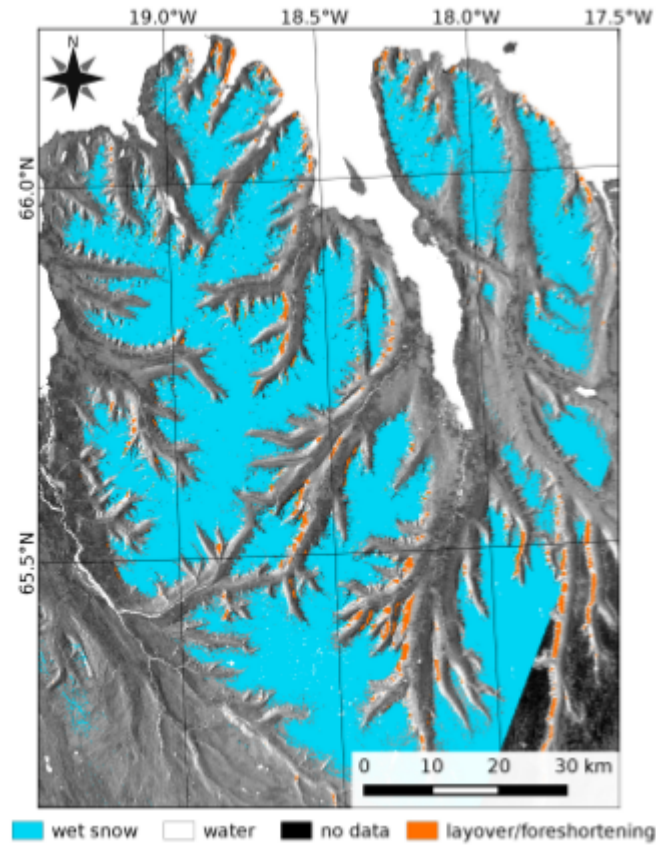
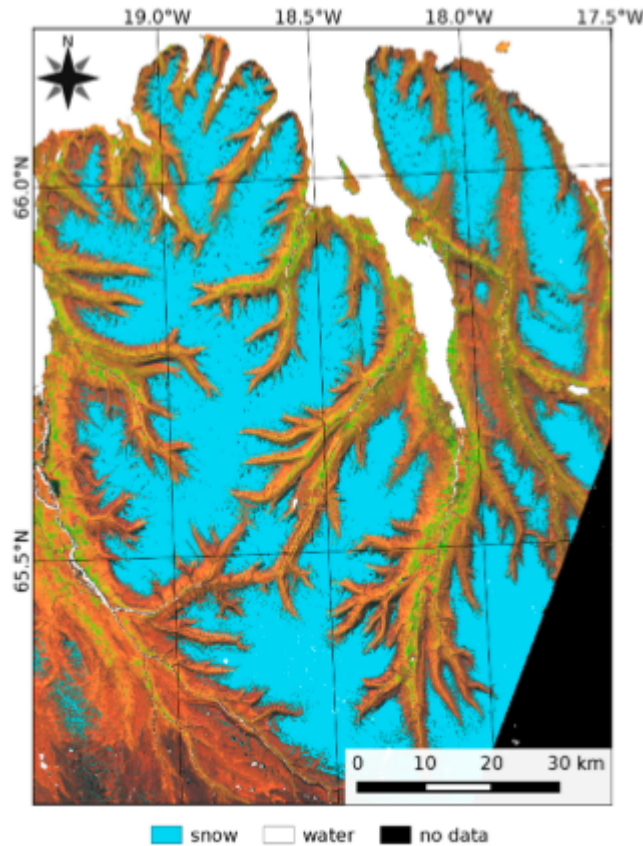


# Example of S1 Snow Melt and Landsat TM Snow Extent – Tröllaskagi Peninsula, Iceland

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Landsat-8, 27 June 2015;

Sentinel-1, 26 June 2015;



	$R_c$		
	S1-S	S1-F	AR
LS-S	94.6	5.4	0.972
LS-F	0.2	99.8	

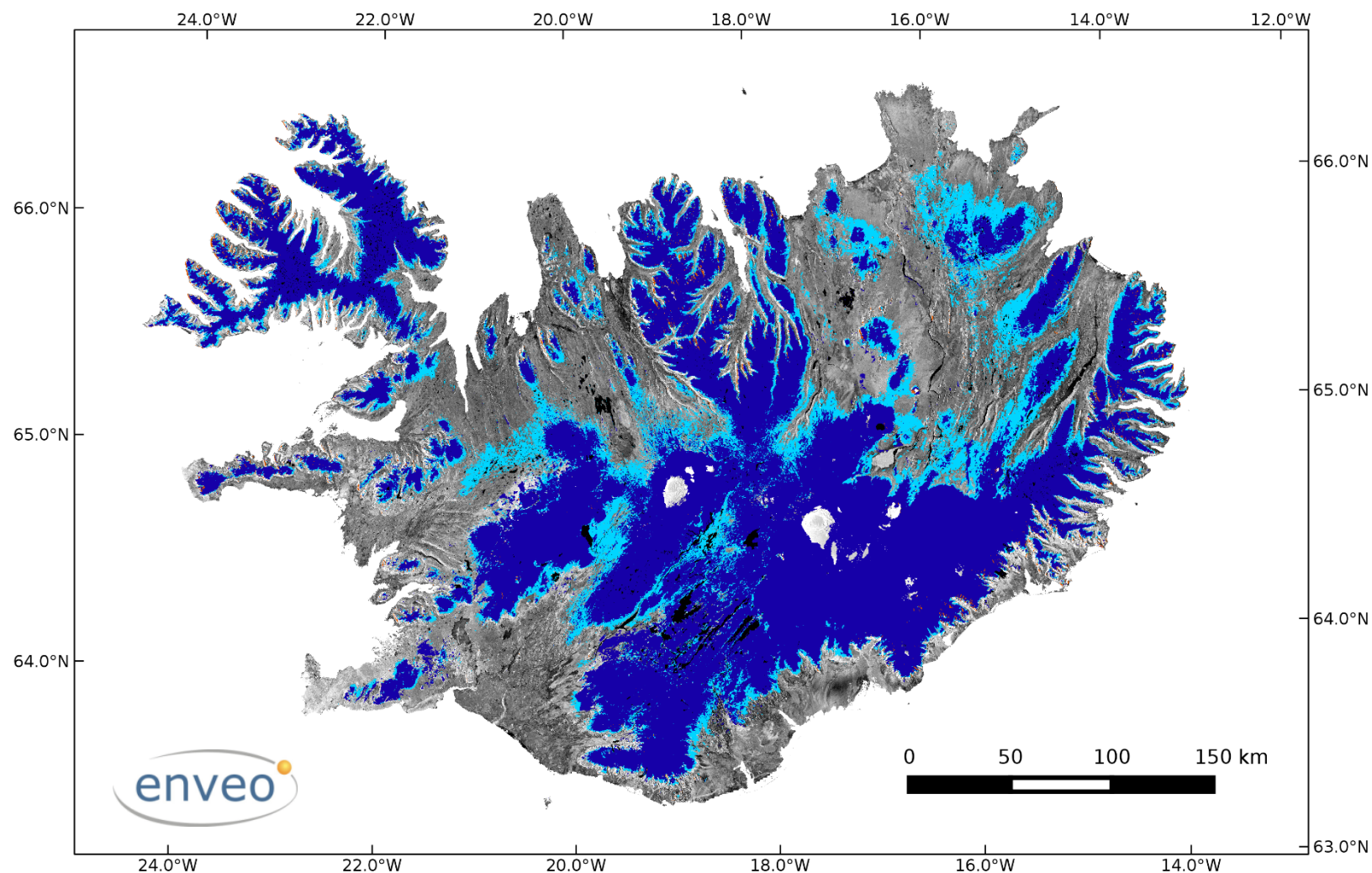
Confusion matrix for the classes snow (S) and snow-free, for snow classification based on Landsat (LS) and Sentinel-1 (S1) data. S1 results are shown for snow maps based on . — overall agreement rate ().

Figure courtesy Thomas Nagler





# Monitoring melting snow using Sentinel-1



decrease of wet snow extent from May 17 and 22 to June 10 and 15, 2015

wet snow extent on June 10 and 15, 2015

layover / foreshortening





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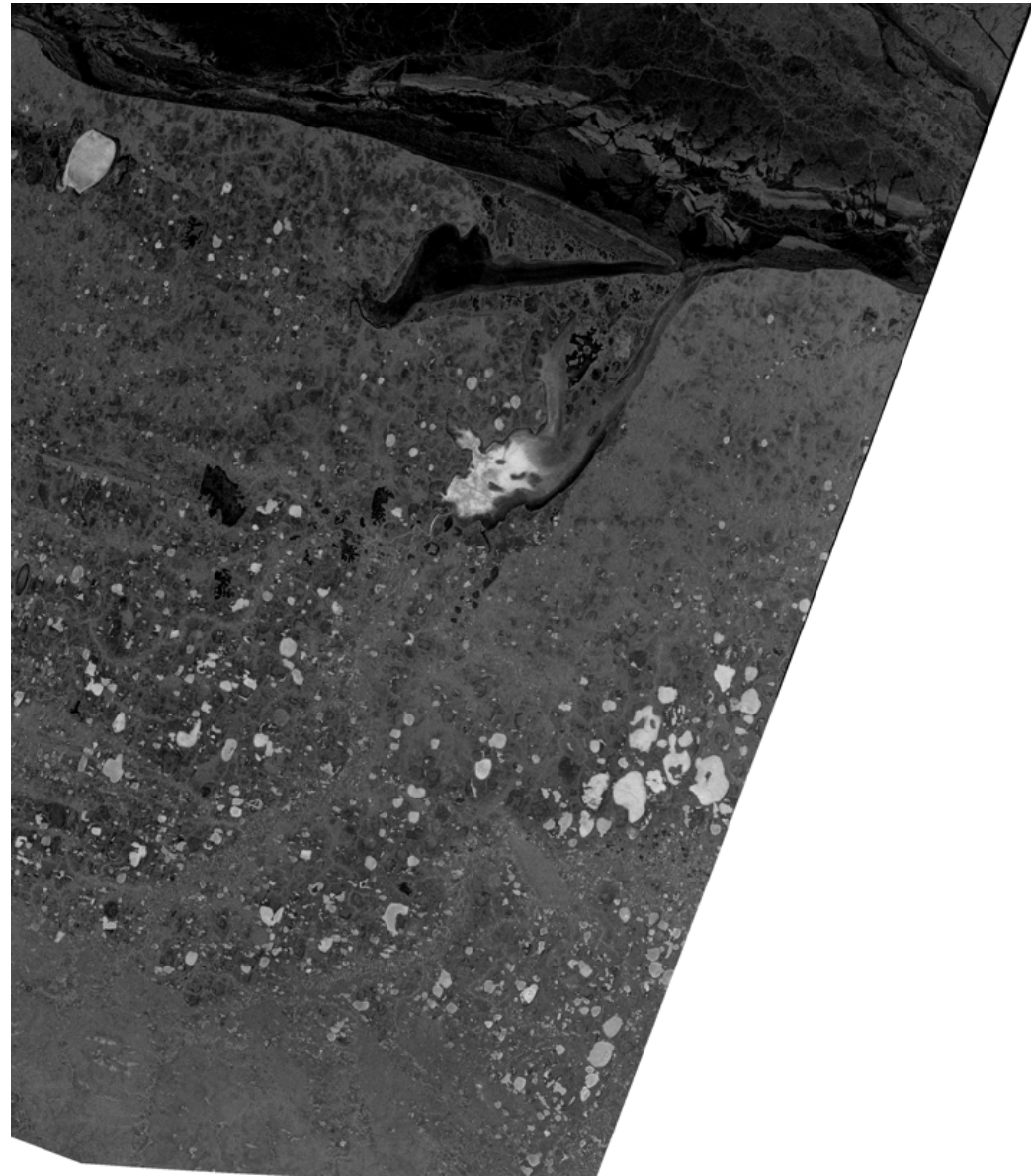
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## Kytalyk, Siberia

Sentinel-1 EW HH-pol.  
Backscatter

20150412





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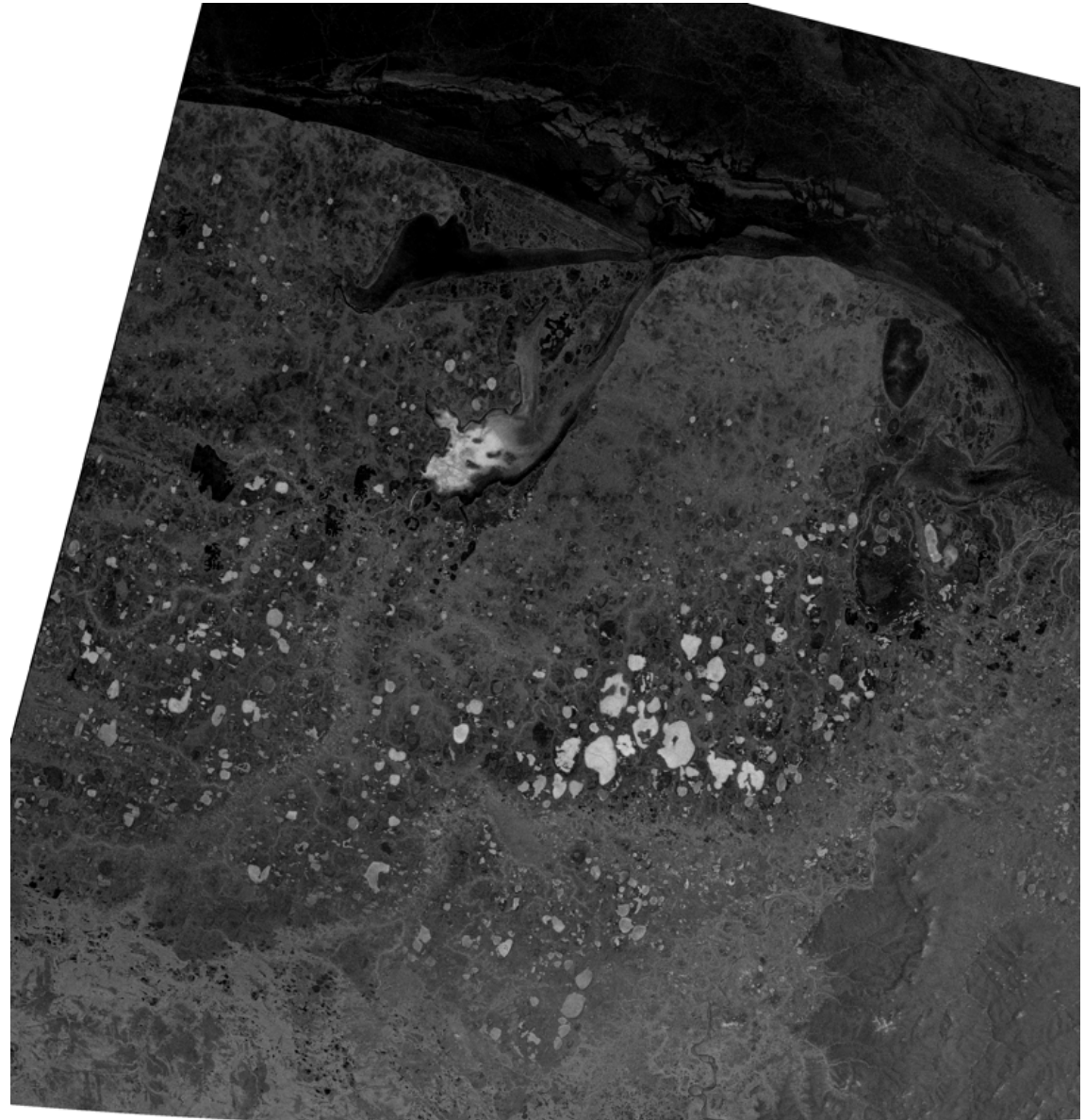
Dept. of Geography / Remote Sensing Laboratories

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## Kytalyk, Siberia

Sentinel-1 EW HH-pol.  
Backscatter

20150527





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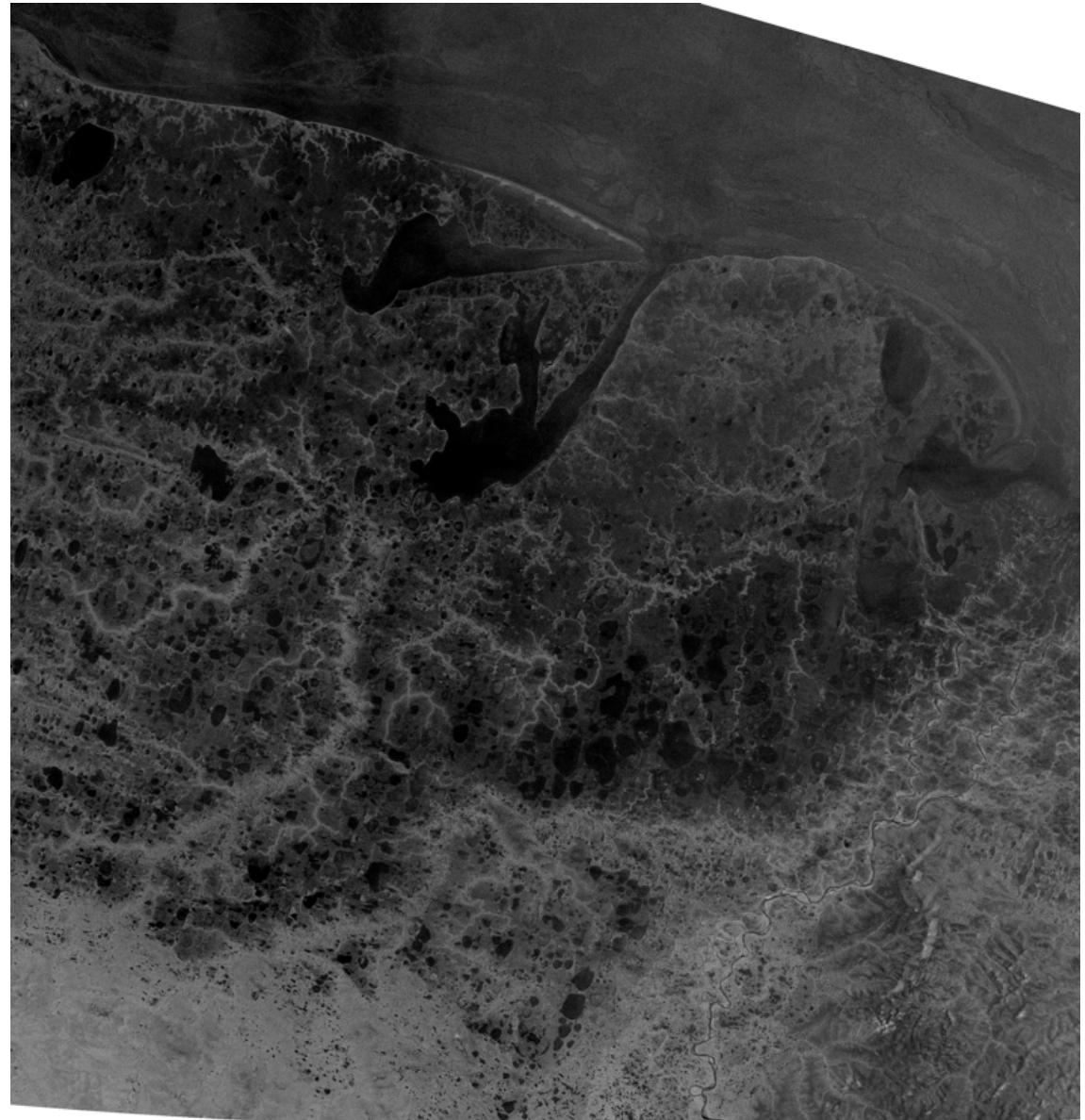
Dept. of Geography / Remote Sensing Laboratories

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## Kytalyk, Siberia

Sentinel-1 EW HH-pol.  
Backscatter

20150601







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## Ground- based sensing

e.g.  
Phenocam  
in Kytalyk,  
Siberia

Movie courtesy  
G. Ghielmetti,  
UZH-RSL





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## Coastal British Columbia Backscatter Composites

**S1A IW VV**

12 day delta

24 day window

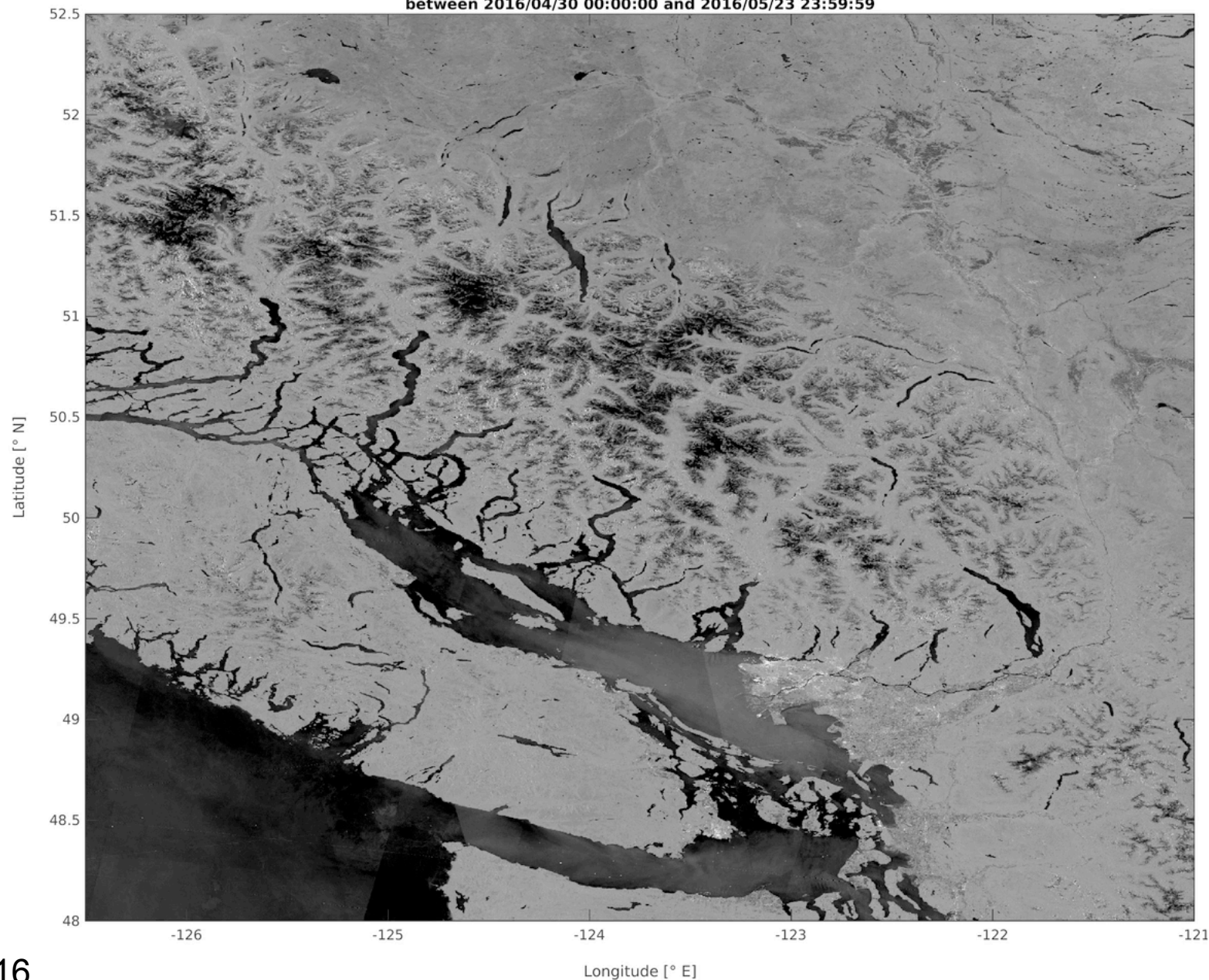
N.B.

*Increased **dual-**  
pol VV/VH  
acquisitions in  
last months)*

Jan. – Aug. 2016

Contains modified Copernicus Sentinel data (2016)

Composite backscatter from 37 scenes  
between 2016/04/30 00:00:00 and 2016/05/23 23:59:59







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## Ellesmere Island Backscatter Composites

RS2 SCWA

HH

2 day delta

4 day window

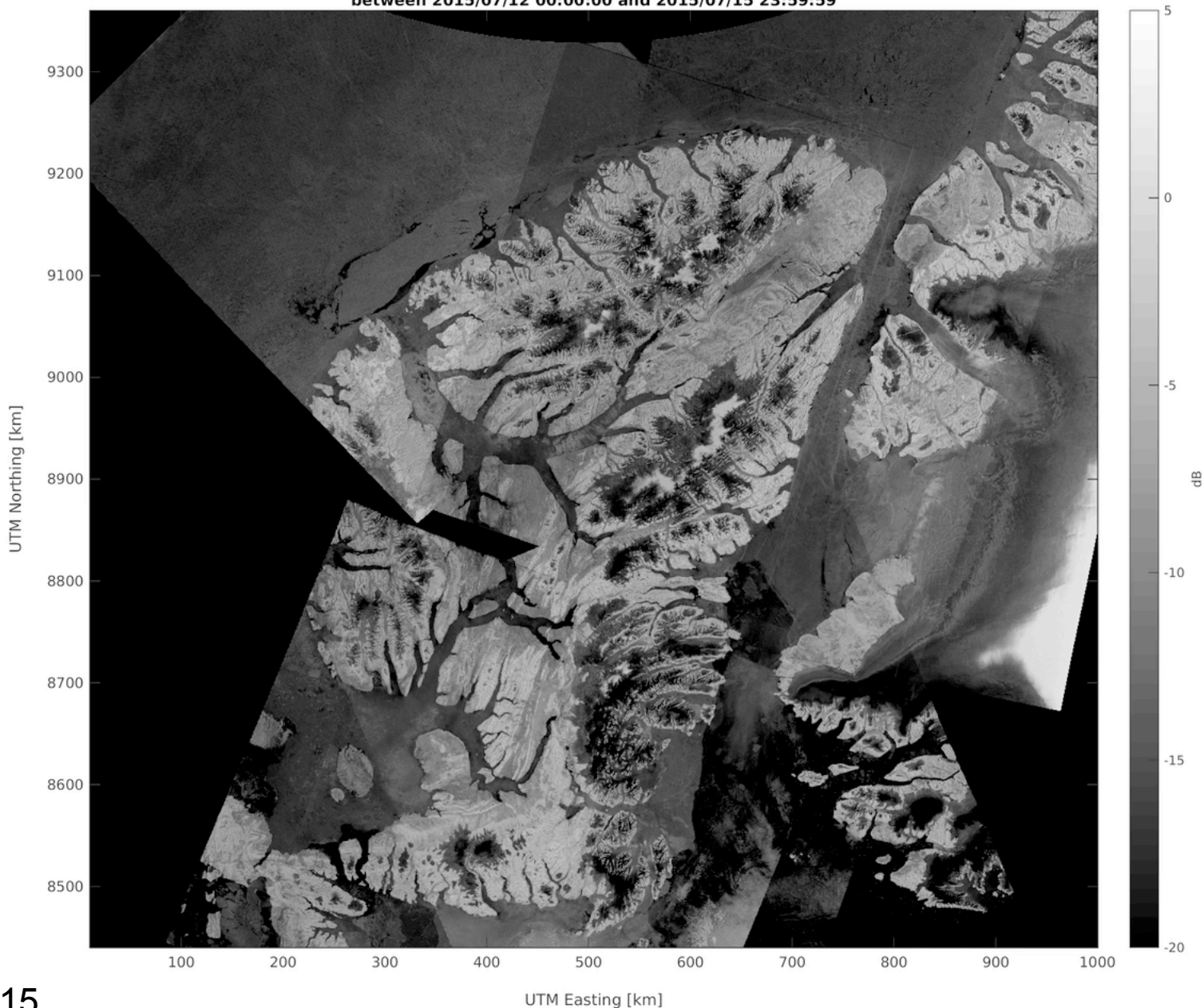
N.B.

8 bit radiometry

CDEM

May – Aug. 2015

Composite backscatter from 7 scenes  
between 2015/07/12 00:00:00 and 2015/07/15 23:59:59



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## Ellesmere Island Backscatter Composites

RS2 SCWA  
HV

4 day delta

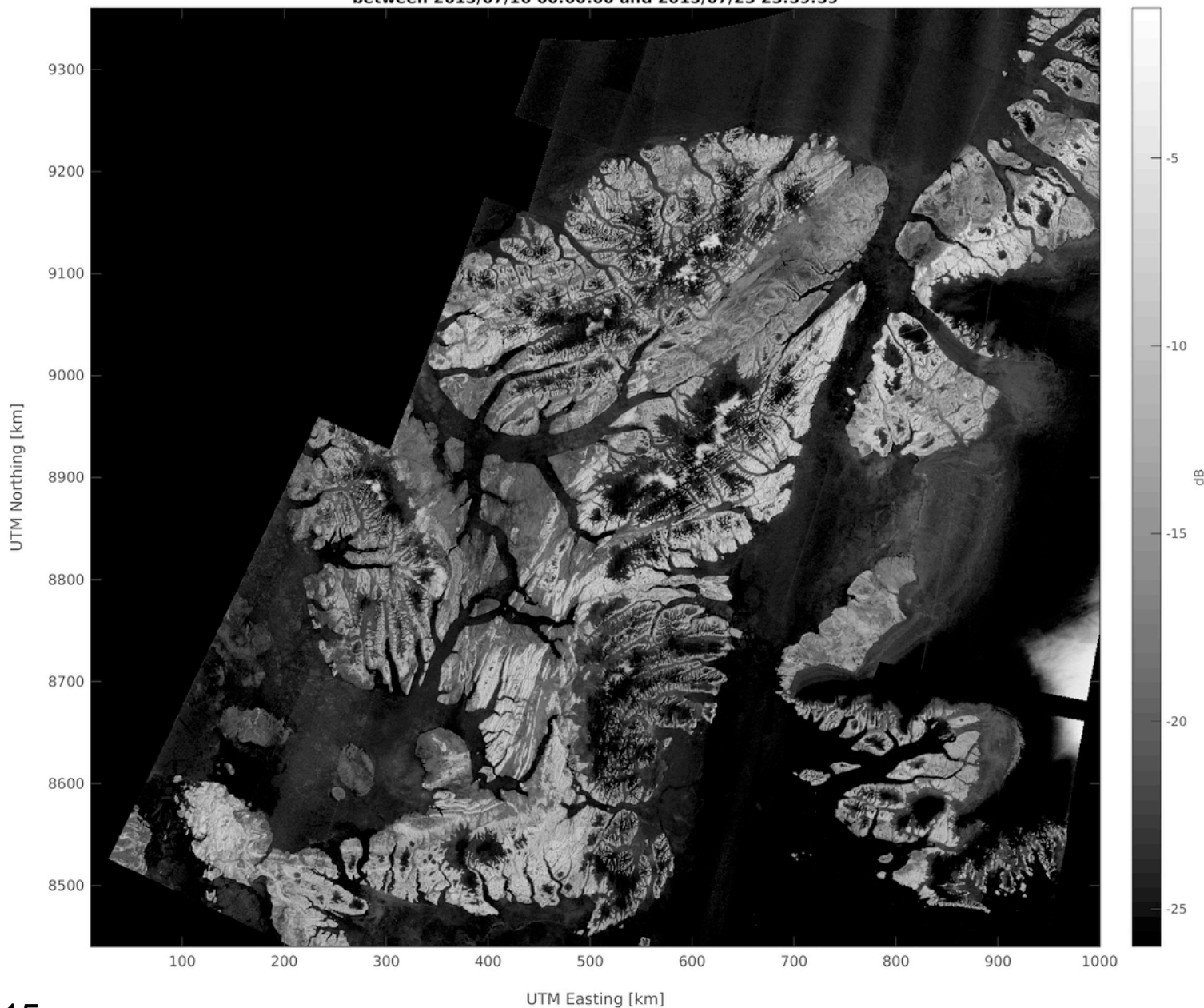
8 day window

N.B.

8 bit radiometry

CDEM

Composite backscatter from 10 scenes  
between 2015/07/16 00:00:00 and 2015/07/23 23:59:59



May – Aug. 2015

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# Ellesmere Island Backscatter Composites

**S1A EW HV**

2 day delta

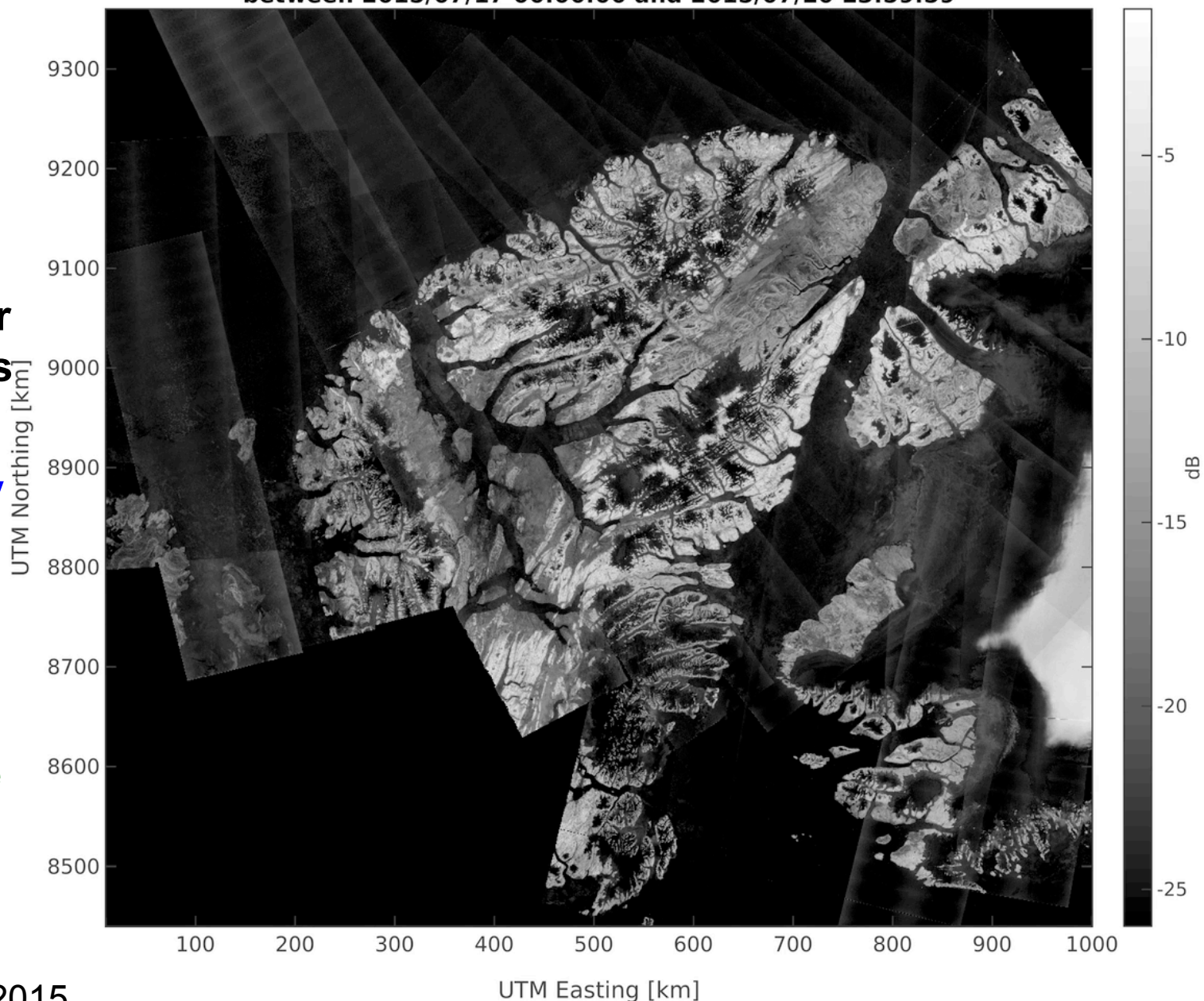
4 day window

N.B.

HH also available

CDEM

Composite backscatter from 21 scenes  
between 2015/07/17 00:00:00 and 2015/07/20 23:59:59



May – Sept. 2015

Contains modified Copernicus Sentinel data (2015)





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## Ellesmere Island, Canada

Sentinel-1 EW 4d  
Composite HV:

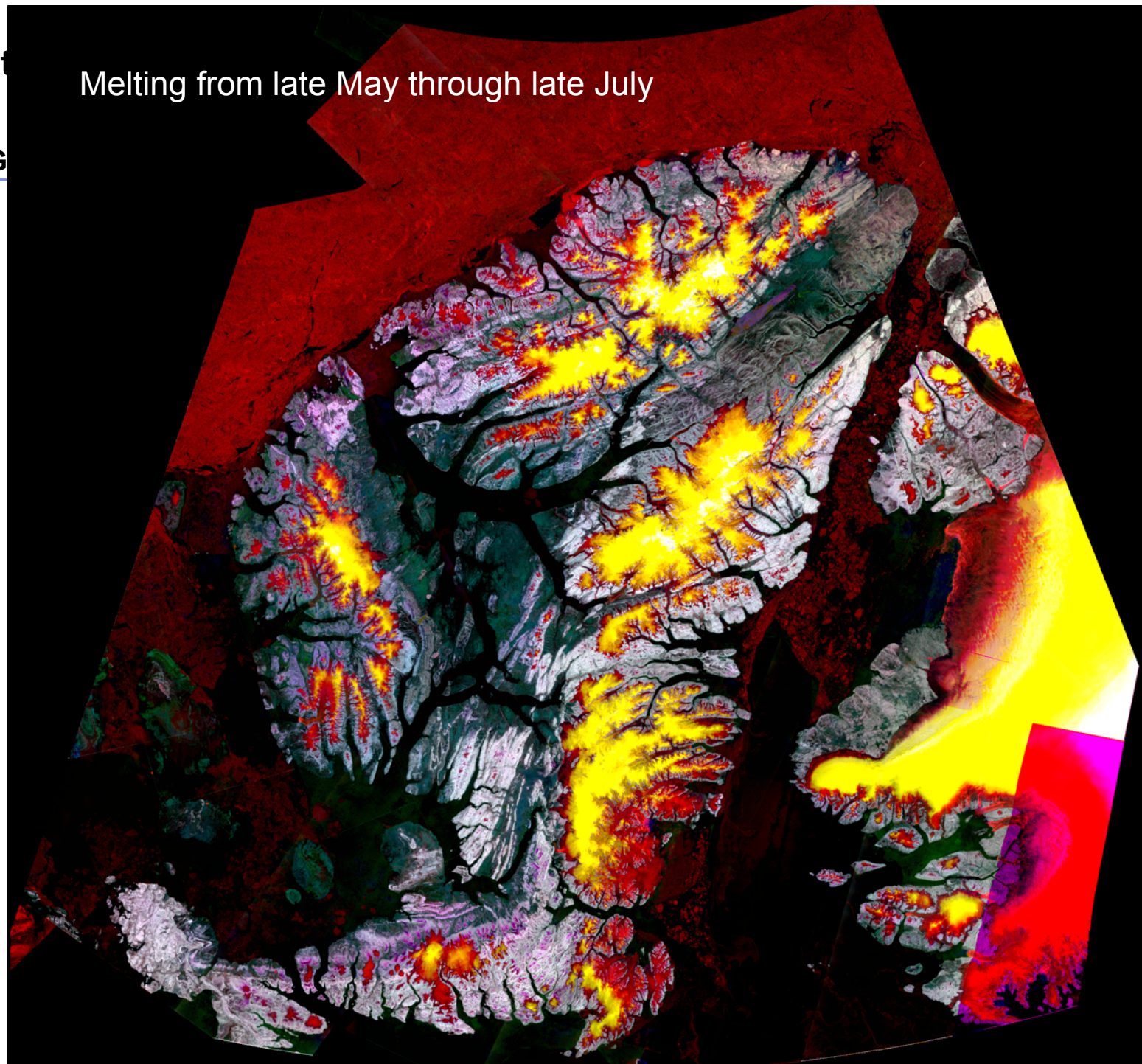
May 24-27,  
June 29-July 2,  
July 23-26

$\gamma_T^0$  HV-pol.

-23dB -6dB

Contains modified  
Copernicus  
Sentinel data (2015)

Melting from late May through late July







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## Ellesmere Island, Canada

Sentinel-1 EW 4d  
Composite HV:

July 25-28,

Aug. 12-15,

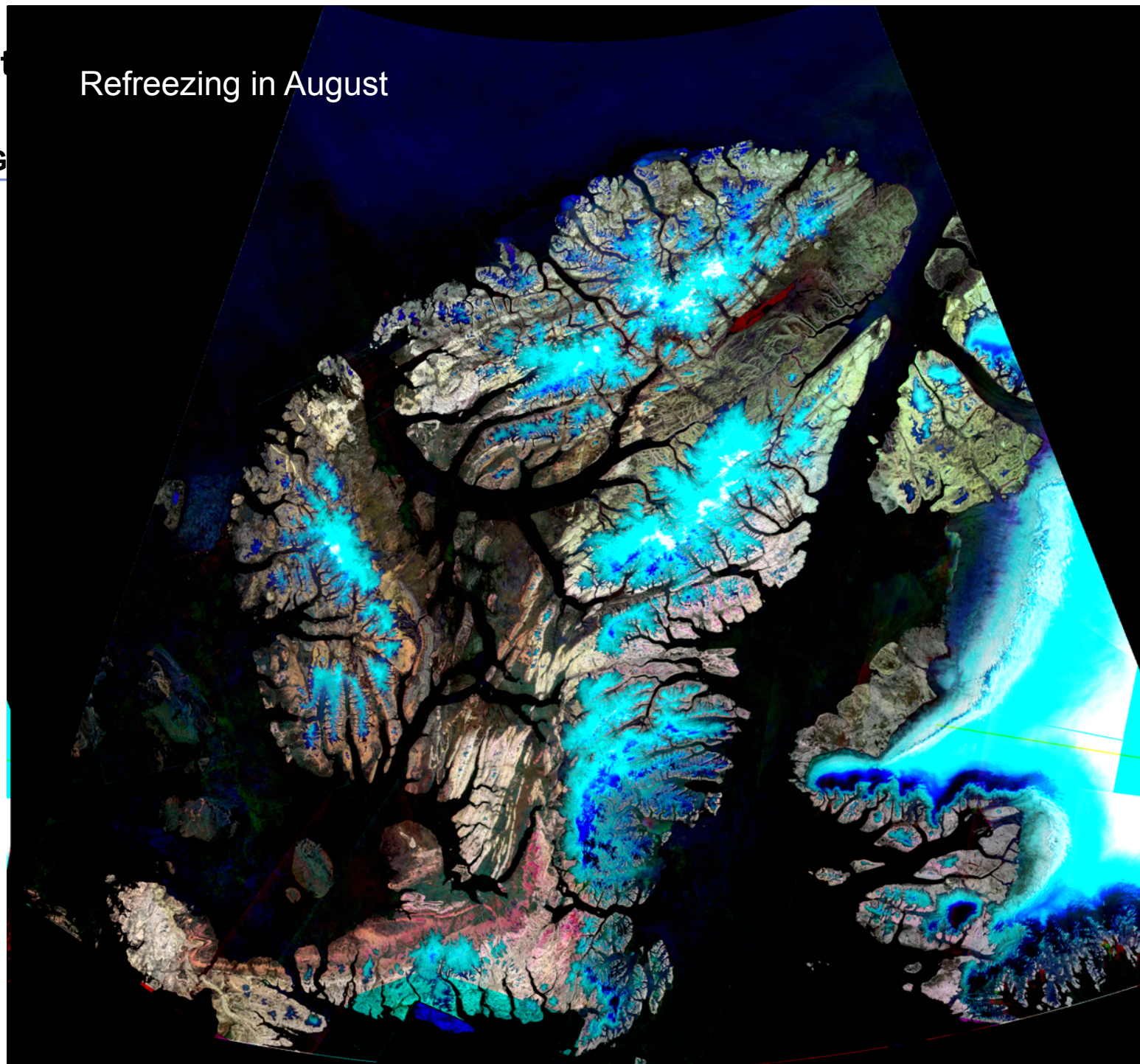
Aug. 24-27

$\gamma_T^0$  HV-pol.

-23dB -6dB

Contains modified  
Copernicus  
Sentinel data (2015)

### Refreezing in August





## Backscatter Composites

- **Demonstrations of Local Resolution Weighting with Sentinel-1A & Radarsat-2**
  - Geometric and radiometric effects of topography strongly reduced
  - Backscatter composite product properties more homogenous across product, also in presence of terrain
  
- **Sensor Integration**
  - **Not limited to a single sensor:** Local Resolution Weighting (LRW) useful for integrating multi-track and multi-mode, but also multi-sensor data streams (e.g. S1 + RS2)
  - **Higher time-resolution coming:** Width of time window can be *narrowed* while still supporting full coverage as more data becomes available (S1B opening this month, RCM-1, -2, -3 in 2018?)
  
- **Importance of Calibration**
  - Composite LRW backscatter stable due to dependable and highly accurate S1A geometric and radiometric calibration



## Conclusions

- **Snow wetness clear strong signal in C-band SAR imagery**
- **Snow depth and Snow Water Equivalent (SWE) currently not accessible in single-date C-band SAR data**
  
- **Series of Sentinel-1 satellites opening a new era of multimodal multi-wavelength data integration**
  - Contributions from other data suppliers (NASA/USGS, CSA, JAXA, DLR, ASI) welcome
    - E.g. SARs: Radarsat-2, Radarsat Constellation Mission, TerraSAR-X, Cosmo-Skymed, PAZ
      - Future: Paz, NiSAR, TanDEM-L
    - E.g. VIS/IR: MODIS, Landsat, Sentinel-2, Sentinel-3





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Copernicus  
Sentinel data (2015)

## Acknowledgments

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T. Jonas, WSL-SLF

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Radarsat-2 data via SOAR programme and MURF with Env. Canada