

-The SMOS Mission-

N.REUL

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Oceanography from Space Laboratory





•Why?

- How ?
- •So What ?
- •Aftermath







Salinity of the Ocean: What do we know about it ?



Ifremer What is ocean salinity?

Ocean salinity S= ionic salt concentration in sea water At the sea surface it is referred to as "SSS" (Sea Surface Salinity)

Unity = PSU (Practical Salinity Unit) 1 PSU≈ 1 g/kg.





Chloride (Cl-):19 gSodium (Na+):11 gSulphate (SO4--):3 gMagnesium (Mg++):1.5 gCalcium (Ca++):0,35 gPotassium (K+):0,35 gOthers :0,00.. g



(Mean chemical composition)

Total \approx 35 g/kg

99% of oceanic waters have salinity between 33.1 and 37.2: =>a global variation in salt concentration between 3.31% and 3.72% !





Historical Density of surface observations 1874-2002

Number of Observations by 1° Square



White - N < 10 Blue - 10 < N < 100 Green- 100 < N < 1000 Red - 1000 < N

F. Bingham et al, 2002

1.3 million SSS observations distributed over the global ocean since 125 years:

✓ No data in 27% elementary oceanic $1^{\circ} \times 1^{\circ}$ area, not accounting for arctic zones.

 \checkmark 70% of these surfaces present at most 10 historical observations

 $\checkmark 28\%$ of all observations were sampled in the coastal domain

✓ Up to 1960, there was no more than 10,000 observations/year ⇔ 1 observation per 4°X4° cell

✓ Since 2002, very net increase in the density of measurements (ARGO network)

Global distribution of the SSS

Monthly climatology of the sea surface salinity:



Atlantic Ocean saltier than Pacific and Indian oceans

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✓ Low variability particularly in the Southern seas and North Pacific
✓ But higher variability around large river run off (Amazon
Congo, Yang Tse, Ganga..), largest currents (Gulf Stream, kuroshio,
Agullas, ...) & in the Tropical bands



Salt versus temperature at the Ocean Surface

Temperature

Salinity



Global and seasonnal distributions of surface salinity strongly differ from the surface temperature one. It is because the processes involved and sources responsible for their own variability are different:

>Ocean are heated in the Tropics and lose heat at higher latitutes

 \triangleright Salinity is modified by dilution-concentration processes associated with the fresh water fluxes. The latter result from the balance between precipitation, evaporation, ice melting/pounding and river run off.

Why Measuring Ocean Salinity From Space ?





Salinity S & Temperature T are indicators of the water masses density ρ :

 $\rho_{sw}(S,T) = \rho_{fw}(T) + b(T)S + c(T)S^{3/2} + dS^2$

« State equation »

Similar to temperature and humidity for the atmosphere





The Thermo-Haline Circulation

Idealized global thermohaline circulation (~1000 years)



- → Warm surface currents
- Deep cold currents

The higher salinity in the Atlantic sustains the oceanic deep overturning circulation



- Conveyor belt . Return period ~1000 years.
- Density differences
- Global scale circulation

Oceanic Fronts Monitoring

SSS SMOS Nov



Equatorial Warm Pool Edge



Rodier et al. (JPO, 2000)

The Oceanographer's Water Cycle

Global Water Cycle





Large River run offs

Geotraces West Atlantic cruise leg 2(RV Pelagia)



- 86% of evaporation over the ocean
- 78% of precipitations over the ocean
- => Ocean is a main component of the earth water cycle
- Sea surface salinity is a tracer of the fresh water flux:





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Surface salinity distributions are closely tied to E-P patterns



Evaporation increases in the Sub-tropical areas and so does the surface salinity

Where precipitation dominates, surface salinity decreases. (Equatorial convection zone & mid-latitudes



Salinity wake behind hurricanes





Heavy rainfall

=>

Fresh water lenses ?

Very large winds=>mixing And upwellings

Salty or Fresh water response Of the upper ocean in the wakes of Tropical Cyclones?

Sea Surface salinity: a climat change indicator

Trends in Sea Surface Salinity in Pacific and Atlantic Oceans

Gordon & Guilivi, Oceanography, 2008. 38.0 37.6 40°N 37.2 36.8 30°N 36.4 36.0 35.6 20°N 35.2 34.8 10°N 34.4 34.0 33.6 **0°** 33.2 120°E 140°E 160°E 180°W 140°W 120°W 100°W 60°W 40°W 20°W **0°** 160°W 80°W Pacific / Atlantic interactions ~+0.03 psu/yea 37.5 increased convection poleward 37.0 Atlantic SSS of high salinities 35.5 ~-0.03 psu/yea salinities warm 35.0 Pacific 34.5 -1950 1955 1960 1965 1970 1975 1980 1985 1990 1995 2000 Years 1950 1007 2000

Near-Surface Salinity as Nature's Rain Gauge to Detect influence of Climate Changes on the Tropical Water Cycle

Trends in the observed SSS from in situ data over 1970-2002 (PSU/century)



SSS changes over the past decades exhibit a strong Pacific freshening and Atlantic salinity increase leading to a strengthening of the mean SSS interbasin contrast, which reflects to a large extent the mean pattern of freshwater fluxes.

=>We observe a recent increase in the marine tropical hydrological cycle strength (*Terray et al, 2012*).

Trends in SSS in the Antarctic Ocean





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10 sections /year since 1993







(Morrow et al., 2008)

Trends in Sea Surface Salinity in the Western Tropical Pacific Warm Pool

The surface extension of the Warm Pool (equivalent to Europe Area with temperatures>28°C) is associated with a surface salinity freshening



~-0.013 psu/year

Cravatte et al., 2009

Temperature

0



Salinity





The spicy waters that change El Niño



0.02

-0.02

90°W

Eastern Pacific El Niño Central Pacific El Niño 20"N 10°N SST Latitude 0 10"5 2015 150°E 180 " 150°W 120°W 90"W 1.50°E Longitude SSS Cluster 3 (44/379 maps)

SSS Cluster 4 (132/379 maps) 0.2 150°E 180 * 150°W 120°W 90°W Longitude

150°W

Longitude

120°W

180*

Christophe Maes, 2002 Singh et al., 2010

90°W

Salinity of the upper ocean play an important role (barrier-layer effects) in the on-set

SSS

of the phenomenon. Monitoring this variable will help in better predicting El Niño.

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1.80 *

150°W

Longitude

120°W

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20°N 10°N

10"5

2018

150°E

Latitude

Surface Salinity & Marine Bio-chemistry

3

Ocean is a major sink for atmospheric CO2 It absorbs ~25% of human emission in the atmosphere, however it is saturated and start



Through its links with carbonate chemistry and surface water masses monitoring, Sea surface salinity data will improve estimates of air-sea CO_2 fluxes. SMOS will help in better quantifying ocean acidification (corals reefs, tropical ecosystem) & ocean-atmosphere CO_2 exchanges in some key areas <u>Surface Salinity and Marine Biolo</u>

Salinity is one of the key environmental factor for the living of fishes and marine biology

Iframa



Why measuring SSS from Space?

•Salinity is a key parameter of ocean dynamics and Climate:

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✓Thermo-haline Circulation

✓Global Water Cycle

(Fresh water flux, E-P-R)

✓ Oocean-atmosphere Coupling

(e.g., ENSO, en rate of CO2 absorbtion)

•Salinity is a key parameter for ocean Biochemistry and Biology

•Lack of SSS measurements

Implications on climatology

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Accuracy ~0.1 psu/monthly

Spatial scale: 100-200 km²



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How can we measure sea surface salinity From Space?

Physics principle Of SSS measurements From Low Microwave frequency radiometry









Basics of SSS measurements from Space

T_b=εT



- At electromagnetic frequency f < 20 GHz, sea water dielectric constant ε is a function of
 - SSS and sea surface temperature SST. $\varepsilon = \varepsilon(SSS,SST)$.

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ε = f(S, T, Freq,
Incidence)
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• The sea surface brightness temperature T_{ant} as measured by a radiometer is thus related to salinity



Given SST(T) & Tb data=> one can deduce SSS (S) in theory



Tb sensitivity to SSS as function of Electromagnetic Frequency

Brightness temperature Sensitivity to Salinity as function of Electromagnetic Frequency



A weakly sensitive principle



The sensitivity of the brightness temperature at L-band to SSS remains small. It depends on the sea surface temperature (SST):

-0.3 K/psu in cold waters (~o°C) à

-0.7 K/psu in warm seas (~30°C)

An istantaneous accuracy on SSS of 0.1 psu would require a radiometer TB measurement accurate to within: ~0.03 K for an SST=0°C! ~0.07 K for an SST=30°C!

The one from AMSR-E & WindSat at 6 GHz: 0.5-0.6 K => technological challenge ! → 4th ESA ADVANCED TRAINING ON OCEAN REMOTE SENSING

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First attempts during NASA Mission SkyLab in 1973



Fig. 1 – Location of radiometers on Skylab

First trial 1973 SkyLab S-194:

L-band radiometer data

Low accuracy & Spatial resolution

Not enough data

Technical limitations due to the size of antennas at L band









Antenna size: the limiting technological factor for SSS remote sensing developement until the 1990s

To obtain a spatial resolution on ground of 50 km (at nadir), from an altitude of 750 km and for an electromagntic wavelength of 20 cm (f=1.4 GHz), a real aperture radiometer must have a characteristic antenna size of ~4 m



Technological Evolution associated with Antenna

SMOS (Soil Moisture & Ocean Salinity) Launch date: November 2^{sd}, 2009

L band radiometer required: No existing device

How to by-pass the antenna size technical difficulty ?: Antenna deployed in space and Interferometry







Goal of both missions:

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• SSS measurments with an accuracy of 0.1-0.2 psu and a spatial resolution of 100x100 km every 10 days (GODAE requirements).

Soil Moisture and Ocean Salinity

The sensor: L band interferometric synthetic aperture Radiometer (1.4 GHz)



a) SMOS artist view

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b) SMOS swath

Brightness temperature measurements at different incidence angles $(0^{\circ} - 60^{\circ})$ Ground resolution: 35-80 km Global coverage every 3 days Spot accuracy (instantaneous) ~1 psu

Aquarius/SAC-D

L band (1.4 GHz) radiometer with 3 incidence angles + L band scatterometer



a) Aquarius artist view

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b) Swath

Spatial resolution: 100 km every 10 days. Accuracy ~0.5 psu





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THE FIELD OF VIEW AT ANTENNA LEVEL

The earth boundary passes outside fundamental hexagon (magenta) and thus earth aliases appear inside the fundamental hexagon.





THE FIELD OF VIEW PROJECTED ONTO THE EARTH

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Here is an example of the scene brightness model for one coastal snapshot. Direct celestial sky brightness has been multiplied by ten in the plot to show it more clearly, and land brightness temperature has been set to 280 K arbitrarily.





Fig. 1 – SMOS field of view (solid line). The red ellipses represent 3 dB synthetic antenna pattern contours (Blackman exact apodization window). The 40-km and 90-km major axis limits appear in green. The blue circles represent the locations of the incidence angles 10° , 25° , 40° and 55° . One dwell line is also shown (dashed line).

⁰¹² Zine et al. 2007





CONCEPT OF A DWELL LINE

THE CONCEPT OF A DWELL LINE





















Ifremer THE MultiAngular SMOS Aquisition



Ifrem SMOS orbital characteristics

Morning ascending passes Over terminator to minimise sun contamination

Parameter	Mean Value
Semi-major axis	a = 7134.552 km
Eccentricity	e = 0.00116
Inclination (sun-synchronous)	i = 98.445°
Argument of perigee	ω = 90°
Mean Local Solar Time	Ω = 06:00 AM
Repeat cycle / cycle length	149 days, 2144 orbits
Orbital duration	6004.478 s



1 day of SMOS data over the ocean



How in Practice ?

Brightness Temperature





Tb^{meas} corrected for OTT

Tb^{mod} is computed at antenna level (Tx, Ty) in order to avoid interpolation for getting same incidence angle.

A change of state in the Hydrogen atom energy generates micro-wave electromagnetic radiations at a frequency of 1420 MHz (L band) ≡length 21 cm known as the « Hydrogen line »

21 cm

Hydrogen being one of the first constituent of the sun And of most of the stars, Earth is constantly illuminated by L-band radiations



Sun

APPARENT TEMPERATURE IONOSPHERE

ATMOSPHERE

IONOSPHERE

ATMOSPHERE

OCEANIC SURFACE

TEMPERATURESALINITY ROUGHNESS



What can SMOS achieve in salinity observations? A technological and scientific challenge



STO





SMOS SSS from one ascending and one descending semi-orbit on 24 Jul 2012 after removing flagged data (http:// www.argans.co.uk/smos/pages/ products.php)

Along track Level 2 SSS qualitatively OK but very noisy.

Average needed to check mission requirements (Level 3)

SMOS Level 2 SSS products





Passes ascendantes



SMOS simple averaging L3 products

10 days –ascending passes

10 days –descending passes



SMOS refined Level 3 SSS products

SSS 10-Day Composite from May 01 through May 10-2010-1°x'



CATDS research CEC products

The SMOS products used: Centre Aval de Traitement des Données SMOS (Level 3)





CNES/CESBIO/IFREMER French ground segment For SMOS L3 & L4

http://www.catds.ifremer.fr/



	CEC-IFREMER	CEC-LOCEAN
SSS retrieval method	SSS retrieved from first Stokes parameter (Reul and Tenerelli 2011)	SSS retrieved from polarized Tbs along dwell lines using an iterative retrieval (see ESA L2OS ATBD)
Region of the instrument field of view (FOV) considered for SSS retrieval	Alias free field of view only	Alias free field of view (AFFOV) and extended AFFOV along dwell lines with at least 130 Tb data samples in AFFOV ($\sim \pm 300$ km from the swath center)
Tb filtering method	Determined from interorbit consistency in incidence angles classes and thresholding	Determined from consistency along dwell lines as reported in ESA level 2 products
Galactic model	Geometrical optics model	Kirchoff's approx. scattering at 3 m/s
Roughness/foam models	Empirical adjustment of Tb dependencies to wind speed	Empirical adjustment of parameters in roughness model and foam coverage models (Yin et al. 2012)
Calibration	Single ocean target transformation (OTT) + daily $5^{\circ} \times 5^{\circ}$ adjustment wrt World Ocean 2001 SSS climatology	Variable OTT (every 2 weeks synchronized with noise injection radiometer as defined in ESA reprocessing)
Average	Simple average	Average weighted by theoretical error on retrieved SSS and spatial

Table 1 Summary of characteristics of CATDS-CEC SSS level 3 products

resolution



SMOS-CATDS-CEC Level 3 product (see http://www.catds.fr/): Monthly Composite





IFREMER-CEC Stronger RFI filtering than ESA L2 Strong constraints wrt SSS climatology

LOCEAN (ESA L2 binned SMOS SSS)

So, several SSS products exists but needed because none of them is perfect and parallel efforts & progresses are required

SMOS Level 3 product: 10 days / 1º optimally interpolated ocean salinity map for 15 – 24 January 2012

Sea Surface Salinity

 $1^{o} \times 1^{o}$ Optimal interpolated map - 15/24 January, 2012 - BEC product



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Ifremer SMOS-ESA Documentation

- Newsletter (every ~2months)
- Available on https://earth.esa.int/web/guest/missions/esa-operational-eo-missions/ smos/newsletter
- =>Highlights (RFIs, new results etc...)
- =>Data availability (anomalies + calibration) (see also: https://earth.esa.int/web/guest/missions/esa-operational-eo-missions/smos/available-data-processing)
- =>Upcoming meetings
- =>Data access
- L2OS release note (when new processed data delivered)
- => Recommendations about the use of released data (quality/defaults)
- Release note for v62x salinity data available on:
- https://earth.esa.int/documents/10174/1854503/SMOS_L2OSv622_release_note
- Access to SMOS data
- See https://earth.esa.int/web/guest/-/how-to-obtain-data-7329





Other Informations:

http://www.salinityremotesensing.ifremer.fr/ http//www.cesbio.ups-tlse.fr/us/indexsmos.html http://www.argans.co.uk/smos http://www.locean-ipsl.upmc.fr/smos

Special SMOS issue in IEEE TGRS, 2012 Special SMOS-AQUARIUS issue JGR 2014